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## Atmospheric CO<sub>2</sub> Variations in Two Contrasting Environmental Sites Over India

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**ABSTRACT:** We analyzed the influence of environmental parameters on the temporal variation of atmospheric carbon dioxide (CO<sub>2</sub>) mixing ratios in two environmentally contrasting Indian sites, Dehradun (30.1°N, 77.4°E, humid subtropical station) and Gadanki (13.5°N, 79.18°E, dry tropical station), from October 2010 to September 2011. The annual range of mixing ratios is low in Gadanki as compared to those of Dehradun because of relatively less monthly variation in temperature and relative humidity (RH) at Gadanki. At both the stations, the minimum mixing ratios are present during the high ecosystem productivity seasons in the afternoon hours. The maximum values are in the early morning hours. However, low wind speed conditions control the unexpected afternoon high mixing ratios in Gadanki during the pre-monsoon season. The early morning maximum is high during monsoon and post-monsoon seasons in Dehradun and Gadanki, respectively, whereas morning inflexion occurred earlier in Gadanki compared with Dehradun. The effect of cloudiness on the CO<sub>2</sub> uptake depends on the canopy cover.

**KEYWORDS:** atmospheric carbon dioxide, water vapour correction, diurnal variation, cloudiness impact

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### 1. Introduction

Carbon dioxide (CO<sub>2</sub>) is the most important anthropogenic greenhouse gas present in the atmosphere. This gas contributes about 64% of the total radiative force created by the long-lived greenhouse gases.<sup>1</sup> CO<sub>2</sub> concentrations have increased by 40% since pre-industrial times, primarily from fossil fuel emissions and secondarily from net land use change emissions.<sup>2</sup> Atmospheric CO<sub>2</sub> measurements are very important tools for understanding the carbon cycle because CO<sub>2</sub> mixing ratios in the atmosphere are strongly affected by photosynthesis, respiration, oxidation of organic matter, biomass and fossil fuel burning, and air–sea exchange process.<sup>3</sup> Local temperature and moisture conditions also affect the diurnal cycle and seasonal variations in atmospheric CO<sub>2</sub>. CO<sub>2</sub> mixing ratio measurements from World Meteorological Organisation (WMO) Global Atmosphere Watch (GAW) program are important to understand the fluxes

using atmospheric inversion models,<sup>4</sup> whereas flux measurements using fast response sensors in eddy covariance method are critical for monitoring global and regional CO<sub>2</sub> mixing ratios.<sup>5</sup> Near-surface CO<sub>2</sub> mixing ratios have been documented in several cities across the world (Vancouver, Canada; Kuwait City, Kuwait; Mexico City, Mexico; Basel, Switzerland; Nottingham, UK; Phoenix, USA) to evaluate the dynamics of atmospheric CO<sub>2</sub> over short periods of time.<sup>6,7</sup>

In India, long-term analysis of atmospheric CO<sub>2</sub> using the flask sample measurement technique at Cape Rama is given in Refs.<sup>8–10</sup> The continuous measurements of CO<sub>2</sub> at Hanle and Pondicherry are carried out by Center for Mathematical Modeling and Computer Simulation (C-MMACS) group Bangalore.<sup>11</sup> Flux measurements over agricultural site using Moderate Resolution Imaging Spectroradiometer (MODIS)-based vegetation indices and land surface temperature

information are carried out by Patel et al.<sup>12</sup> whereas analysis of CO<sub>2</sub>, water vapor, and energy fluxes over mixed deciduous forest is reported by Jha et al.<sup>13</sup>

The large diversity of climatic resources, like vegetation; land use changes; densely populated regions with high anthropogenic sources; and energy hot spots of CO<sub>2</sub> emission (thermal plants, fossil fuel combustion) results in CO<sub>2</sub> source and sink spatial and temporal variability. Such variability has not been characterized in developing an understanding of net carbon balance of the country.

As a part of the program of Indian Space Research Organisation (ISRO)—Geosphere Biosphere Programme (GBP)—an integrated study of the carbon cycle has been initiated. These CO<sub>2</sub> mixing ratio measurements at multisites will be integrated with forest and soil carbon pool for quantifying national carbon balance. Recently, Sharma et al.<sup>14</sup> reported the influence of the environmental factors and biospheric processes on the temporal variations of atmospheric CO<sub>2</sub> in Dehradun. Here, we report an analysis of continuous measurements of CO<sub>2</sub> mixing ratios at two environmental contrasting sites of India, which are separated by ~17° in latitude. One is Dehradun (30.1°N, 77.4°E) in Uttarakhand with a subtropical humid climate, and the other is Gadanki (13.5°N, 79.18°E) in Andhra Pradesh with a tropical dry climate.

