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# The Barremian heteromorph ammonite *Dissimilites* from northern Italy: Taxonomy and evolutionary implications

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A new acrioceratid ammonite, *Dissimilites intermedius* sp. nov., from the Barremian (Lower Cretaceous) of the Puez area (Dolomites, northern Italy) is described. *Dissimilites intermedius* sp. nov. is an intermediate form between *D. dissimilis* and *D. trinodosum*. The new species combines the ribbing style of *D. dissimilis* (bifurcating with intercalating single ribs) with the tuberculation style of *D. trinodosum* (trituberculation on entire shell). The shallow-helical spire, entirely comprising single ribs intercalated by trituberculated main ribs, is similar to the one of the assumed ancestor *Acrioceras*, whereas the increasing curvation of the younger forms resembles similar patterns observed in the descendant *Toxoceratoides*. These characters support the hypothesis of a direct evolutionary lineage from *Acrioceras* via *Dissimilites* to *Toxoceratoides*. *D. intermedius* sp. nov. ranges from the upper Lower Barremian (*Moutoniceras moutonianum* Zone) to the lower Upper Barremian (*Toxancyloceras vandenheckii* Zone). The new species allows to better understand the evolution of the genus *Dissimilites*. In the *Kotetishvilia compressissima* Zone, two morphological forms developed: smaller forms very similar to *Acrioceras* and forms with very long shaft and juvenile spire like in *D. intermedius* sp. nov. The latter most likely gave rise to *D. subalternatus* and *D. trinodosum* in the *M. moutonianum* Zone, forms which were probably ancestral to the genus *Toxoceratoides*.

Key words: Ammonoidea, Acrioceratidae, Dissimilites, Barremian, Cretaceous, Alps, Italy.

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## Introduction

The ammonoids from the Lower Cretaceous sediments in the Puez area (Dolomites, South Tyrol, North Italy) have been mostly utilized so far for the stratigraphical purposes (Lukeneder and Aspmair 2006; Lukeneder 2010, 2011, 2012b), a task that has been recently summarized by Lukeneder (2012b). The present paper provides a systematic revision of one of the stratigraphically important forms, the heteromorph acrioceratid ammonite *Dissimilites*. All known occurrences of *Dissimilites* are listed and dicussed and a reconstruction of its evolution is proposed.

*Institutional abbreviations.*—NHMW, Natural History Museum Vienna, Austria; STMN, South Tyrol Museum of Nature Sciences (PZO and CP inventory numbers), Bozen, Italy.

Other abbreviations.—CP, Col Puez; Dsp, spire diameter; Hf, whorl height at flexus; Hp1, whorl height at start of

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proversum; Hp2, whorl height at end of proversum; Hr, max. whorl height of retroversum; ir, intercalatory ribs; L, max. length; Lp, proversum length; Lr, retroversum length; min Hsp, minimal whorl height of spire; mr, main ribs; P, Puez; Pp1, proversum part 1; Pp2, proversum part 2; Pp1b, proversum part 1 breadth; Pp2b, proversum part 2 breadth; sp, spines; Spb, spire breadth;  $\alpha$ , angle of retroversum;  $\beta$ , angle formed by pro- and retroversum;  $\gamma$ , open angle on inner side of proversum;  $\delta$ , angle at end of proversum; -x- marks position of main tuberculated ribs on spire and proversum and bifurcating ribs on retroversum in given rib-formulae; 1 or 2 shows value of intercalatories.

## Material

The ammonite specimens discussed herein originate from the Puez locality (Dolomites; Fig. 1). The material collected

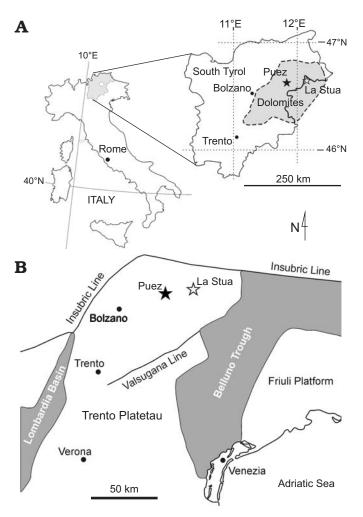


Fig. 1. Locality maps of the study area. **A**. Location of South Tyrol in northern Italy and Puez area (black star) and neighbouring La Stua section (white star) within the Dolomites. **B**. Position of the Puez and La Stua localities on the Trento Plateau.

by Christian Aspmair (Bozen, Italy) is stored in the STMN. The bed-by-bed collected material assembled by the authors is stored at the NHMW. Ammonites are moderately well preserved, mostly as compressed inner molds without shell. In total, fourteen specimens of *Dissimilites* have been examined (Figs. 2–4).

The entire Upper Hauterivian to Upper Barremian ammonite fauna from Puez consists of 44 genera in 17 families (Lukeneder 2012b). Acrioceratidae are represented by *Acrioceras* Hyatt, 1900 and *Dissimilites* Sarkar, 1954, Ancyloceratidae by *Toxancyloceras* Delanoy, 2003 and *Audouliceras* Thomel, 1965, and Heteroceratidae by *Moutoniceras* Sarkar, 1954.

## Geological setting and biostratigraphy

The ammonite-bearing outcrop at Puez is situated on the southern side of the Piz de Puez, Puez-Odle-Gardenaccia

Plateau in the Dolomites, northern Italy (Lukeneder 2010). During Early Cretaceous times the locality was situated in the northern part of the Trento Plateau. La Stua locality is located 20 km to the east of the Puez area (Stöhr 1993, 1994), in the eastern part of the Trento Plateau, geographically near the northern border of the Veneto district. During the Barremian, the Trento Plateau constituted a part of the so-called Adriatic Plate. The Adriatic Plate (also called Apulian Plate) of the South Alpine-Apennine Block was bordered by the Penninic Ocean (= Alpine Tethys) to the north and the Vardar Ocean to the south-east (Cecca 1998; Dercourt et al. 1993; Stampfli and Mosar 1999; Scotese 2001; Stampfli et al. 2002).

Lower Cretaceous sediments on the Trento Plateau are represented mostly by limestone, marly limestone, and marl. The Puez Formation in the Puez area comprises three members, i.e., from bottom to top, the Puez-Limestone, Puez-Redbed, and Puez-Marl Member (Lukeneder 2010). A detailed description of the geology and lithostratigraphy is given by Lukeneder (2010). Ammonites, including *Dissimilites*, derive from the lowermost part of the formation, the Puez-Limestone Member. An ammonite-based stratigraphical framework has recently been refined by Lukeneder (2012b).

During the late 19th and early 20th century, a rich cephalopod fauna was collected from Lower Cretaceous sediments from this area by Haug (1887, 1889), Hoernes (1876), Uhlig (1887), Rodighiero (1919), Reithofer (1928), and Pozzi (1993). Additionally, microfacies and ammonites were investigated from the "Alpe Puez" by Cita and Pasquaré (1959) and Cita (1965), who assumed a Hauterivian to Barremian age for the sediments outcropping in the Puez area. After these publications no further investigations were undertaken at the main locality of Puez, except for a few descriptive papers dealing with some small ammonoid faunas from different localities near Puez, most of all from La Stua (Baccelle and Lucchi-Garavello 1967; Stöhr 1993, 1994). The latter contributions compared the faunas from La Stua with the Puez ammonite faunas reported by Haug (1887, 1889) and Uhlig (1887). The most recent works on the Lower Cretaceous from this area were focused on stratigraphy (Lukeneder and Aspmair 2006), palaeoecology (Lukeneder 2008), lithostratigraphy (Lukeneder 2010, 2011), and ammonite biostratigraphy (Lukeneder 2012b).

