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New radiolarian species from Late Jurassic chert-terrigenous deposits of the Taukha Terrane, Southern Sikhote-Alin

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Abstract. Late Jurassic and Early Cretaceous well preserved radiolarians are obtained from the Taukha Terrane in the southern Sikhote-Alin, Russia. The Taukha Terrane consists of three tectono-stratigraphic units, the Erdagou, Gorbousha and Skalistorechenka units. Three new species (Stichocapsa (?) pseudoconvexa, Sethocapsa taukhaensis and Cinguloturris primorika) are described from Late Jurassic chert-terrigenous sequences of the Gorbousha Unit cropping out in the Roudnaya River and Koreyskaya River areas, and their stratigraphic ranges are made precise.

Key words: Late Jurassic, new species, Radiolaria, Russia, Sikhote-Alin, Taukha Terrane

Introduction

The Taukha Terrane, located in the southeastern part of the Sikhote-Alin folded belt, represents a fragment of an Early Cretaceous accretionary prism that was accreted to the eastern margin of the Khanka-Bureya paleocontinent at the end of the Early Cretaceous (Barremian-Albian) (Figure 1A). The structure and genesis of the terrane was decoded by careful studies including lithological, structural and especially biostratigraphic research of individual sections of the terrane and their correlation with each other.

During this research, huge amounts of data about Jurassic to Early Cretaceous radiolarian fossils were accumulated (Kemkin *et al.*, 1992; Kemkin and Rudenko, 1993, 1994; Kemkin and Khanchuk, 1994; Kemkin, 1996a, 1996b; Kemkin *et al.*, 1997; Kemkin and Kemkina, 1998, 1999; Kemkin *et al.*, 1999 and others). Especially detailed studies of radiolarian assemblages were carried out in the Roudnaya River area of Dalnegorsk Town and the Koreyskaya River area (Kemkin and Kemkina, 1999; Kemkin *et al.*, 1999). The investigated assemblages of Late Jurassic and Early Cretaceous radiolarians are characterized by a large number of well preserved specimens and rich species diversity. Most of them are also well

known in other regions (Europe, America, Japan etc.) and allow us to determine the age of the deposits rather confidently. Nevertheless, some species have not been described yet. We have found some new species from the chert-terrigenous deposits of the Gorbousha Unit distributed in the Koreyskaya River and Roudnaya River areas. The descriptions and occurrences of these new radiolarian species are given.

Geological setting

The Taukha Terrane is separated from the neighbouring Samarka and Zhouravlevka terranes by faults (Figure 1A). According to new data (Kemkin and Kemkina, 2000), the terrane consists of three tectonostratigraphic units which overlie each other and are similar in lithology and structure, but differ in the age of the rocks (Figure 1B). Each unit is composed of paleo-oceanic deposits (mainly cherty and carbonate facies) in the lowest part. Toward their middle part, the units gradually change from these facies to terrigenous rocks derived from the continental margin. The terrigenous rocks are then replaced by olistostrome.

The lower unit (Erdagou Unit) is represented by Late Jurassic to Early Cretaceous (Berriasian) cherts and clayish cherts overlying Middle Jurassic (Callovian) basalts (Erdagou Suite) and followed by

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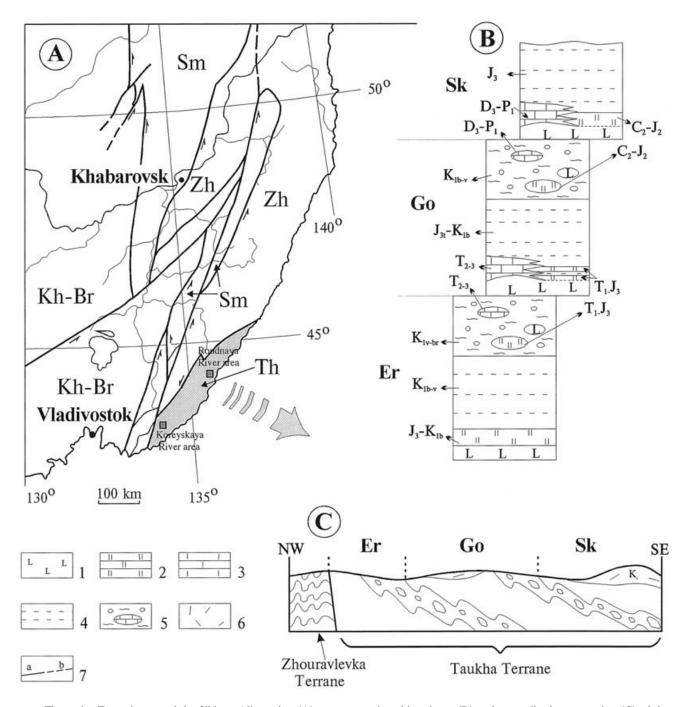


Figure 1. Tectonic map of the Sikhote-Alin region (A), tectonostratigraphic column (B) and generalized cross-section (C) of the Taukha Terrane (after Kemkin et al., 1999).