The structure of the paper is organized as follows: Section 2 introduces the site and instrument used for the measurements. Sections 3 and 4 describe the data and methodology applied

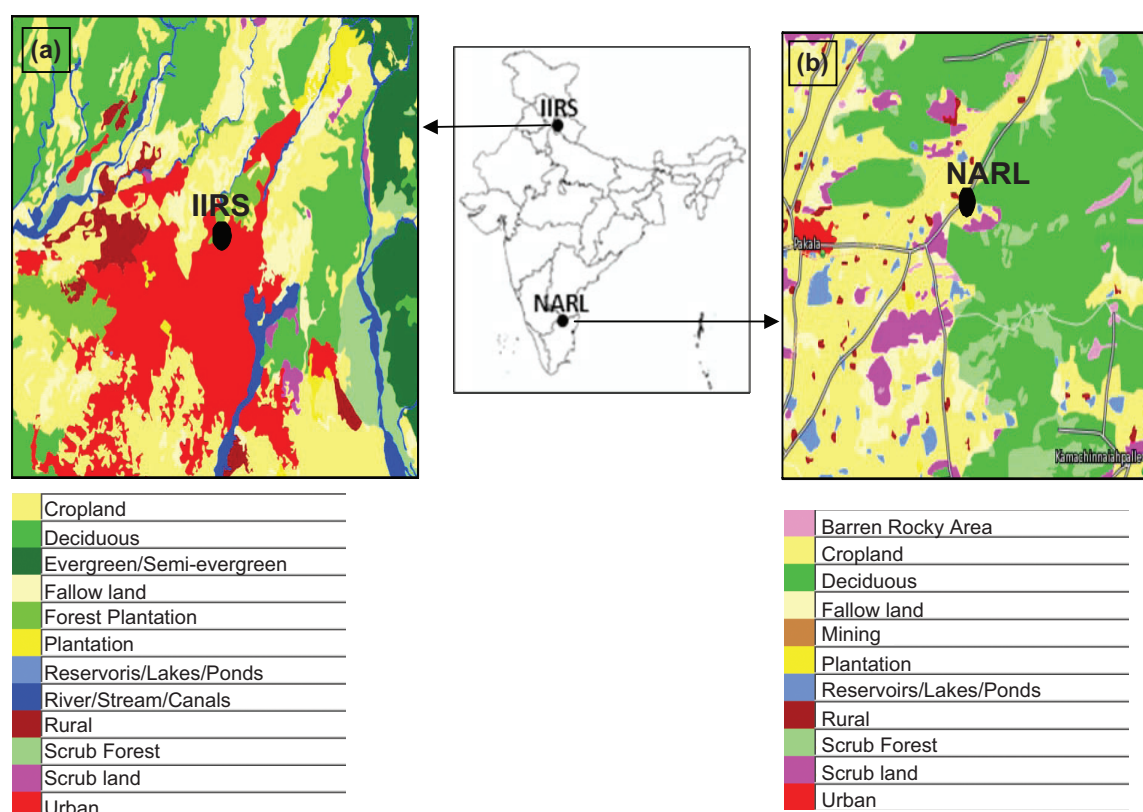
to the analysis, respectively. Impact of environmental parameters is discussed in Section 5 and summary of the results in Section 6.

## 2. Instrument and Site

**2.1. Instrument.** CO<sub>2</sub> observations at both the stations are recorded using a Vaisala GMP-343 probe through a data logger.<sup>15</sup> The instrument is of diffusion type and based on non-dispersive infrared (NDIR) technology (detail working principle of the sensor is given by Rigby et al.)<sup>16</sup> It provides an acceptable compromise between size, response time, accuracy, and stability.<sup>14,16</sup> Studies show reliable results using this sensor in atmospheric, limnology, and oceanographic investigations.<sup>14,16–18</sup>

**2.2. Site.** Dehradun is in a valley located at a mean altitude of 700 m with a humid subtropical climate. This city has a large area covered under vegetation and trees. On the other hand, Gadanki is a tropical semi-arid station located near Tirupati in the Chittoor district of Andhra Pradesh. It is a complex hilly terrain surrounded by dense forest and a very irregular mix of agriculture, and has small-scale rural population.

At both the locations, the sensors are mounted on the top of buildings, sufficiently above the roof to avoid heating effects of the building. The building height is approximately 15 m from the ground. One instrument is in the campus of Indian Institute of Remote Sensing (IIRS), Dehradun, and the other at the National Atmospheric Research Laboratory (NARL), Gadanki. The location of the sites is shown in Figure 1. Both



**Figure 1.** Location of observation sites: (a) IIRS, Dehradun and (b) NARL, Gadanki.

the institutes are significantly away from the main traffic, and hence the impact of public vehicular pollution is expected to be negligible on CO<sub>2</sub> measurements.

**2.3. Meteorology of the stations.** The mean monthly variation of temperature (°C) and RH (%) in both the stations are shown in Figure 2(a) and (b), respectively. Dehradun is a cold city with air temperature less than 28°C throughout the year. The minimum temperature of 10.4°C is observed during January and a maximum of 27.9°C during May. Gadanki, being a hot station, the temperature is above 20°C throughout the year with a minimum temperature of 21.3°C during December and a maximum of 30.7°C during May. The annual temperature range in Dehradun is 17.5°C, whereas it is 9.4°C in Gadanki.

Relative humidity (RH) in Dehradun crosses more than 80% during July–September, whereas in Gadanki it crosses this value only in November. From pre-monsoon to monsoon seasons, a significant increase in RH is observed in Dehradun, and it remains high until October. On the contrary, the increase in RH during monsoon season is not significant in rest of the year in Gadanki.

### 3. Data

CO<sub>2</sub> mixing ratios were recorded at 15 minutes interval, at both the stations, during October 2010–September 2011. Temperature (°C), RH (%), wind speed (m/second), and cloud amount (in tens) are obtained from

- India Meteorological Department (IMD), Dehradun, at three-hour intervals. These data (except the cloud amount) are linearly interpolated to hourly values. The cloud amount observations are at 0730 and 1730 hours. If both the observations are zero, we considered the day to be cloud free, and if both the observations are greater than or equal to nine, we considered the day as cloudy (C).