In this paper we follow the ammonite zonation established by the Cretaceous Ammonite Working Group (Reboulet et al. 2009) which has been refined for the Puez area by Lukeneder (2012b). Assessing the evolutionary patterns of *Dissimilites* requires careful stratigraphic control. Therefore we found it necessary to add up-to-date definitions of the ammonite zones in order to indicate exact stratigraphic positions of occurence and avoid confusion with former interpretations in earlier papers.

Nicklesia pulchella and Kotetishvilia compressissima zones.

—The *N. pulchella* Zone was defined by Lukeneder (2012a), ranging from the bottom of bed P1/105 up to the top of bed P1/111 (Fig. 5). New findings show that *M. moutonianum* goes down to bed P1/103, which means that the *N. pulchella* 

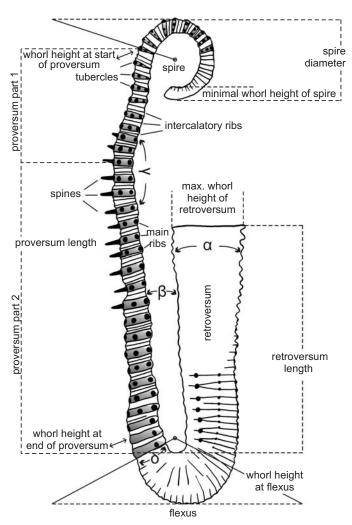


Fig. 2. Dimensions and terms used in *Dissimilites intermedius* sp. nov. from Puez, based on the holotype, the most entire specimen (NHMW 2012/0002/0001).  $\alpha$ , angle of retroversum;  $\beta$ , angle formed by pro- and retroversum;  $\gamma$ , open angle on inner side of proversum;  $\delta$ , angle at end of proversum.

and *K. compressissima* zones are even more condensed (i.e., 3.5 m) than it was interpreted so far. No members of Acrioceratidae (*Acrioceras*, *Dissimilites*) occur here. The *K. compressissima* Zone extends up to the top of bed P1/102.

**Moutoniceras moutonianum Zone**.—The first occurrence of the index ammonite *M. moutonianum* has previously been thought to occur in bed P1/119 (Lukeneder 2012b). More recent findings show that *M. moutonianum* appears already in bed P1/103 associated with *Macroscaphites tirolensis*, extending therefore the *M. moutonianum* Zone into an interval from P1/103 up to P1/137. The Acrioceratidae occur with *Dissimilites intermedius* sp. nov. (bed P7/-3.4 m = bed P1/103, P7/ -3.2 m = P1/105, P1 bed 110) and *Dissimilites trinodosum* (beds P1/130). Heteroceratidae are present at the base of the zone for the first time with the zonal index ammonite *M. moutonianum*. Macroscaphitidae occur with singletons of two species, i.e., *Costidiscus nodosostriatum* and *M. tirolensis*. Toxancyloceras vandenheckii Zone.—The T. vandenheckii Zone reaches from the bottom of bed P1/138 up to the top of bed P1/193 (Fig. 5). The appearance of the index ammonite T. vandenheckii in bed P1/138 marks the lower boundary of the T. vandenheckii Zone. The T. vandenheckii Zone is subdivided into the T. vandenheckii Subzone and the younger B. barremense Subzone. Acrioceratidae occur with D. trinodosum (beds P1/146, 169, 172). Dissimilites is accompanied by the family Ancyloceratidae with the zonal index species T. vandenheckii and Toxancyloceras sp. Within the family Leptoceratoididae, Karsteniceras sp. occurs in the lower part of the zone. Hamulinidae occur only in the lower part with Hamulina sp., Anahamulina cf. uhligi and Anahamulina sp., and with Pychohamulina cf. acuaria in both parts. The lower part yields members of Duyeina cf. kleini and Duyeina sp., whereas the upper part is characterized by the occurrence of the hamulinid member Vasicekina cf. pernai. Macroscaphitidae occur with a single specimen at the base, along with Costidiscus sp.

## Systematic palaeontology

The authors follow the basic classification of Cretaceous Ammonoidea by Wright et al. (1996), Vermeulen (2004, 2006), and Klein et al. (2007).

Class Cephalopoda Cuvier, 1798 Order Ammonoidea Zittel, 1884 Suborder Ancyloceratina Wiedmann, 1966 Superfamily Ancyloceratoidea Gill, 1871 Family Acrioceratidae Vermeulen, 2004 Subfamily Helicancyclinae Hyatt, 1894 Genus *Dissimilites* Sarkar, 1954

*Type species: Hamites dissimilis* d'Orbigny, 1842 from Escragnolles, Lower Barremian, SE France; subsequent designation by Sarkar 1954: 619; lectotype designated by Busnardo in Gauthier 2006: pl. 20: 8.

Dissimilites intermedius sp. nov.

#### Figs. 3A-C, F, G, I, 4A-C, F, G, I.

1955 *Dissimilites* cf. *dissimilis* d'Orb. sp. (forme n° 1); Sarkar 1955: 138, pl. 4: 17.

1994 Ancyloceras; Costamoling and Costamoling 1994: text-fig. 93.

- 2006 Acrioceras (Dissimilites) dissimilis (d'Orbigny); Lukeneder and Aspmair 2006: pl. 6: 7.
- 2006 Dissimilites (Dissimilites) trinodosum (d'Orbigny); Lukeneder and Aspmair 2006: pl. 7: 1.

2012 Dissimilites trinodosum; Lukeneder 2012b: text-fig. 6B.

*Etymology*: After the intermediate morphology between *Dissimilites dissimilis* and *Dissimilites trinodosum*.

*Type material:* Holotype, the specimen NHMW 2012/0002/0001 (Figs. 3A, 4A), mainly flattened specimen, in partly limonitic sculpture mould, with perfectly preserved sculpture. Only one side of the holotype is available for observation and retains perfect conch outlines without any trace of fragmentation. The embryonal part of the shell and suture lines are not visible. The holotype is relatively large, which indicates the presence of the body chamber and thus an adult stage of

growth. Paratypes, specimens NHMW 2012/0002/0002 (Figs. 3B, 4B), NHMW 2012/0002/0003 (Figs. 3C, 4C) and NHMW 2011/0159/0027 (Figs. 3G, 4G).

*Type locality*: Southern side of the Piz de Puez (Fig. 1A; E 011°49'15", N 46°35'30") on the Puez-Odle-Gardenaccia Plateau in the Dolomites, Trentino-Alto Adige, South Tyrol (Lukeneder 2010, 2011, 2012b).

*Type horizon:* Puez log P7, interval P1/-3.0 to P1/-3.5, bed P7/-3.4 m (= bed P1/103). This interval in P7 is equivalent to beds P1/102–105. *Moutoniceras moutonianum* Zone, Barremian, Lower Cretaceous.