Kh-Br: Khanka-Bureya Superterrane, Sm: Samarka Terrane, Zh: Zhouravlevka Terrane, Th: Taukha Terrane, Er: Erdagou Unit, Go: Gorbousha Unit, Sk: Skalistorechenka Unit.

1: basalt, 2: chert, 3: limestone, 4: turbidite, 5: olistostrome, 6: Late Cretaceous volcanic rock, 7: character of contact between the lithological units; conformable (a) and estimated (b).

Berriasian-Valanginian turbidite deposits (Silinka Suite). The thickness of the cherts together with basalts is about 150 m, whereas the thickness of terrigenous rocks is estimated at 2500 m (Golozubov and Khanchuk, 1995). However, it is possible that the turbidite part is repeated tectonically several times. The turbidites conformably and gradually replace chert rocks by a set of siliceous mudstone and mudstone. Valanginian-Barremian olistostrome also conformably overlies the turbidites. The thickness of the olistostrome layers fluctuates from 100–200 to 400 m in different areas.

The middle unit (Gorbousha Unit) consists of Middle to Late Triassic limestone (Tetyukha Suite) with high-titanium alkaline basalts at the base (400–500 m thickness). Early Triassic to Late Jurassic cherts and clayish cherts (about 100 m) analogously change to late Tithonian to Berriasian turbidites (Gorbousha Suite), and then to Berriasian-Valanginian olistostrome. The thickness of the turbidites in different slides is 350–700 m (Volokhin *et al.*, 1990; Bragin, 1991 and others). The thickness of the olistostrome deposit is the same as that of the Erdagou Unit.

The upper unit (Skalistorechenka Unit) is composed of Late Devonian to Early Permian limestone (Skalistorechenka Suite) associated with high-titanium alkaline basalts (about 400 m) and Carboniferous to Middle Jurassic chert rocks overlapped by Late Jurassic turbidites (Pantovyi Creek Suite). The thickness of the chert and clastic deposits is not clear because they only have fragmentary outcrops.

The apparent structure of the Taukha Terrane shows inverted stratigraphy because the older marine formations and correspondingly overlapping terrigenous rocks (including olistostromes) comprise the upper structural levels of the terrane, and younger units comprise the base (Figure 1C). However, within each structural unit the stratigraphic succession is normal (from older to younger). Such structure of the terrane results from successive accretion of different-placed units from sea-floor spreading centre sites of a paleo-oceanic plate (Kemkin and Kemkina, 2000).

Radiolarian assemblages

In the Roudnaya River area (Dalnegorsk Town), twelve successive Jurassic to Early Cretaceous radiolarian assemblages were distinguished from the chert-terrigenous sequence of the Gorbousha Unit outcropping along the left bank of the Roudnaya River (Kemkin and Kemkina, 1998, 1999). They are the Parahsuum longiconicum, Hsuum altile, Parahsuum grande, Parahsuum hiconocosta, Hsuum mirabundum,

Hsuum matsuokai, Triactoma tithonianum, Archaeodictyomitra minoensis, Pseudodictyomitra okamurai, Xitus gifuensis, Pseudodictyomitra carpatica and Cecrops sp. assemblages. However, we think that the Pseudodictyomitra okamurai assemblage should be deleted from this list and included in the Archaeodictyomitra minoensis assemblage because the species composition of the two assemblages is rather the same. The age of each assemblage and the stratigraphical position of each sample of the section are shown in Figure 2. Among the eleven assemblages, the species composition and age assignment of Late Jurassic and Early Cretaceous assemblages are as follows:

The Archaeodictyomitra minoensis assemblage was extracted from the boudinaged fragments of cherts (sample P-6) in the contact zone of cherts and cherty mudstones and green cherty mudstones (samples P-5, 86-14). This radiolarian assemblage is characterized by rich specific diversity and contains mainly nassellarian genera. They are Archaeodictyomitra, Xitus, Mirifusus, Hsuum, Sethocapsa, Pseudodictyomitra, Stichocapsa and others. The age of the assemblage is middle Oxfordian to late Kimmeridgian. The assumed lower age boundary is based on the first appearance of the index species (Baumgartner et al., 1995). The upper age limit is restricted by the last appearance of Wrangellium okamurai (= Pseudodictyomitra okamurai Mizutani) (Baumgartner et al., 1995). The assemblage is comparable in age with that of the upper part of the Stylocapsa (?) spiralis zone and Hsuum maxwelli zone (Matsuoka, 1995) described in Japan. It should be noted that within the assemblage there are several morphological forms that are more common in the Early Cretaceous time. They are similar to Pseudodictyomitra nuda and Pseudodictyomitra leptoconica.