- Automatic Weather Station (AWS), Gadanki, at one-hour interval.

### 4. Methodology

At both the sites, sensors are installed after a factory calibration by Vaisala at 25°C and 1,013 hPa. The two instruments are installed at different years. Hence, they have been inter-calibrated with a known concentration by adding biases to the measurements. The bias correction was applied linearly from the date of installation to the calibration date. As Dehradun is located at 700 m above sea level, CO<sub>2</sub> measurements are corrected for the ambient temperature and pressure using an ideal gas equation.<sup>15</sup> Sharma et al<sup>14</sup> have given the detailed procedure for this correction. No correction has been applied to the observations at Gadanki, as its height is 375 m above the mean sea level. However, CO<sub>2</sub> observations were corrected for atmospheric moisture, at both the stations, using the following equation, which relates diluted and dry gas mole fraction<sup>19</sup>:

$$\frac{C_{\text{dilution}}}{C_{\text{dry}}} = 1 - 0.01 H_{\text{act}} \quad (1)$$

where  $C$  is the mole fraction of CO<sub>2</sub> and  $H_{\text{act}}$  is the actual water mole fraction (in percentage).

To know  $H_{\text{act}}$  in the atmosphere, we first calculated  $H_{\text{saturated}}$  using the following equation<sup>20</sup>:

$$\ln\left(\frac{p}{p_c}\right) = \frac{(a_1\tau + a_2\tau^{1.5} + a_3\tau^3 + a_4\tau^{3.5} + a_5\tau^4 + a_6\tau^{7.5})T_c}{T} \quad (2)$$

where  $p = H_{\text{saturated}}$  (saturated vapor pressure),  $T_c$  (critical temperature) = 647.096 K,  $p_c$  (critical pressure) = 22.064 MPa,  $a_1 = -7.859\ 51783$ ,  $a_2 = 1.844\ 082\ 59$ ,  $a_3 = -11.786\ 6497$ ,

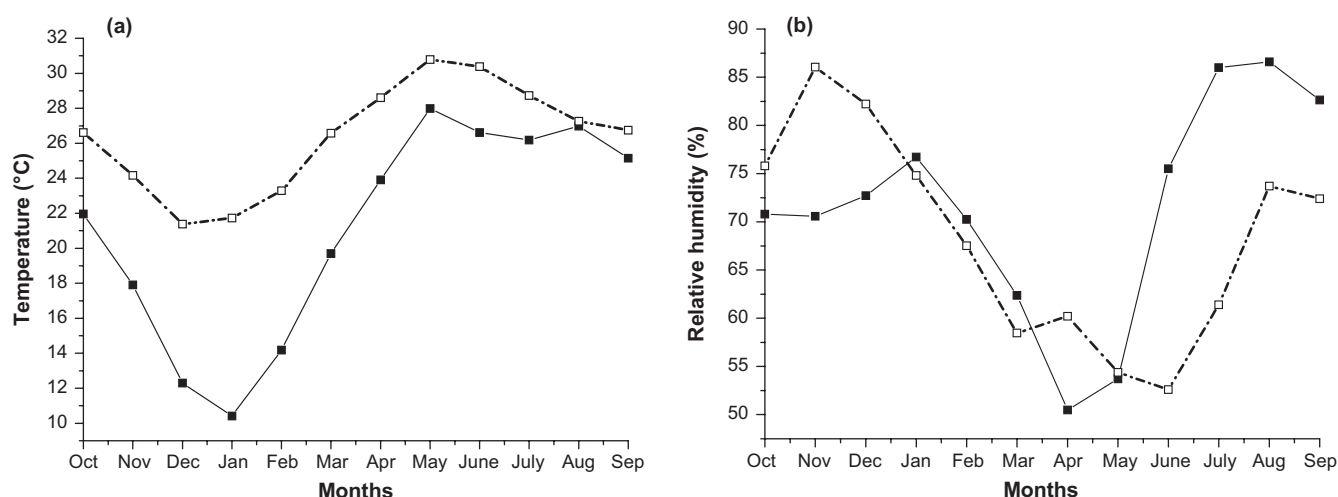


Figure 2. Monthly variations in (a) temperature (°C) and (b) RH (%) in Dehradun (solid line) and Gadanki (dash line).

$a_4 = 22.680\ 7411$ ,  $a_5 = -15.961\ 8719$ ,  $a_6 = 1.801\ 225\ 02$ , and  $\tau = 1 - (T + 273.15)/T_c$ .

With known RH and calculated  $H_{\text{saturated}}$ ,  $H_{\text{act}}$  is calculated using the following equation:

$$\text{RH} = \left( \frac{H_{\text{act}}}{H_{\text{saturated}}} \right) \times 100 \quad (3)$$

Then average hourly mixing ratios were calculated using the 15-minute observations at both the stations from which monthly means were computed. Similarly, monthly means of meteorological observations are also computed.