*Material.*—Type material and two additional specimens from the type locality (PZO-CP56 and NHMW 1886/0008/0017).

*Diagnosis.*—Small acrioceratid coiled shells with a weak torsion in spire. Whorl section elliptical to subrounded. Shell tripartite with juvenile spire, proversum and adult final retroversum connected by a flexus (i.e., hooked). The sculpture comprises relatively dense, sharp, simple and bifurcating ribs, each bearing three tubercles throughout the shell.

#### Measurements.—See Table 1.

Rib formulae.-Holotype, NHMW 2012/0002/0001, spire x-x-x-1-x-1-x-2-x-1-x-1-x-1-x-, proversum Pp1 is -2-x-1-x-1-x-1-x-1-x-1-x-1-x-. The second part of the proversum (Pp2) shows 19 tuberculated ribs. Rib formula of Pp2 is -2-x-1-x-2-x-1-x-1-x-2-x-2-x-1-x-1-x-2-x-1-x-1-x-1-x-1-x-0-x-0-x- and retroversum -x-x-x-x-1-x-x-2-x-1-x-2-x-4-. NHMW 2012/0002/0002, proversum NHMW 2012/0002/0003, proversum -x-1-x-1-x-1-x-1-x-1-x-1-x-1-x-1-x-1-x-1-x-1-x-x-x-x-x-and retroversum -x-x-3-x-1-x-9-. NHMW 2011/0159/0027, 1-x-1-x-x-x-x-x-. NHMW 1886/0008/0017, proversum and retroversum -x-x-x-, specimen possesses a rib anomaly with rib-displacement on one side (e.g., forma verticata after Hölder 1956), at the start of flexus. This anomaly causes trifurcate ribbing on flexus. PZO-CP56, proversum -x-1-x-1-x-1-x-1-x-1-x-1-x-1- and retroversum -x-x-x-1x-1-1-1-x-1-1-1 (specimen shows a suture line with trifid lateral lobus).

*Description.*—Small, acrioceratid coiled shell with weak torsion in spire. Whorl section subrounded. Shell tripartite with juvenile spire, proversum and adult final retroversum

connected by a flexus. The sculpture comprises relatively dense, sharp, simple and bifurcating ribs, partly fibulous on proversum. Suture line not preserved. The holotype (NHMW 2012/0002/0001) is 131 mm long. Ribs cross the venter without interruption except at the flexus, where a ventral depression or a smooth, shallow furrow occurs. All ribs are trituberculated. Lateral and ventral tubercles decrease in size from proversum over flexus to retroversum.

*Spire.*—The spire starts at a whorl height of 3 mm. On its first third, single fine ribs are visible (approx. 14). Ribbing is slightly retroverse. At a whorl height of 4 mm the first trituberculated ribs occur. Tubercles are located at a dorso-lateral (umbo-lateral), lateral and ventro-lateral position. Ten of these trituberculated main ribs occur on the spire up to a diameter of 6 mm. Tuberculated ribs are intercalated by 1 or 2 thinner, single ribs. Rib formula starts from the first main rib (tuberculated) after 14 single ribs (see rib formula). All ribs cross the venter without interruption. The spire seems to have weak torsion. The spire ends at a maximum diameter of 21 mm (Fig. 2).

Proversum.--The proversum starts at a whorl height of 6 mm. It is subdivided in a first part (Pp1) and a second part (Pp2). The change from Pp1 to Pp2 is marked by a distinctly open angle at the first third of the proversum (approx. 160–168°). The maxium length of the entire proversum is 110 mm. The whorl height opens with a value of approx.  $\delta$ 3°. The proversum shows thick (up to 2 mm), single, trituberculate main ribbing intercalated by single, sharp, thinner ribs (one third of main ribs) without tubercles. Ribbing is radial on Pp1 and is followed on Pp2 by slightly proverse ribbing. Main ribs start from the dorso-lateral tubercles. Ribs cross the venter without interruption. On the dorsal part of the proversum, fine ribs occur at the height of the dorso-lateral tubercles (twice as many ribs as on the flanks) and cross the dorsal area. In the mid part of the proversum, 10 spines (up to 4 mm in length) are visible starting from the ventro-lateral tubercles. Note that each tubercle gives rise to a spine. Spines are tiny, thin and are normally missing due to sampling or preparation. This was also proven by computer tomographical analysis of the holotype (Lukeneder 2012a). Pp1 shows 9 tuberculated ribs. Pp2 bears 19 tuberculated ribs (see rib formulae). The proversum ends at the last thick, tuberculated main rib at a whorl height of 11 mm. The last 3–6 main ribs show a fibulation (Figs. 2, 3), notable on every specimen

Table 1. Measurements (in mm). Abbreviations: Dsp, spire diameter; Hf, whorl height at flexus; Hp1, whorl height at start of proversum; Hp2, whorl height at end of proversum; Hr, max. whorl height of retroversum; L, max. length; Lp, proversum length; Lr, retroversum length; min Hsp, minimal whorl height of spire; Pp1b, proversum part 1 breadth; Pp2b, proversum part 2 breadth; Spb, spire breadth;  $\alpha$ , angle of retroversum;  $\beta$ , angle formed by pro- and retroversum;  $\gamma$ , open angle on inner side of proversum;  $\delta$ , angle at end of proversum.

Inventory number	L	Lp	Lr	Dsp	minHsp	Hp1	Hp2	Hf	Hr	Spb	Pp1b	Pb2b	α	β	γ	δ
NHMW 2012/0002/0001, ht	131	110	65	21	3	6	11	13	19	1-2	3	5	14–15	4	160–162	3
NHMW2012/0002/0002, pt	82	69	26	-	-	3	9	10	12	-	-	4	3	5	160-165	6
NHMW 2012/0002/0003, pt	83	71	51	-	-	4	9	8	15	-	2	3	7	-4	160-163	4
NHMW 2011/0159/0027, pt	65	55	-	-	-	-	8	9	-	-	_	2	-	-	160-162	5
NHMW 1886/0008/0017	53	42	26	-	-	-	10	11	13	-	_	9	-	_	-	3
PZO-CP56	49	_	38	_	_	_	9	10	12	_	_	4	3	-7	_	2

starting from the lateral to the ventro-lateral tubercles, instead of intercalatory ribs. The last main rib marks approximately the starting position of the body chamber.

Paratype NHMW 2011/0159/0027 (bed P1/110; Figs. 3G, 4G) shows the fibulation on the last 10–12 trituberculated main ribs on the proversum more precisely. Fibulation occurs between the lateral and ventro-lateral tubercles, instead of intercalatory ribbing. The same pattern is visible in paratypes NHMW 2012/0002/0002 (bed P7/ -3.2 m = P1/105; Figs. 3C, 4C) and NHMW 2012/0002/0003 (bed P7/ -3.2 m = P1/105; Figs. 3B, 4B).