The Xitus gifuensis assemblage was also found in green cherty mudstones (samples P-4, P-3). Archaeodictyomitra and Xitus predominate among the numerous radiolarians of this assemblage. Some spumellarians such as Tritrabs, Emiluvia and Pantanellium are also found. The geological age of the assemblage is presumed to be late Kimmeridgian to middle Tithonian. The lower age boundary corresponds to the first appearance of Xitus gifuensis (Baumgartner et al., 1995). The upper boundary is restricted by the lower boundary of the upper assemblage. The assemblage is correlated with that of the Pseudodictyomitra primitiva zone (Matsuoka, 1995) and also contains some species which are similar to Pseudodictyomitra nuda and Pseudodictyomitra leptoconica.

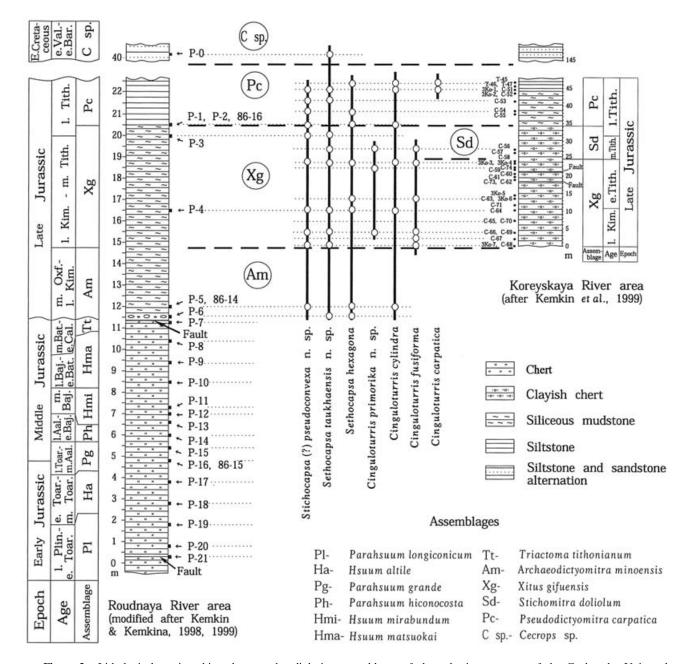


Figure 2. Lithological-stratigraphic column and radiolarian assemblages of chert-clastic sequences of the Gorbousha Unit and stratigraphic range of new radiolarian species.

The *Pseudodictyomitra carpatica* assemblage was found in black mudstones (samples P-1, P-2, 86-16) very close to the contact with green cherty mudstones. The assemblage is represented by numerous and diverse radiolarians correlated to the end of the Jurassic and the beginning of the Cretaceous time. Nassellarians such as *Pseudodictyomitra*, *Archaeodictyomitra*, *Xitus*, *Thanarla* and *Sethocapsa* are predominant. Spumellarians such as *Tritrabs*, *Emiluvia*,

Pantanellium and others are also contained therein. The age of the assemblage is late Tithonian. The lower age boundary corresponds to the first appearance of Pseudodictyomitra carpatica (Matsuoka, 1992, 1995). The upper boundary is determined by the last appearance of Archaeodictyomitra minoensis, Ristola altissima and Protunuma japonicus (Baumgartner et al., 1995). This assemblage also contains some species similar to Pseudodictyomitra nuda and Pseudodictyo-

mitra leptoconica. The assemblage is comparable in age with that of the lower part of the *Pseudodictyomitra* carpatica zone (Matsuoka, 1995) established in Japan.

The Cecrops sp. assemblage was distinguished in the black mudstones (sample P-0) that are 40 m above the horizon P-1 of the section. The assemblage is characterized by diverse and abundant radiolarian species. Most of them have a wide age interval. However, the presence of Cecrops sp. and Stichocapsa ex gr. altiforamina allows us to determine the age as early Valanginian to early Barremian (Baumgartner et al., 1995). This assemblage is comparable in age with that of the Cecrops septemporatus zone and the lower part of the Acanthocircus carinatus zone (Matsuoka, 1995).