## 5. Results

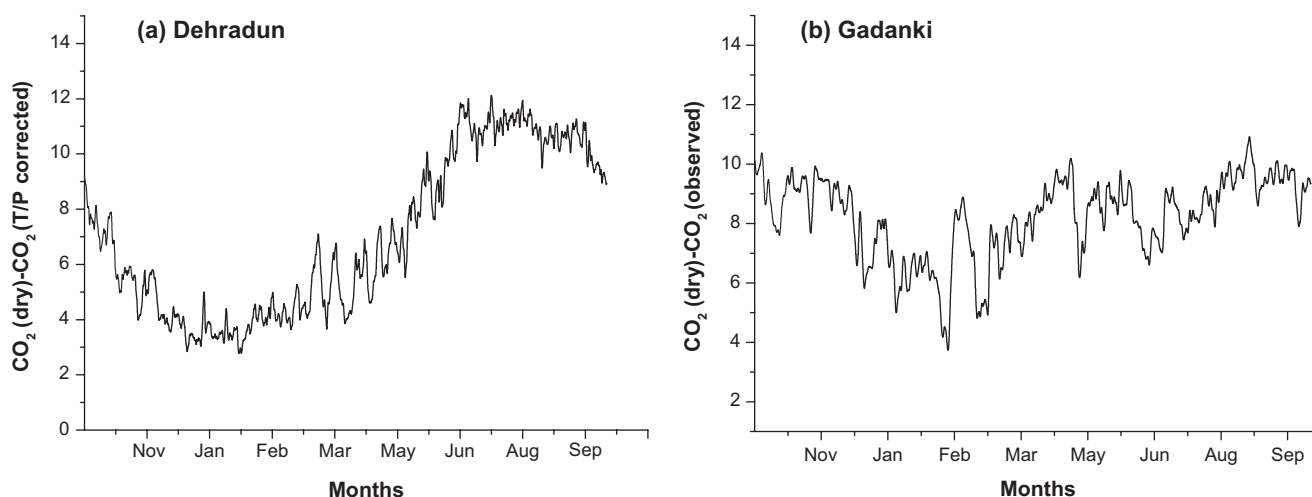
Impact of water vapor correction on the dry concentrations of  $\text{CO}_2$  at the two environmentally contrasting stations is discussed here. Diurnal variation of  $\text{CO}_2$  with respect to the meteorological parameters such as wind speed and cloudiness is also discussed.

**5.1. Water vapor corrections.** The differences between dry (water vapor corrected) and wet (without water vapor correction)  $\text{CO}_2$  mixing ratios in Dehradun and Gadanki are shown in Figure 3. The cycle is smoothed by applying adjacent point averaging technique. The dilution of mixing ratios depends upon the amount of water vapor and temperature conditions of the atmosphere. The monthly variations of difference between dry and observed values are different in both the stations because of the differences in the atmospheric water vapor. In case of Dehradun, the pattern of this difference follows almost that of humidity, indicating that in a wet station, the correction is mainly controlled by the humidity. On the other hand, such a correspondence with humidity is not observed in Gadanki, which shows that other factors like temperature also influence in the dry station.

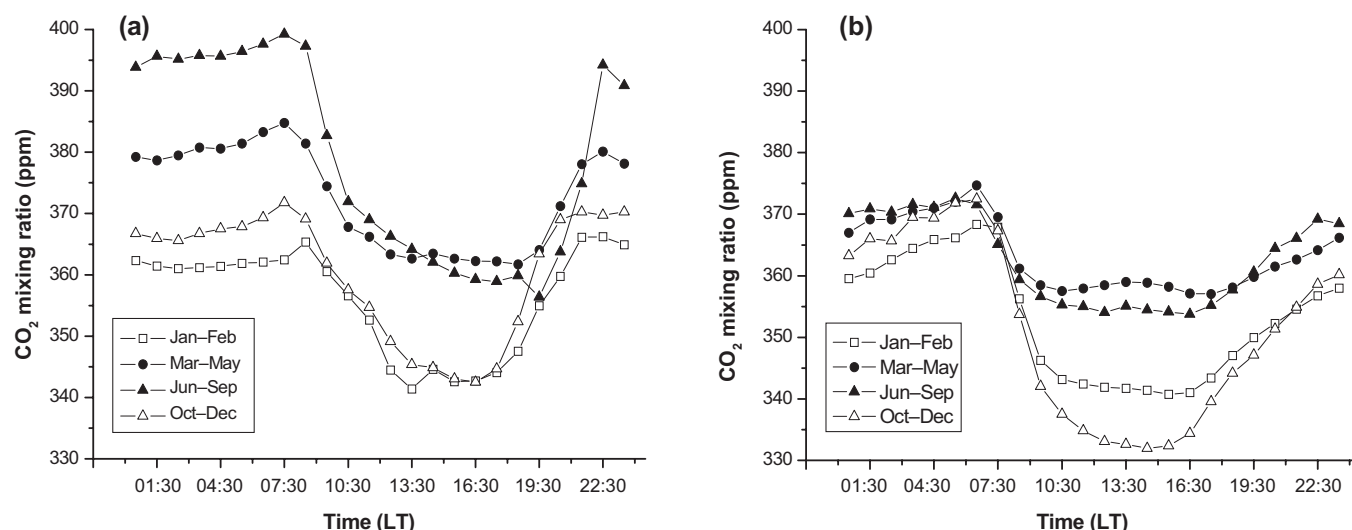
The mixing ratios have increased by 2–14 ppm in both the stations after applying the water vapor correction. As Dehradun has more water vapor in the atmosphere in the monsoon season, this difference is more in this season compared with winter and post-monsoon seasons (Fig. 3). On the other hand, the monthly variations in water vapor content are less in Gadanki (a dry station). Thus, the difference between dry and observed mixing ratios is not much during the study period, although relatively less difference is observed in winter and more during pre-monsoon and monsoon seasons.