The three paratypes are almost equal in size (65 mm, 82 mm, 83 mm) and adult (as indicated by the almost entire retroversum), but somewhat smaller than the holotype (131 mm); this size difference may reflect sexual dimorphism (e.g., micro- vs. macroconchs) or a decrease of elongation of the proversum in evolution, resulting in a shorter proversum shaft, since the holotype (bed P7/-3.4 m) was found approximately 20 cm below the paratypes (bed P7/-3.2 m). Paratypes show the typical curvation of the first part of the proversum and its shallow-helical whorl morphology, ornamented with trituberculated main ribs. All specimens exhibit this characteristic feature of the species: trituberculation from spire over flexus to final retroversum combined with irregular alternation of single and bifurcating ribs on retroversum, flat ventral side on flexus (Figs. 3I, 4I,) and fibulation of main ribs on Pp2 of proversum. The paratypes show a comparable change in angle from the Pp1 to the Pp2, which starts from the flexus (most external point of flexus) in paratypes NHMW 2011/0159/0027 at 54 mm, in NHMW 2012/0002/0002 at 60 mm and in NHMW 2012/0002/0003 at 69 mm.

*Flexus.*—No thick main ribs occur. Ribbing is sharp and starts from 5–6 less pronounced, elongated tubercles to clavae from the umbilical region. Ribs bifurcate from the umbilical tubercles or at the mid flank from small, tiny tubercles. The ribs are still trituberculated at this point. Umbilical and ventro-lateral tubercles are bigger than lateral tubercles. Lateral tubercles occur on mid flank on every branch of the bifurcating ribs. A shallow, almost smooth area occurs on the venter throughout the flexus (Figs.  $3I_3$ ,  $4I_3$ ). The flexus shows a maximum whorl height of 13 mm. Fine ribs cross the dorsal area as described for the proversum. As noted above, all tubercles give rise to tiny spines, which has been proven by computer tomography (Lukeneder 2012a).

*Retroversum.*—The retroversum opens from the proversum at an angle  $\beta$  of approx. 4° and shows a maximum length of 78 mm (Fig. 2; see Table 1). The distance between pro- and retroversum is the same as the whorl height of the proversum or half the retroversum whorl-height. Shell height increases by  $\alpha$  14–15° up to the aperture, ending with a peristome. Ribbing is characteristic for the new species. Ribs are sharp and radial. Ribs bifurcate from umbilical tubercles and cross the venter without interruption. Bifurcating ribs are intercalated irregularly by straight single ribs also crossing the venter and showing the same thickness. All ribs on the retroversum (body chamber) show trituberculation. The umbo-lateral tubercle is the most prominent one, followed by a smaller ventro-lateral tubercle and by the tiny lateral tuberculation. The lateral tuberculation is situated at approximately twothirds of the flanks height. The position is the same as on the anterior flexus. The aperture (peristome) is marked by a collared last rib. The dorsal area shows the same ribbing as the proversum and flexus, with fine ribbing starting from dorso-lateral tubercles or at the umbilical region, crossing the dorsal area without interruption. A single dorsal rib starts from each tubercle and is intercalated by 2 single short ribs.

*Remarks.*—The newly described species, *Dissimilites intermedius* sp. nov., differs from all other descendant species of the genus by its unique combination of tuberculation and ribbing style, i.e., the trituberculation on bifurcate ribs and intercalatory ribs throughout the entire shell.

The type material (20 specimens) of Dissimilites dissimilis derives from the collection of d'Orbigny (1842). D'Orbigny's (1842: 529, pl. 130: 4-7) syntypes of Hamites dissimilis were collected from Lower Barremian of Escragnolles, SE France. Busnardo (2006) revised the species, including the specimens by d'Orbigny (1842), figured two of d'Orbigny's specimens of D. dissimilis, designated one of them as a lectotype (Busnardo 2006: pl. 20: 8) and figured one paralectotype (Busnardo 2006: pl. 20: 9). However, the original figure in d'Orbigny (1842: 529, pl. 130: 4) and the lectotype of Busnardo (2006: pl. 20: 8) differ markedly. Busnardo's (2006) lectotype shows three to sometimes four strong, branching ribs extending from one umbilical tubercle on the retroversum, in contrast to single or bifurcate ribbing in d'Orbigny's (1842: pl. 130: 4-7) original drawing. The discrepancy most probably results from subjective interpretation and drawings of d'Orbigny (1842). All in all D. intermedius sp. nov. differs both from the specimens figured by d'Orbigny (1842) as well as by Busnardo (2006) by its trituberculation on the entire shell.

D'Orbigny (1852) described a similar specimen of Dissimi*lites dissimilis* (= *Hamulina dissimilis* in d'Orbigny 1852: 212) from Barremian beds near Escragnolles in southern France (Ravin de Saint-Martin). He (d'Orbigny 1842: 529) noted the specific features, 3-5 intercalatory ribs on the proversum "... ces côtes tuberculeuses se trouvent d'abord cinq petites côtes annulaires simples, réduites à quatre, puis à trois en approchant du coude.", no trituberculation on hook and retroversum, and mostly 2-3 ribs branching on retroversum from an internal tubercle. Spines shown in d'Orbigny's specimens (1852: pl. 3: 2, 3) are all restored on plates (also noted by Busnardo 2006) and spines could be detected on the holotype herein (Fig. 3A). The Puez specimens differ by trituberculation of flexus and retroversum on the single and branching ribbing. Additionally, the specimens from Puez exhibit only 1 (dominant) or 2 (rare) intercalatory ribs on the proversum (see rib formulae). The proversum from d'Orbigny's (1852) specimen is straighter than in the Puez specimens, where an angle is observable in the middle part of the proversum.

*Dissimilites subalternatus* (Breskovski, 1966) and *Dissimilites brestakensis* (Dimitrova, 1967) (= *Hamulina brestak-* *ensis* in Dimitrova 1967) show "looping" of trituberculate main ribs on proversum, only a single intercalatory rib on proversum and the disappearance of lateral tubercles on hook and retroversum. According to Avram (2001) and Klein et al. (2007), *D. brestakensis* (juvenile stage) is a junior synonym of *D. subalternatus*. *D. dissimilis* features "looping" primary ribs, different numbers of intercalatory ribs and the presence of lateral tubercles on the hook.

Following this synonymy, the specimen illustrated by Obata and Matsukawa (2007) from the Barremian of Japan as *Hamulina* aff. *brestakensis* has to be named *Dissimilites* aff. *subalternatus*. In Obata and Matsukawa (2007: fig. 2) the distribution of *D*. aff. *subalternatus* in Japan is indicated as Upper Barremian. *D. intermedius* sp. nov. from Puez differs from *D*. aff. *subalternatus* by trituberculation on the entire shell.

Avram (1976) described an "almost entire" specimen (i.e., without spire and early stages of proversum) of *Dissimilites dissimilis* from the Barremian of Romania (Comarniac, Outer Eastern Carpathians) showing 2–3 intercalatory ribs on the proversum, no lateral tubercles on hook and retroversum, and loss of internal tuberculation on adult shaft. The main differences to *D. intermedius* sp. nov. are the missing trituber-culation on hook and retroversum, higher numbers of intercalatory ribs (2–3 vs. 1–2 in *D. intermedius* sp. nov.) and more rounded tubercles on the hook in Avram's (1976) specimen.