In the Koreyskaya River area, three successive Late Jurassic radiolarian assemblages have been distinguished from the chert-clastic sequence of the Gorbousha Unit exposed along the bank of the Koreyskaya River (Kemkin et al., 1999). They are the Xitus gifuensis, Stichomitra doliolum and Pseudodictyomitra carpatica assemblages. The age of each assemblage and the stratigraphical position of each sample of the section are also shown in Figure 2. The species composition and age assignment of the assemblages are as follows:

The Xitus gifuensis assemblage was extracted from the greenish grey clayish cherts (samples 3Ko-3, 3Ko-4, 3Ko-5, 3Ko-6, 3Ko-7, C-58, C-61, C-63, C-64, C-65, C-66, C-67, C-68, C-69, C-70, C-71, C-74). This assemblage contains numerous well preserved radiolarian faunas which are characterized by rich species diversity. They are Archaeodictyomitra, Cinguloturris, Hsuum, Parvicingula, Pseudodictyomitra, Sethocapsa, Spongocapsula, Stichocapsa, Stichomitra, Thanarla, Xitus, Wrangellium and others. The geological age of the assemblage is presumed to be late Kimmeridgian to early Tithonian. The lower age boundary corresponds to the first appearance of Xitus gifuensis, Archaeodictyomitra excellens, Obesacapsula verbana and Sethocapsa horokanaiensis (Mizutani, 1981; Kawabata, 1988; Baumgartner et al., 1995). The upper boundary is restricted by the last appearance of such species as Eucyrtidiellum ptyctum, Parvicingula dhimenaensis, Sethocapsa funatoensis, Wrangellium okamurai, Triactoma blakei and Tritrabs exotica. This boundary corresponds to the early Tithonian (Pessagno, 1977; Mizutani, 1981; Aita, 1987; Baumgartner et al., 1995). The assemblage is correlated with the upper part of the *Hsuum maxwelli* zone and the lower part of the Pseudodictyomitra primitiva zone (Matsuoka, 1995).

The Stichomitra doliolum assemblage was found in the identical greenish-grey clayish chert (sample C-

56). The assemblage is represented by numerous and diverse well preserved radiolarians related to the end of Jurassic and the beginning of Cretaceous time. Nassellarians such as Pseudodictyomitra, Cinguloturris, Parvicingula, Archaeodictyomitra, Xitus, Thanarla, Stichocapsa, Stichomitra and Sethocapsa are predominant. The assemblage also contains spumellarians such as Tritrabs and Emiluvia. The range of the assemblage is middle Tithonian to early late Tithonian. The lower age boundary is based on the absence of characteristic species of the early Tithonian (see Xitus gifuensis assemblage), as well as the presence of Stichomitra doliolum, the first appearance of which is described from the middle Tithonian (Aita and Okada, 1986). The upper age boundary is controlled by the lower boundary of the next assemblage, as well as by the last appearance of Ristola altissima, which corresponds to early late Tithonian (Baumgartner et al., 1995). The assemblage is comparable in age with that of the Pseudodictyomitra primitiva zone (Matsuoka, 1995).