**5.2. Diurnal cycles.** The diurnal cycles of  $\text{CO}_2$ , mixing ratios in Dehradun and Gadanki for different seasons, are shown in Figure 4(a) and (b), respectively. Both the stations exhibit similar cycles of mixing ratios with a morning maximum followed by an afternoon minimum. The values again increase from evening onward reaching a maximum in the early morning hours. Mixing ratios sharply decrease after sunrise because of atmospheric process and drops to a maximum in the afternoon because of the net ecosystem uptake and increasing boundary layer. During night, mixing ratios increase because of the absence of photosynthetic activity, and stable atmospheric boundary layer. These types of diurnal cycles are also observed in other stations.<sup>11,21</sup>

At both the stations, the minimum (maximum) mixing ratios are present during the seasons with high (low) ecosystem productivity. Monsoon and post-monsoon seasons have high ecosystem productivity, and winter and pre-monsoon seasons have low productivity. The day time uptake of mixing ratios is more during winter and post-monsoon seasons at both the sites. However, the amplitudes of diurnal cycles during pre-monsoon and monsoon seasons are more or less similar in both the stations. The nighttime increase is more significant in Dehradun especially during pre-monsoon and monsoon seasons because of more moisture content in the soil, whereas such an increase is not significant in Gadanki.



**Figure 3.** Water vapor correction on  $\text{CO}_2$  mixing ratios in (a) Dehradun and (b) Gadanki.



**Figure 4.** Diurnal cycles of mixing ratios (ppm) during different seasons in (a) Dehradun and (b) Gadanki.

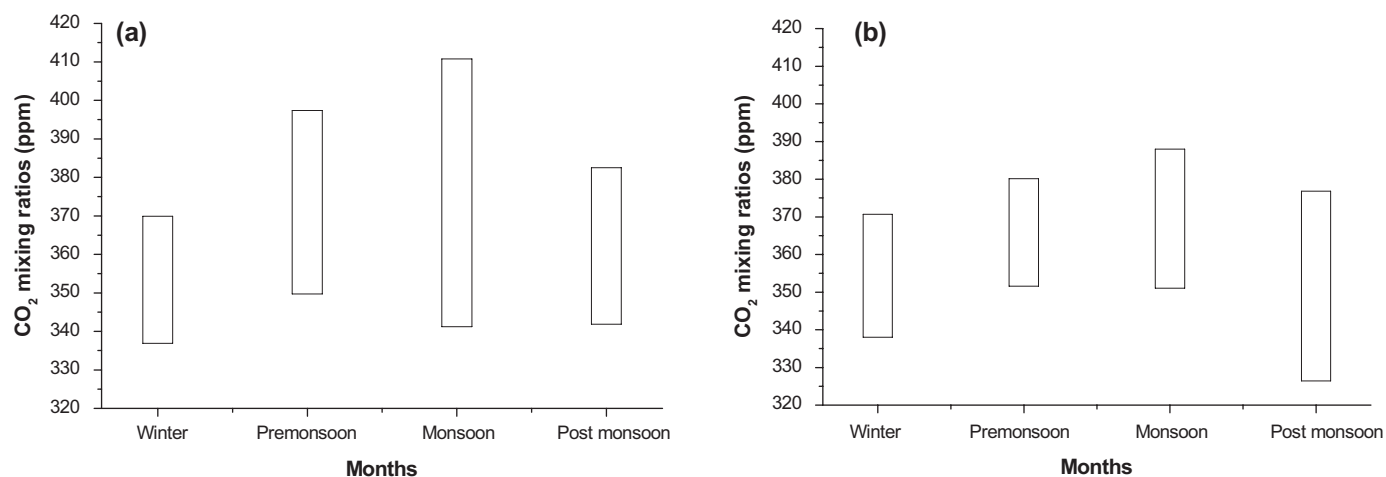
The seasonal variations of minimum and maximum mixing ratios at Dehradun and Gadanki are shown in Figure 5(a) and (b), respectively. Annually, the range of mixing ratios is high in Dehradun as compared to those of Gadanki. The maximum range (~69 ppm) is observed during June–September in Dehradun, whereas it is ~50 ppm during October–December in Gadanki. The minimum and maximum values are nearly similar during January–December at both the stations.

Both the stations exhibit similar diurnal cycles of mixing ratios. However, the main differences are in (a) wind speed influence, (b) early morning inflexion, and (c) impact of cloudiness. The seasonal changes are minimized in Gadanki, a dry station, which can be attributed to the minimum changes in the soil moisture. Similarly, the day and night changes are minimum in this station.