The specimen of *Dissimilites* cf. *dissimilis* ("forme n° 1": Trigance, SE France) illustrated in Sarkar (1955: pl. 4: 17) is morphologically equivalent (although fragmented) to the herein presented form, and thus regarded synonymous. It shows the typical, weak trituberculation from proversum, over the hook and on the retroversum. All other individuals presented by Sarkar (1955) as *D*. cf. *dissimilis* ("forme n° 2") and *D*. *piveteaui* differ mainly in the missing trituberculation on hook and retroversum.

Major similarities exist with an individual from the Lower Barremian Nicklesia pulchella Zone of SE-France (Moriez, coupe Notre Dame) published as D. dissimilis by Ebbo et al. (2000), which is well comparable because of the fact that the specimen is almost complete. The specimen from Moriez shows the same buckle of the proversum as the specimen from Puez. Most probably, this feature relates to an ontogenetic change in growth direction. However, the new species also shows remarkable differences to the specimen from France, such as trituberculate ribs on the flexus and retroversal ribs, which are formed by irregularly alternating single and bifurcating ribs. In the Puez specimen, the distance between the pro- and retroversum, which are almost parallel, is half of the retroversum height. This contrasts with the specimen of Ebbo et al. (2000) showing an opening angle  $\beta$  of approx. 20° (vs. 4° in D. intermedius sp. nov.) and a distance between pro- and retroversum measuring 1 to 1.4 times the height of the retroversum. Similarly, the distance from the most juvenile part of the spire to the proversum is 18 mm in D. aff. dissimilites and only 5 mm in *D. intermedius* sp. nov. The spire in *D. intermedius* sp. nov. shows trituberculate ribbing, whereas Ebbo et al. (2000) reported no visible ornamentation on the spire, perhaps due to bad preservation. The latter authors noted the onset of ribbing at a diameter of approximately 11.7 mm in shell height. Both species show the same number of intercalatory ribs (1–2) on the proversum. The periostome is not preserved in the French specimen but marked by a final thickened rib in *D. intermedius* sp. nov.. The suture line is not visible in either of the type specimens, but the last chamber is visible immediately before the hook. The suture line, showing a trifid lobus, is visible in specimen PZO-CP56.

Dissimilites intermedius sp. nov. differs from *D. duboise* Vermeulen, 2009 by its extended trituberculation on the retroversum and on the final part of the proversum, and the absence of a bituberculation phase. Additionally, the stratigraphic occurrence of the two species is different, with the *Kotetishvilia compressissima–Moutoniceras moutonianum* to lower *Toxancyloceras vandenheckii* Zone for *D. intermedius* sp. nov. and the lower *Nicklesia pulchella* Zone for *D. duboise*.

The specimen from the Pizes de Puez (i.e., the same locality) figured by Costamoling and Costamoling (1994: 94, fig. 93) and identified as *Ancyloceras* shows exactly the same features as our specimens, especially the tiny trituberculation on the hook and additionally on the bifurcating and singular intercalatory ribbing on the retroversum.

Geographic and stratigraphic range.—Dissimilites intermedius sp. nov. is so far known only from the type locality and from Trigance (SE France) where it occurs in *M. moutonianum* Zone of upper Lower Barremian to the lower Upper Barremian (*T. vandenheckii* Zone).

#### Dissimilites trinodosum (d'Orbigny, 1852)

Figs. 3D, E, H, J-N, 4D, E, H, J-N.

- \*1852 Hamulina trinodosa; d'Orbigny 1852: 215, pl. 4: 1, 2, 3.
- 1889 Crioceras trinodosum (d'Orb) Uhl.; Haug 1889: 208, pl. 13: 3.
- 1966 Hamulina trinodosa d'Orbigny, 1851; Breskovski 1966: 83, pl. 4: 4.
- 1967 Hamulina dissimile d'Orbigny, 1851; Nagy 1967: 66, pl. 4: 2, 3.
- 1993 Dissimilites trinodosum (d'Orbigny); Stöhr 1993: 72, pl. 3:7.
- 1994 Dissimilites trinodosus (d'Orbigny); Avram 1994: pl. 15: 3, 4, 5.
- 2001 Dissimilites trinodosus (d'Orbigny); Avram 2001: pl. 1: 2, 3, 4.
- 2006 Acrioceras (Dissimilites) trinodosum (d'Orbigny); Lukeneder and Aspmair 2006: pl. 6: 8.
- 2006 Dissimilites (Dissimilites) trinodosum (d'Orbigny), Lukeneder and Aspmair 2006: pl. 6: 9.
- 2012 Dissimilites trinodosum; Lukeneder 2012b: text-fig. 6A.
- 2012 Dissimilites trinodosum; Lukeneder 2012b: text-fig. 8G.
- 2012 Dissimilites trinodosum; Lukeneder 2012b: text-fig. 8H.

*Holotype*: Specimen illustrated by d'Orbigny (1852: 215, pl. 4: 1, 2, 3); holotype by monotypy.

*Material.*—Two specimens derive from Puez, Early Barremian (top *Moutonianum moutonianum* Zone), log P1 bed 130 (NHMW 2011/0159/0026) and Late Barremian, P1 bed 146 (NHMW 2011/0159/0056). Two additional specimens derive from Late Barremian, P1 beds 169 and 172, top of the middle *T. vendenheckii* Zone (NHMW 2011/0159/0055 and NHMW 2005z0245/0036). Four additional specimens NHMW 2012/0003/0001–3, PZO-CP68) derive from La Stua, only regarded as Barremian in age because they were not collected bed-by-bed (Fig. 1). Eight predominantly flattened sculpture moulds with perfectly preserved sculpture were found within Barremian beds at Puez and La Stua. Most of the specimens are preserved on matrix and observable on one side only. In specimens NHMW 2011/0159/0036 and NHMW 2011/0159/0056 it was possible to prepare the proversum, flexus, and retroversum on both sides. The spire is not visible in these specimens, but early stages of proversum are preserved.

#### Measurements.—See Table 2.

*Description.*—Small, coiled acrioceratid shells with a weak torsion in spire. Whorl section is subrounded. Shell is tripartite as in *Dissimilites intermedius* sp. nov. The sculpture comprises simple, thick main ribs, single sharp ribs and bifurcating ribs. The suture line is not preserved. Ribs cross the venter without interruption except at the flexus (Figs.  $3D_4$ ,  $E_3$ ,  $J_4$ ,  $4D_3$ ,  $E_2$ ,  $J_4$ ), where a ventral depression or smooth, shallow furrow may occur. Main ribs on proversum are trituberculated. Specimens are mostly compressed; whorl breadths are approximate values due to compression.

*Spire.*—Only in specimen NHMW 2011/0159/0055 parts of the spire, connected to the proversum, are preserved. In the latter specimen the spire shows exclusively an alternation between sharp trituberculated ribs and single intermediate ribs (Figs. 3N, 4N). Ribs are slightly prosiradiate, and cross the venter without interruption. Trituberculate ribbing occurs throughout the specimen. Spire starts at 2.5 mm whorl height. Spire opens in a wide bow at a whorl height of 5 mm.