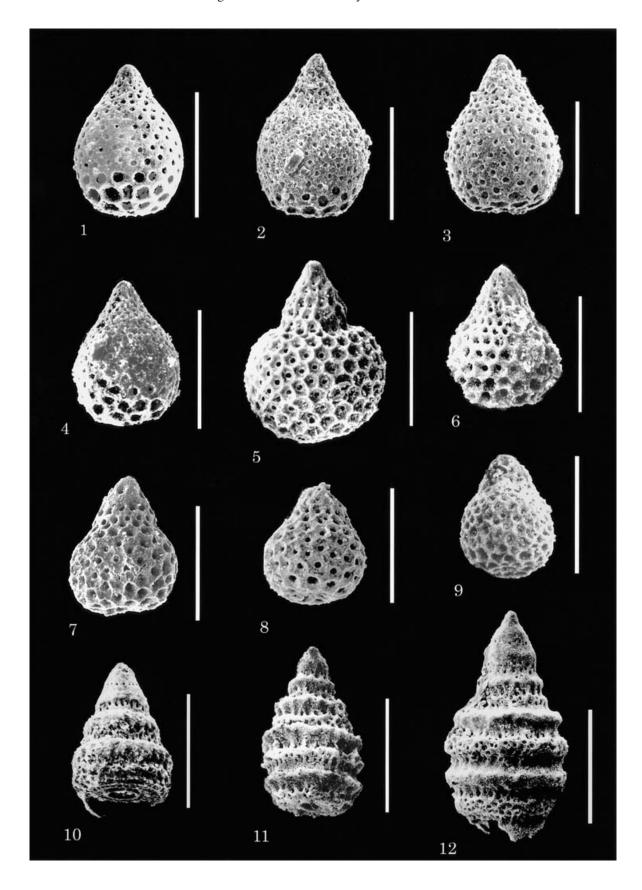
The Pseudodictyomitra carpatica assemblage was revealed in greenish-grey siliceous mudstone and black mudstone (samples T-45, T-46, 3Ko-1, 3Ko-2, C-51, C-53, C-54). Abundant well preserved radiolarians of the latest Jurassic to earliest Cretaceous were extracted from these samples. Representatives of nassellarians such as Archaeodictyomitra, Cinguloturris, Eucyrtidiellum, Hsuum, Parvicingula, Pseudodictyomitra, Sethocapsa, Stichocapsa, Stichomitra, Thanarla, Wrangellium, Xitus and others are predominant among them. The geological age of the assemblage is late late Tithonian. The lower age boundary is accepted according to the first appearance of Pseudodictyomitra carpatica and corresponds to late late Tithonian (Matsuoka, 1992). The upper boundary is determined by the last appearance of Archaeodictyomitra minoensis, Spongocapsula perampla, Protunuma japonicus, Eucyrtidiellum pyramis and Sethocapsa yahazuensis (Mizutani, 1981; Matsuoka and Yao, 1985; Aita, 1987; Baumgartner et al., 1995). The assemblage is correlated with the lower part of the Pseudodictyomitra carpatica zone (Matsuoka, 1995).

Systematic paleontology

Three new species are described and compared with co-occurring species of the same genera.

Sublass Radiolaria Order Nassellaria Family Theoperidae Genus *Stichocapsa* Haeckel, 1881

Stichocapsa Haeckel, 1881, p. 439



Type species.—Stichocapsa jaspidea Rüst, 1885

Stichocapsa (?) pseudoconvexa sp. nov.

Figure 3.1-3.4

Tricolocapsa sp. nov.; Kemkin et al., 1997, pl. 5, fig. 6. Sethocapsa sp. nov.; Kemkin et al., 1999, fig. 6.46.

Holotype.—0341-09 - C-51

Definition.—Shell is probably of four segments, lacking strictures. It is composed of two parts. Upper part of shell regularly and gradually increases in width, with segmental partitions only slightly visible outside. Cephalis spherical, small and poreless, without apical horn. Thorax subcylindrical with sparsely displaced small circular pores. Abdomen trapezoidal also with sparsely displaced small circular pores. Postabdominal segment globe-shaped, covered by sparse circular pores. The terminal portion of the test is closed, with 2–3 circular rows of large circular pores. The diameter of these pores is 3–5 times bigger than of the other pores.

Remarks.—Shell consists of upper conical and lower globose parts as well as whole genus, but it lacks apparent distal aperture. Therefore it is questionable that this form belongs to Stichocapsa. However, since its outline and surface are similar to Stichocapsa convexa Yao, this form is tentatively described as Stichocapsa. Stichocapsa (?) pseudoconvexa n. sp. is very similar to Stichocapsa convexa Yao by having almost the same shape and a smooth surface, but it differs from the latter by possessing a smaller cephalis and thorax chambers, and by having 2–3 circular rows of large circular pores in the terminal portion of the test.

Etymology.—Stichocapsa (?) pseudoconvexa means false (or mistaken) Stichocapsa convexa Yao.

Measurements (in µm).—(based on 5 specimens)

Height: 120–146 Max. width: 86–108

Type locality.—Right bank of Koreyskaya River, Gorbousha Unit of the Taukha Terrane, Sikhote-Alin, Russia.

Occurrence.—Archaeodictyomitra minoensis, Xitus gifuensis, Stichomitra doliolum and Pseudodictyomitra carpatica assemblages (middle Oxfordian to late

Tithonian).

Family Sethocapsidae Genus *Sethocapsa* Haeckel, 1881

Sethocapsa Haeckel, 1881, p. 433

Type species.—*Sethocapsa cometa* (Pantanelli); Rüst, 1885; Subsequent designation by Foreman, 1973.