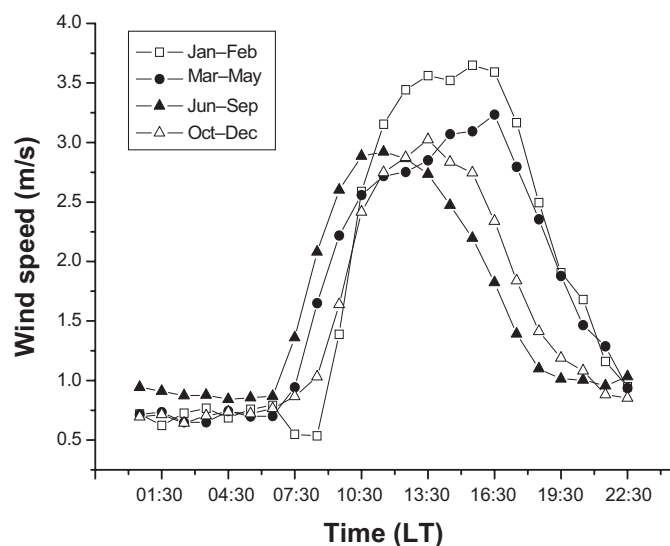
(a) *Wind speed influence.* Generally, mixing ratios during 1130–1530 hours have the lowest values of the day. However,

in Gadanki during pre-monsoon season mixing ratios are not the lowest during this period. Such values are not observed in other seasons in Gadanki and in any winter, monsoon and post monsoon seasons in Dehradun. As wind speed is an important parameter that affects the atmospheric boundary layer mixing, we analyzed the average diurnal cycle of wind speed in different seasons in Gadanki (Fig. 6). The wind speeds during 1130–1330 hours in the pre-monsoon seasons are low as compared to those in other seasons in Gadanki. The afternoon low wind speed reduces the atmospheric mixing because of which CO<sub>2</sub> is poorly mixed during pre-monsoon season resulting in higher values of CO<sub>2</sub>. Thus, wind speed played an important role in the CO<sub>2</sub> concentrations during this time. As these observations are monthly means, the impact of winds on the mixing ratios is significant.

(b) *Early morning inflexion.* CO<sub>2</sub> mixing ratios start increasing from evening onward and reach a maximum before sunrise of the next day from where it starts decreasing.



**Figure 5.** Monthly range (maximum and minimum) of mixing ratios in (a) Dehradun and (b) Gadanki.



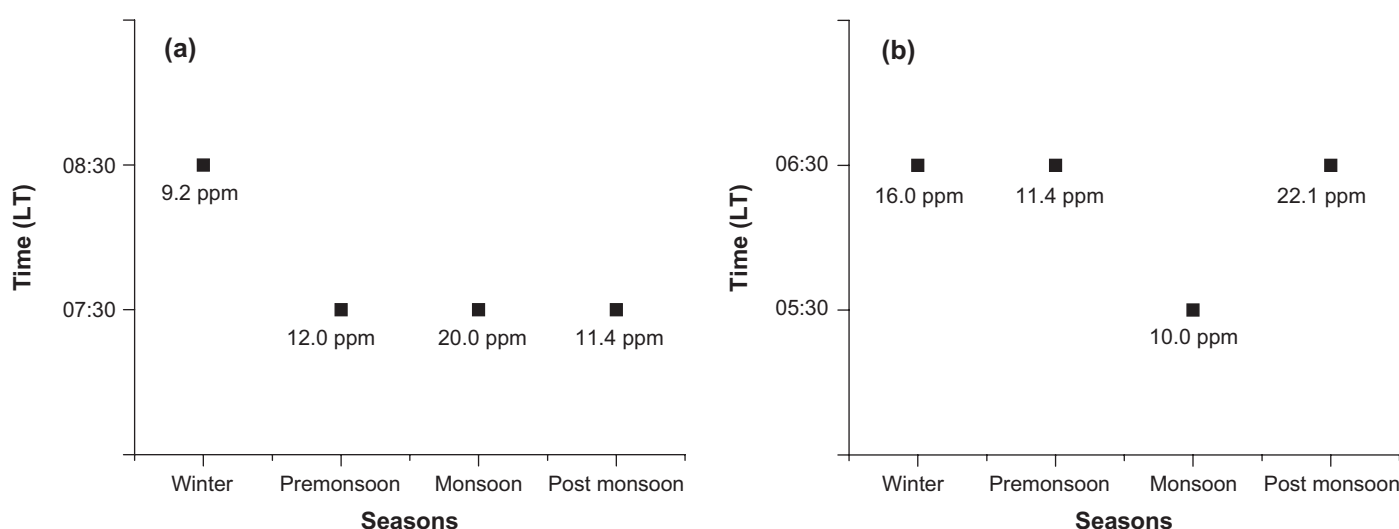
**Figure 6.** Diurnal cycle of wind speeds during different seasons in Gadanki.

The time when it reaches to morning maximum before the decrease is considered as the morning inflexion time. The time of morning inflexion and the value of maximum mixing ratios at that time are also different during different seasons at the two stations. The timings of inflexion during different seasons in Dehradun and Gadanki are shown in Figure 7. The numbers shown in these figures are the amounts of increase in mixing ratios at that time with respect to the seasonal mean values. The morning inflexion is 0730 hours in Dehradun from pre-monsoon to post-monsoon seasons, and it is one hour later (0830 hours) during winter. In Gadanki, the inflexion points are at 0630 hours during winter, pre-monsoon and post-monsoon seasons whereas they are at 0530 hours during monsoon season. The time of morning inflexion depends upon