*Proversum.*—The proversum of specimen NHMW 2011/0159/ 0026 starts at a whorl height of 5 mm. It is distinctly curved up to the beginning of the flexus at a whorl height of 8 mm at the last trituberculate rib (Figs. 3D, 4D). The maximum length of the preserved proversum is 57 mm. Whorl height opens with a value  $\delta$  of approx. 4°. The proversum shows thick (up to 2 mm), single, trituberculate main ribs with only single, sharp, much thinner (one third of main ribs) intercalated ribs lacking tubercles. Ribbing is almost radial on the first part of the proversum and becomes more prosiradiate on the second part. Main ribs start from dorso-lateral tubercles. Ribs cross venter without interruption (Figs. 3D<sub>2</sub>, E<sub>2</sub>, J<sub>2</sub>, 4J<sub>2</sub>). On the dorsal part of the proversum, fine ribs occur from the height of dorso-lateral tubercles (twice as many ribs as on the flanks) and cross the dorsal area. Apparently, 2-3 of the later dorsal ribs end in a single dorso-lateral tubercle which gives rise to the thick trituberculated main ribs. Between these 2-3 tiny ribs the intercalatory ribs from the flank pass uninterruptedly throughout the entire shell. The last 12 main ribs (NHMW 2011/0159/0026; Figs. 3D, 4D) show fibulation from the lateral to the ventro-lateral tubercles; additionally, intercalatories appear up to the flexus. The last main rib also marks the beginning of the body chamber. As noted for D. intermedius sp. nov. above, all tubercles give rise to tiny spines, which are missing due to the preparation process.

Numerous specimens show similar fibulation of the last 5–12 main ribs on the proversum (NHMW 2011/0159/0056; Figs. 3E, 4E; PZO-CP68, Figs. 3H, 4H), a comparable curvature of the proversum, and sharp, trituberculated ribbing on the flexus. Specimen NHMW 2005z0245/0036 (Figs. 3J, 4J) is probably the most typical *D. trinodosum*. At 5mm whorl height, it shows the characteristic trituberculated (prosiradiate) main ribs, intercalated with single, fine ribs up to a whorl height of 8 mm, when the proversum starts with the appearance of solely trituberculate main ribs (8 in a row) up to the beginning of the flexus at 9 mm whorl height.

*Flexus.*—No thick main ribs occur. In specimen NHMW 2005z0245/0036 (Figs. 3J, 4J), ribs start from 6 elongated tubercles or clavae at the umbilical region and increase in thickness towards the external flank. Ribs may bifurcate (only once) from an umbilical tubercle. Ribs are trituberculate. Umbilical tubercles are less prominent than ventro-lateral and lateral ones. Ventro-lateral tubercles and lateral tubercles show the same size. Ribs are more prominent than in *D. intermedius* sp. nov. Lateral tubercles occur on mid flank. A shallow, almost smooth area occurs on the venter throughout the flexus (Figs. 3D<sub>4</sub>, E<sub>3</sub>, J<sub>4</sub>, 4D<sub>3</sub>, E<sub>2</sub>, J<sub>4</sub>). The flexus shows a maxium whorl height of 12 mm in NHMW 2005z0245/0036, 8 mm in NHMW 2011/0159/0026. The dorsal area is almost smooth, with tiny ribs as in *D. intermedius* sp. nov., crossing the dorsal area

Table 2. Measurements (in mm). Abbreviations: Dsp, spire diameter; Hf, whorl height at flexus; Hp1, whorl height at start of proversum; Hp2, whorl height at end of proversum; Hr, max. whorl height of retroversum; L, max. length; Lp, proversum length; Lr, retroversum length; min Hsp, minimal whorl height of spire; Pp1b, proversum part 1 breadth; Pp2b, proversum part 2 breadth; Spb, spire breadth;  $\alpha$ , angle of retroversum;  $\beta$ , angle formed by pro- and retroversum;  $\gamma$ , open angle on inner side of proversum;  $\delta$ , angle at end of proversum.

Inventory number	L	Lp	Lr	Dsp	minHsp	Hp1	Hp2	Hf	Hr	Spb	Pp1b	Pb2b	α	β	γ	δ
NHMW 2011/0159/0026	64	52	49	-	-	-	8	9	14	-	_	4	8	-	-	5
NHMW 2011/0159/0056	58	48	-	-	-	-	7	9	5	-	-	4	_	_	-	3
NHMW 2005z0245/0036	73	65	44	-	-	-	9	11	15	-	-	3	14	3	-	4
NHMW 2011/0159/0055	48	65	44	-	2	5	-	-	-	-	-	3	14	3	-	2
PZO-CP68	43	_	34	-	2	-	10	15	16	_	—	5	6	-2	-	3

without interruption. In NHMW 2011/0159/0026 (Fig. 3D, 4D), fibulation seems to appear also on the flexus between the first three ribs between lateral and ventro-lateral tubercles.

Retroversum.--The retroversum opens from the proversum almost subparallel with an angle  $\beta$  of approx. 2–3° and shows a maximum length of 54 mm (NHMW 2005z0245/0036) and 42 mm (NHMW 2011/0159/0026). The distance between proand retroversum is half of the whorl height of the proversum or one third of the retroversum whorl height. Shell height increases slowly by  $\alpha$  6° up to the aperture. The ribs are sharp and slightly retroradiate (NHMW 2005z0245/0036) to almost radial (NHMW 2011/0159/0026). Only the first trituberculated rib bifurcates from an umbilical tubercle (NHMW 2005z0245/0036; Figs. 3J, 4J), and the evolving branches cross the venter without interruption (Figs. 3D<sub>2</sub>, J<sub>2</sub>, 4D<sub>2</sub>, J<sub>2</sub>). After that bifurcating rib, only straight, single, trituberculated ribs occur (15 trituberculated ribs), which cross the venter and are of similar thickness (see rib formulae). Umbo-lateral tuberculation is virtually similar in size as the ventro-lateral tuberculation and the lateral tuberculation. The distance between the umbo and lateral tubercle is visibly larger than between lateral and ventro-lateral tubercles. Measured distances on the first single ribs are 5 mm vs. 3 mm. Lateral tuberculation is situated at approx. two thirds of the flank height. The position is similar as on the anterior flexus. The aperture seems to be marked by a collared last rib (NHMW 2011/0159/0026; Figs. 3D, 4D). The dorsal area shows the same ribbing pattern as proversum and flexus, with a single tiny rib starting from each dorso-lateral tubercle and crossing the dorsal area in a provers bow. These bows are intercalated by parallel bows of tiny ribs, starting on the umbilical area and crossing the dorsal area without interruption (NHMW 2005z0245/0036).

The ribbing pattern varies, as observable in NHMW 2011/0159/0026 (Figs. 3D, 4D). From the umbilical tubercle, the last bifurcating rib gives rise to a branch of two ribs; the front rib crosses the venter and ends at the following umbilical tubercle of the opposite flank, forming the back rib of that branch.