Sethocapsa hexagona Hori, 1999

Figure 3.5

Sethocapsa sp.; Aita and Okada, 1986, pl. 5, fig. 13.Sethocapsa hexagona Hori, 1999, p. 74–75, Fig. 6.12–6.16, Fig. 11.1a, b.

Occurrence.—Archaeodictyomitra minoensis, Xitus gifuensis, Stichomitra doliolum and Pseudodictyomitra carpatica assemblages (middle Oxfordian to late Tithonian).

Sethocapsa taukhaensis sp. nov.

Figure 3.6-3.9

Sethocapsa sp.; Kemkin et al., 1997, pl. 6, fig. 6.

Holotype.—0377-06 – C-58

Definition.—Shell pear-shaped, probably of four segments as for all other species of the genus. Upper part of shell conical, regularly and gradually increases in width, with segmental partitions only slightly visible outside. Cephalis is subspherical, small, poreless, without apical horn. Thorax and abdomen are trapezoidal, with sparsely displaced small circular pores. Fourth chamber closed, globose and much larger than first three segments. Pores of the fourth segment circular and have hexagonal or pentagonal pore frames. Terminal portion of the test has 1–2 circular rows of large circular pores, diameters of which are 2–3 times larger than of the other pores.

Remarks.—Sethocapsa taukhaensis n. sp. is similar to Sethocapsa hexagona Hori, but differs from it by having a much smaller test, by lacking a stricture between the first three segments and the fourth segment, and also by having smaller-sized pores on the whole.

5. Sethocapsa hexagona Hori (C-54)

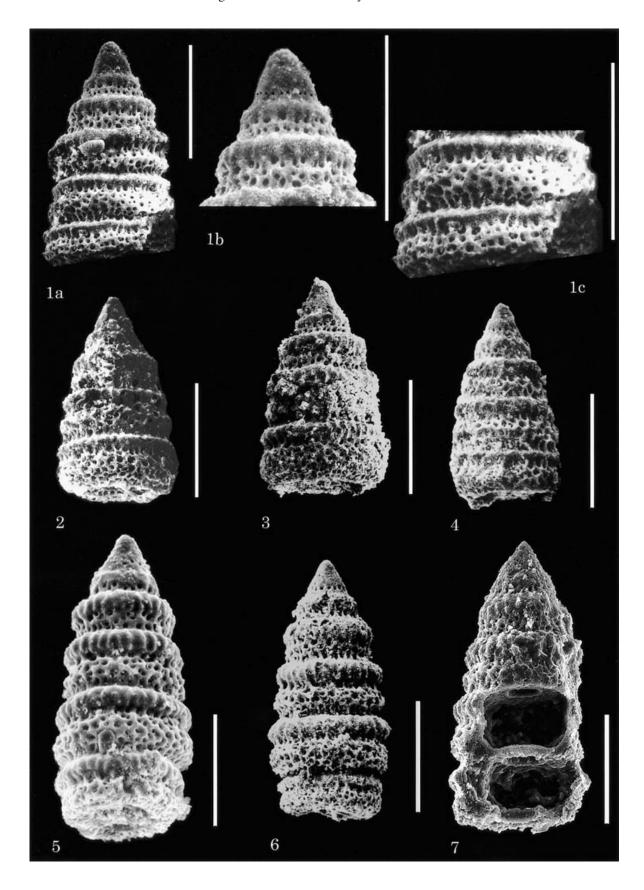
6–9. Sethocapsa taukhaensis sp. nov. (6 – C-58, holotype; 7 – P-3, 8 – P-6; 9 – C-56)

10. Cinguloturris carpatica Dumitrica (C-51)

11, 12. Cinguloturris fusiforma Hori (11 – C-69; 12 – C-56)

Figure 3. Late Jurassic radiolarians from chert-clastic sequence of the Gorbousha Unit. Scanning electron photomicrographs. All markers are 100 μm.

 $[\]textbf{1-4. } \textit{Stichocapsa} \ (?) \ \textit{pseudoconvexa} \ \text{sp. nov.} \ (1-\text{C-51, holotype}; \ 2-\text{P-3}; \ 3-\text{P-4}; \ 4-86\text{-}16)$



Etymology.—The species name is derived from the Taukha Terrane of the Sikhote-Alin region, Far East of Russia.

Measurements (in µm).—(based on 5 specimens)

Height: 103–127 Max. width: 83–103

Type locality.—Right bank of Koreyskaya River, Gorbousha Unit of the Taukha Terrane, Sikhote-Alin, Russia.