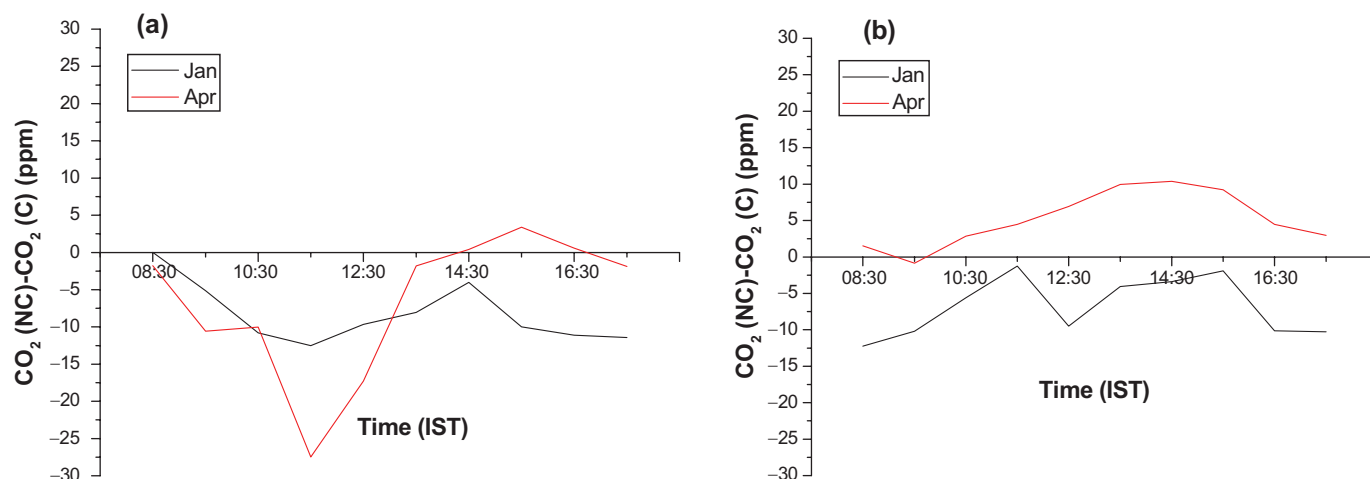
the sunrise time of the station. In Dehradun, being at high latitude, the sunrise time is later compared to Gadanki. The magnitude of maximum mixing ratios at the time of inflexion exhibits a strong seasonality at both the stations. For example, the maximum of 20.0 ppm (22.1 ppm) is observed during the monsoon (post-monsoon) season in Dehradun (Gadanki). The increase in vegetation and soil moisture during the monsoon and the post-monsoon seasons increases the accumulation of  $\text{CO}_2$  mixing ratio from evening to early morning. Gadanki receives more rain during the post-monsoon season compared to those during monsoon. Hence, it shows high morning maximum during this season.

(c) *Impact of cloudiness.* The amount of cloudiness influences the solar radiation reaching the ground, which in turn affects the  $\text{CO}_2$  exchange between the terrestrial ecosystem and the environment. Diffuse radiation increases that of direct radiation when the cloud amount increases. The carbon uptake in the scenario of increased diffuse radiation depends on the vegetation cover. Still et al,<sup>22</sup> Gu et al,<sup>23</sup> and Law et al<sup>24</sup> have shown that increase in diffuse radiation enhances forest canopy photosynthesis. This, in turn, increases the net ecosystem exchange of  $\text{CO}_2$  despite the low irradiance on cloudy (C) days. This effect is driven by a large increase in light-limited shade leaf photosynthesis following an increase in the diffuse fraction of irradiance.<sup>22</sup> On the other hand, photosynthesis reduces in the grass canopy with increasing cloud cover and diffuse fraction because of near-constant light limit of photosynthesis.<sup>22</sup>

Here, we study the impact of cloudiness on  $\text{CO}_2$  at the two stations. We used the cloud amount collected (in octas) at 0830 and 1730 hours. If the cloud amount is 0 at these two times, we considered the entire day as cloud free, and if it is more than or equal to seven, as a cloudy day. Fog is considered as a low cloud, which obscures the incoming radiation.



**Figure 7.** Times of inflexion of  $\text{CO}_2$  mixing ratios during different seasons in (a) Dehradun and (b) Gadanki. The values represent the increase in mixing ratios with respect to the mean.



**Figure 8.** Difference between CO<sub>2</sub> mixing ratio of NC and C days from 0830 to 1630 IST for (a) Dehradun and (b) Gadanki.