*Remarks.*—The described specimens of *Dissimilites trinodosum* differ from all other descendant species of this genus by their trituberculation on the proversum, flexus, and retroversum. Intercalatories are more regular on the proversum, with only one intercalatory throughout. Ribbing on the retroversum is single and trituberculate, except at the start of the retroversum, where 1–2 bifurcating ribs may occur. Tuberculation

in D. trinodosum shows almost the same strength throughout. D'Orbigny (1852) first described D. trinodosum (= Hamulina trinodosa in d'Orbigny 1852) from Barremian beds near Escragnolles in southern France (e.g., Ravin de Saint-Martin). As specific features, d'Orbigny (1852) noted the fewer intercalatory ribs on the proversum (1-2 in contrast to D. dissimilis with 3-5), trituberculation also on hook and retroversum, "... par la moitié moins de petites côtes intermédiaires aux côtes tuberculeuses de la partie droite..."(d'Orbigny 1852: 216), and single ribbing on retroversum. At the same morphological stages (i.e., ribbing style), the specimens from Puez differ only in size. The Puez individuals reach a max. length of 70 mm, which is almost half of the length of the French specimen (e.g., 20 cm given by d'Orbigny 1852). The spines illustrated in d'Orbigny's specimens (1852: pl. 4: 2, 3) are all restored in the drawings. Such spines are observable on the holotype of D. intermedius sp. nov. (Figs. 3A, 4A)

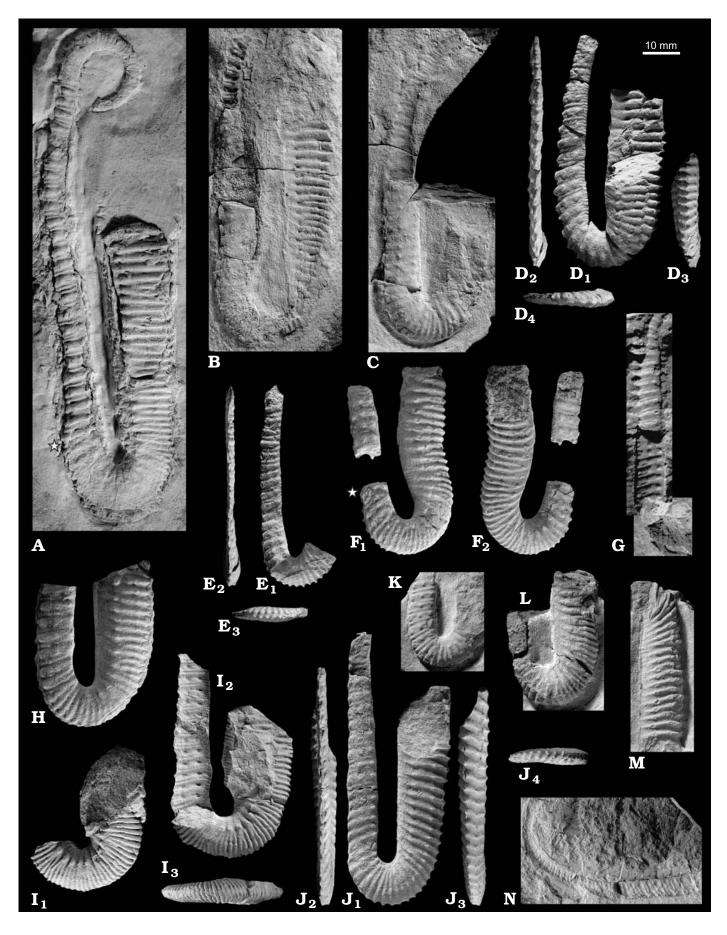
The specimens figured by Nagy (1967) from the Lower Barremian of the Gerecse Mountains (Hungary), named *Hamulina dissimile* (e.g., 1967: pl. 4: 2, 3), are here identified as *D. trinodosum*. A similar specimen was figured by Breskovski (1966) from Upper Barremian beds of Romania as *Hamulina trinodosa* (1966: pl. 4: 4). Specimens illustrated herein are also comparable to specimens figured by Avram (1994: pl. 15: 3, 4, 5) and Avram (2001: pl. 1: 2, 3, 4.) as *D. trinodosus* from Lower and Upper Barremian outcrops in Romania. Moreover, the specimens we present herein are very similar taxonomically to the material from Puez locality figured by Haug (1889: pl. 13: 3) and to a single specimen from La Stua illustrated by Stöhr (1993: pl. 3: 7).

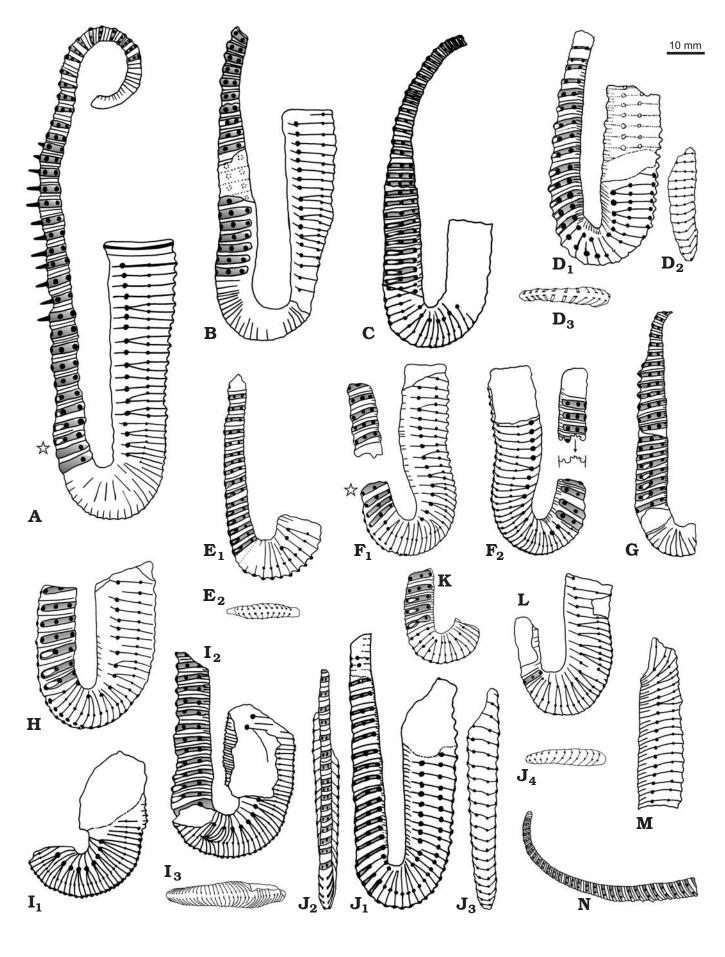
Geographic and stratigraphic range.—Dissimilites trinodosum is so far known from Europe (Italy, France, Bulgaria, Hungary, Romania) where it occurs from lowermost Lower Barremian (*Taveraidiscus hugii* Zone) to the upper Upper Barremian (*Toxancyloceras vandenheckii* and *Gerhardtia* sartousiana zones). D. trinodosum has been found in the Puez P1 section of the Trento Plateau (northern Italy, Southern Alps) from the top of the Moutoniceras moutonianum Zone up to the lower Toxancyloceras vendenheckii Zone.