Occurrence.—Archaeodictyomitra minoensis, Xitus gifuensis, Stichomitra doliolum, Pseudodictyomitra carpatica and Cecrops sp. assemblages (middle Oxfordian to early Barremian).

Family Parvicingulidae Genus *Cinguloturris* Dumitrica in Dumitrica and Mello, 1982

Cinguloturris Dumitrica; Dumitrica and Mello, 1982, p. 22.

Type species.—Cinguloturris carpatica Dumitrica; Dumitrica and Mello, 1982.

Cinguloturris carpatica Dumitrica in Dumitrica and Mello, 1982

Figure 3.10

Theoperidae gen. et sp. indet. I; Aita, 1982, pl. 2, fig. 18. Theoperid gen. et sp. indet.; Aoki and Tashiro, 1982, pl. 2, fig. 9. Cinguloturris carpatica Dumitrica; Dumitrica and Mello, 1982, p. 23, pl. 4, figs. 7-11; Yao, 1984, pl. 2, fig. 28; Aita, 1985, fig. 7.12; Ishida, 1985, pl. 4, figs. 13, 14; Matsuoka and Yao, 1985, pl. 2, fig. 13; Tanaka et al., 1985, pl. 1, fig. 12; Matsuoka, 1986, pl. 2, fig. 16, pl. 3, fig. 11; Matsuoka and Yao, 1986, pl. 2, fig. 14; Aita, 1987, p. 64, pl. 10, fig. 12; Wakita, 1988, pl. 5, fig. 8; Kato and Iwata, 1989, pl. 5, fig. 5, pl. 6, fig. 10; Matsuoka and Oji, 1990, pl. 1, fig. 7; Yao, 1990, pl. 4, fig. 11; Ishida and Hashimoto, 1991, p. 46, pl. 4, fig. 16; Matsuoka, 1992, pl. 4, fig. 1; Kashiwagi and Yao, 1993, pl. 2, fig. 4; Sashida et al., 1993, fig. 6.7; Gorican, 1994, p. 64, pl. 23, figs. 1, 6-11; Kemkin and Rudenko, 1994, p. 1118, pl. 1, figs. 2, 5-8; Baumgartner et al., 1995, p. 142, pl. 3193, figs. 1, 2, 4-6 (not 3); Ohba and Adachi, 1995, pl. 3, fig. 10; Kemkin, 1996a, pl. 1, figs. 1, 2, 5; Sashida and Uematsu, 1996, p. 65, fig. 4.11; Shen et al., 1996, pl. 2, fig. 11; Nishizono et al., 1997, pl. 2, fig. 13; Ishida and Hashimoto, 1997, pl. 2, fig. 5; Takashima et al., 1997, fig. 5.10; Yang and Matsuoka, 1997, pl. 3, fig. 9; Zyabrev and Matsuoka, 1999, pl. 1, fig. 11; Hori, 1999, p. 91, fig.

Dictyomitra sp. B; Ishida, 1983, pl. 5, figs. 3, 4. Unnamed multicyrtoid nassellarian; Yamamoto, 1983, pl. 1, fig. 10. Cinguloturris sp. aff. C. carpatica Dumitrica; Yao, 1984, pl. 3, fig. 19. Cinguloturris sp. cf. C. carpatica Dumitrica; Tanaka et al., 1985, pl. 1, fig. 7; Kurimoto, 1989, pl. 1, fig. 17.

Occurrence.—Pseudodictyomitra carpatica assemblage (late Tithonian).

Cinguloturris cylindra Kemkin and Rudenko, 1993

Figure 4.5-4.7

Unnamed multicyrtoid nassellarian; Adachi, 1982, pl. 2, fig. 9. *Cinguloturris* sp.; Wakita, 1987, pl. 1, fig. 8; Kato and Iwata, 1989, pl. 2, fig. 7.

Cinguloturris sp. aff. C. carpatica Dumitrica; Wakita, 1988, pl. 6, fig. 14; Sashida et al., 1993, fig. 6.8, 6.9.

Cinguloturris carpatica Dumitrica; ?Yasuda, 1989, pl. 1, fig. 14.
Cinguloturris sp. cf. C. carpatica Dumitrica; Kemkin et al., 1992, p. 33, pl. 2, fig. 3.