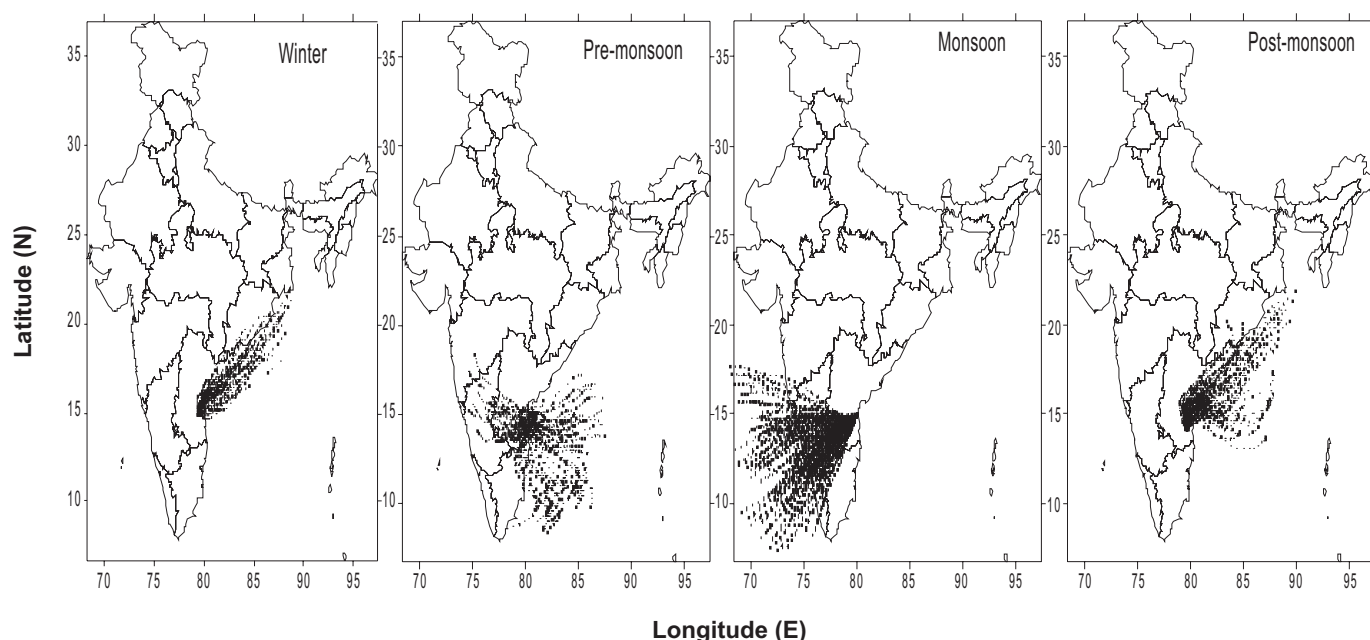
The cloudy (C) and non-cloudy (NC) days are considered in January (representing winter season) and April (representing summer season). In Dehradun, the C days are 3 January and 16 April, and NC days are 5 January and 21 April. Similarly, for Gadanki these days are 4 January and 23 April (C), and 12 January and 13 April (NC). The difference in CO<sub>2</sub> between the NC and C (NC minus C) days for Dehradun and Gadanki is shown in Figure 8(a) and 8(b), respectively.

In Dehradun, the mixing ratios are significantly higher during C days compared to those during NC days, in both the months. On the other hand, though the mixing ratios during the C days are more during January in Gadanki, they are significantly less in April. This could be because the grass canopy, which increases the CO<sub>2</sub> uptake during a NC day (Still et al),<sup>21</sup>

might have dried down in the summer and the uptake is primarily because of the trees present there.

Here, we considered the impact of cloudiness as an average value of the two observations at 0830 and 1730 hours. However, the impact of diurnal variation of cloudiness on CO<sub>2</sub> uptake and the impact of the nighttime cloudiness in increasing the temperature are not studied, which are beyond the scope of this paper.

(d) *Analysis of long-range flow of air mass.* Sharma et al<sup>14</sup> analyzed the influence of long-range transport of air mass from large point sources<sup>25</sup> associated with fossil fuel combustion and agriculture sector for Dehradun station using back trajectory method. It was found that Dehradun is not influenced by the air mass from any CO<sub>2</sub> source regions.



**Figure 9.** Seasonal changes in backward trajectories in Gadanki.



Similarly, the first 7 days back trajectory analysis of all the months is carried out for the Gadanki station. The seasonal flow of air mass toward Gadanki is shown in Figure 9. During winter and post-monsoon seasons, the direction of air mass is from northeast, whereas it is mainly from southeast during pre-monsoon season. However, as expected, trajectories are from the southwest during the monsoon. The analysis shows that Gadanki was also not influenced by the long-range transport of CO<sub>2</sub> from the source regions.

## Summary and Conclusions

Variations in atmospheric CO<sub>2</sub> mixing ratios in two climatically contrasting stations, Dehradun (subtropical humid) and Gadanki (tropical dry), are studied. As these mixing ratios are affected by the amount of water vapor present in the atmosphere, wet corrections are applied to the measured values by which the mixing ratios have increased by 2–14 ppm. Depending upon the moisture content in the atmosphere, the corrections of a humid station, Dehradun, are mainly controlled by the RH of the station. Monthly mixing ratios in a dry station are low compared to those in a wet station. Low mixing ratios are observed during the high vegetation activity showing the influence of photosynthetic activity on CO<sub>2</sub> mixing ratios. Maximum mixing ratios in Gadanki are observed one hour before those in Dehradun, which is because of the difference in sunrise time. After the sunrise, the mixing ratios decrease because of the increased photosynthesis and the atmospheric mixing processes. A distinct seasonality is observed in the variation of mixing ratios. Depending upon the type of canopy cover, the amount of cloudiness affected the mixing ratios.

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## Author Contributions

Conceived and designed the experiments: VKD. Analyzed the data: NS, YK, HG, AS, PM, KM. Wrote the first draft of the manuscript: NS, MMA, VKD. Contributed to the writing of the manuscript: NS, MMA, VKD. Agree with manuscript results and conclusions: NS, VKD, YK, PM, KM, HG, AS, MMA. Jointly developed the structure and arguments for the paper: VKD, NS, MMA. Made critical revisions and approved final version: VKD. All authors reviewed and approved of the final manuscript.

## DISCLOSURES AND ETHICS

As a requirement of publication the authors have provided signed confirmation of their compliance with ethical and legal obligations including but not limited to compliance with ICMJE authorship and competing interests guidelines, that the article is nei-

ther under consideration for publication nor published elsewhere, of their compliance with legal and ethical guidelines concerning human and animal research participants (if applicable), and that permission has been obtained for reproduction of any copyrighted material. This article was subject to blind, independent, expert peer review. The reviewers reported no competing interests.

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