Cinguloturris cylindra Kemkin and Rudenko, 1993, p. 116, pl. 2, figs. 3, 9–15; Kemkin and Rudenko, 1994, p. 1116, pl. 1, figs. 3, 9–15; Baumgartner et al., 1995, p. 144, pl. 6101, figs. 1–4; Kemkin, 1996a, pl. 3, fig. 11; Dumitrica et al., 1997, p. 69, pl. 14, fig. 16; Kemkin et al., 1997, pl. 3, fig. 9, pl. 4, fig. 11, pl. 6, fig. 3; Hori, 1998, pl. 2, fig. 12; Hori, 1999, p. 91–93, fig. 9.2; Kemkin et al., 1999, fig. 4.8, 4.9, fig. 5.9; Matsuoka et al., 2002, fig. 6.7.

Cinguloturris sp. A; Gorican, 1994, p. 64, pl. 23, figs. 3–5. *Cinguloturris arabica* Jud, 1994, p. 69, pl. 5, figs. 15, 16.

Occurrence.—Archaeodictyomitra minoensis, Xitus gifuensis, Stichomitra doliolum and Pseudodictyomitra carpatica assemblages (middle Oxfordian to late Tithonian).

Cinguloturris fusiforma Hori, 1999

Figure 3.11, 3.12

Unnamed multicyrtoid nassellarian; Adachi, 1982, pl. 2, figs. 8, 10. *Cinguloturris carpatica* Dumitrica; Kishida and Hisada, 1986, fig. 2.12; Kawabata, 1988, pl. 2, fig. 10; Tamba Belt Research Group, 1990, pl. 2, fig. 16; Sashida *et al.*, 1993, fig. 6.6.

Cinguloturris sp. A; Kemkin and Rudenko, 1994, p. 1118, pl. 1, fig. 6. Cinguloturris sp.; Hori, 1998, pl. 2, fig. 13.

Cinguloturris fusiforma Hori, 1999, p. 93, fig. 9.3–9.6, fig. 11.6. Cinguloturris sp. nov.; Kemkin et al., 1999, fig. 5.10.

Occurrence.—Xitus gifuensis and Stichomitra doliolum assemblages (late Kimmeridgian to middle Tithonian).

Cinguloturris primorika sp. nov.

Figure 4.1-4.3

Cinguloturris carpatica Dumitrica; Baumgartner et al., 1995, p. 142-

- Figure 4. Late Jurassic radiolarians from chert-clastic sequence of the Gorbousha Unit.
 Scanning electron photomicrographs. All markers are 100 μm.
 - **1–3.** *Cinguloturris primorika* sp. nov. (1 C-69, holotype; 2 C-56; 3 C-74)
 - 4. Cinguloturris sp. aff. C. primorika sp. nov. (C-56)
 - 5-7. Cinguloturris cylindra Kemkin and Rudenko (5 T-47; 6 C-58; 7 Be-9 from the fourth right tributary of Benevka River)

143, pl. 3193, fig. 3; ?Matsuoka, 1992, pl. 3, fig. 2; Hori, 1998, pl. 2, fig. 11.

Holotype.—0409-02 – C-69

Definition.—Multisegmented conical test as with all other species of the genus. The number of segments is seven or more. The diameter of the segments increases rapidly from the proximal to the distal portion of the test. Cephalis subspherical, smooth, poreless and without apical horn. Thorax also smooth, poreless and trapezoidal. Thorax and abdomen are separated from one another by a single row of small circular pores. Each post-thoracic segment begins with a narrow, inflated, poreless and slightly costate circumferential ridge. The remaining segment wall is covered with a thick spongy meshwork with irregularly arranged pores. The ratio of height between ridge and spongy meshwork is 1:3-1:5. Meshed portion of segment wall and ridge are separated from one another by a single row of larger pores.

Remarks.—Cinguloturris primorika n. sp. differs from Cinguloturris cylindra Kemkin and Rudenko and Cinguloturris fusiforma Hori by possessing a distinct conical shell shape. From Cinguloturris carpatica Dumitrica it differs by having very narrow, slightly costate circumferential ridges.

Etymology.—The species name is derived from Primorye area, southern part of the Far East of Russia.

Measurements (in µm).—(based on 6 specimens)

Height: 166–186 Max. width: 103–120

Type locality.—Right bank of Koreyskaya River, Gorbousha Unit of the Taukha Terrane, Sikhote-Alin, Russia

Occurrence.—Xitus gifuensis and Stichomitra doliolum assemblages (late Kimmeridgian to middle Tithonian).

Cinguloturris sp. aff. C. primorika sp. nov.

Figure 4.4

Remarks.—*Cinguloturris* sp. aff. *C. primorika* n. sp. differs from *Cinguloturris primorika* n. sp. by having less prominent circumferential ridges.

Occurrence.—Stichomitra doliolum assemblage (middle Tithonian).

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