## Taxonomic implications

The genus *Dissimilites* was first introduced by Sarkar (1954) in his revision of the Crioceratidae, unfortunately without figures of specimens. In his subsequent paper on heteromor-

Fig. 3. Upper Lower Barremian–lower Upper Barremian acrioceratid ammonoids *Dissimilites intermedius* sp. nov. (A–C, F, G, I) and *Dissimilites trino-*  $\rightarrow$  *dosum* (d'Orbigny, 1852) (D, E, H, J–N) from Dolomites (Southern Alps, Italy), Puez locality (*Moutoniceras moutonianum* Zone [A–D, G], *Toxancyloceras vandenheckii* Zone [E, N], unknown zone [F, H, I, J]), and La Stua locality (L–N). A–C. Holotype NHMW 2012/0002/0001 (A), paratypes NHMW 2012/0002/0003 (B), and NHMW 2012/0002/0002 (C); bed P7/-3.4 m (= bed P1/103) and bed P7/-3.2 m (= P1/105). D. NHMW 2011/0159/0026, bed P1/130, lateral view (D<sub>1</sub>), external views of proversum (D<sub>2</sub>), retroversum (D<sub>3</sub>), and flexus (D<sub>4</sub>). E. NHMW 2011/0159/0056, bed P1/146, lateral view (E<sub>1</sub>), external views of proversum (E<sub>2</sub>) and flexus (E<sub>3</sub>). F. PZO-CP56, lateral views (F<sub>1</sub>, F<sub>2</sub>); white star indicates the start of body chamber. G. Paratype NHMW 2011/0159/0027, bed P1/110. H. PZO-CP68, lateral view. I. NHMW 1886/0008/0017, lateral view (I<sub>1</sub>, I<sub>2</sub>), external view of flexus (I<sub>3</sub>). J. NHMW 2005z0245/0036, lateral view (J<sub>1</sub>), external views of proversum (J<sub>2</sub>), retroversum (J<sub>2</sub>), retroversum (J<sub>3</sub>), and flexus (J<sub>4</sub>). K. NHMW 2012/0003/0002, lateral view. L. NHMW 2012/0003/0003, lateral view. M. NHMW 2012/0003/0001, lateral view. N. NHMW 2011/0159/0055, bed P1/169, lateral view. All specimens in natural size, coated with ammonium chloride before photographing.





phic forms from south-east France, Sarkar (1955) reported different variations of *D. dissimilis* ("forme n° 1" and "forme n° 2"), *D. trinodosum*, and a new species, *D. piveteaui*.

*Dissimilites* is herein included in the subfamily Helicancyclinae Hyatt, 1894 (nom. transl. Casey 1961: 76. ex Helicancyclidae Hyatt, 1894: 565) within the recently defined family Acrioceratidae Vermeulen, 2004; this concept was also followed by Klein et al. (2007).

The problem of integrating the small heteromorph ammonite morphologies from the Barremian (e.g., Acrioceras, Dissimilites, Artareites) and the Aptian (e.g., Helicancylus, Toxoceratoides, Tonohamites) was discussed in detail by Casey (1961) and Aguirre Urreta (1986). Wright et al. (1996) proposed a scheme that incorporated the subfamily Helicancyclinae Hyatt, 1894 in the family Ancyloceratidae Gill, 1871. Casey (1961), Aguirre Urreta (1986), Ebbo et al. (2000), and Vašíček (2006) followed the expanded concept of the family Ancyloceratidae Gill, 1871, whereas Vermeulen (2004, 2006) and Klein et al. (2007) favoured the concept of the family Acrioceratidae Vermeulen, 2004. Bert (2009) included the genus Dissimilites into the family Helicancyclidae Hyatt, 1894 (Acrioceratidae Vermeulen, 2004). The evolutionary lineage from Acrioceras over Dissimilites to Toxoceratoides with a speculative position of the connective genera Artareites and Argvethites with indicated stratigraphy (Reboulet et al. 2009) is given in Fig. 6.

Vermeulen (2004) includes the genera *Acrioceras* Hyatt, 1900, *Dissimilites* Sarkar, 1955 and *Toxoceratoides* Spath, 1924 into one family, Acrioceratidae. The latter author notes that *Dissimilites* is most probably a descendant of *Acrioceras* and appears within the *Nicklesia pulchella* Zone. Starting from the *Kotetishvilia compressissima* Zone, two morphological forms develop: (i) shortened forms which are very similar to *Acrioceras* and (ii) forms with very long shaft and juvenile spire. Concerning the maximal size (e.g., shaft length), the open juvenile spire of the specimen presented herein as *D. intermedius* sp. nov. is similar to that seen in the second morphotype group of Vermeulen (2004). *Toxoceratoides* Spath, 1924 seems to be a descendant of the genus *Dissimilites* Sarkar, 1955 (Vermeulen 2004).

In his phylogenetic paper on heteromorphes Vermeulen (2006) strengthens the arguments for the monophyly of the family Acrioceratidae Vermeulen, 2004. Within this paper, he supports once again ranges given in Vermeulen (2004, 2005) and the evolutionary trend from *Acrioceras* Hyatt, 1900 to *Dissimilites* Sarkar, 1955, followed by *Toxoceratoides* Spath, 1924, and *Argvethites* Rouchadze, 1933.

Klein et al. (2007) placed *Dissimilites* Sarkar, 1954 within the family Acrioceratidae Vermeulen, 2004, whereas Wright et al. (1996) regarded *Dissimilites* as a subgenus of *Acrioceras* and additionally placed the subfamily Ancyloceratinae in the family Ancyloceratidae Gill, 1871. For a detailed discussion of the genus *Acrioceras* see also Delanoy and Bulot (1990) and Thomel et al. (1987, 1990).

Bert (2009) placed the genus *Dissimilites* within the family Helicancyclidae Hyatt, 1894 with its type genus Helicancylus Gabb, 1869. Bert (2009) discusses the members of Helicancyclidae in detail. Within this family, he proposes two lineages, i.e., Acrioceras-Dissimilites (see also Sarasin and Schöndelmayer 1902) and their possible descendant Artareites Bert, 2009 and Toxoceratoides-Helicancylus. Bert (2009) reported Dissimilites trinodosum from the Toxancyloceras vandenheckei Biozone (= T. vandenheckii Zone; subzone of *Barrancyloceras barremense*) from south-eastern France (e.g., Majastre section). Both species, D. dissimilis and D. trinodosum, differ from members of Artareites Bert, 2009 in having stronger and constant trituberculation on at least the proversum. Artareites exhibits only the fibulating rib type on the main ribs (e.g., ventral), but lacks trituberculation. Artareites first appears within the T. vandenheckii Zone (B. barremense Subzone). In contrast to Artareites, the initial spire can be shallowly helical in Acrioceras, Helicancylus, and Toxoceratoides. In the new species D. intermedius the initial spire apparently shows a shallow-helical morphology with trituberculation on the initial spire. Bert (2009) proposed Artareites as an intermediate form between Dissimilites and Toxoceratoides, and therefore Artareites as a descendant of Dissimilites. He strengthens his arguments by noting that in southern France the appearance of Artareites is coeval to the vanishing of Dissimilites. Bert (2009) described the insecurity in determining, from which form or ancestor Artareites derives: from D. trinodosum or from an unknown, earlier non-trituberculate form in the Toxancyloceras vandenheckii Zone (e.g., *Barrancyloceras barremense* Subzone).

The possible existence of dimorphic pairs is not discussed herein due the weakness of the fossil record with regard to that feature. Delanoy et al. (1995) and Bert (2009) discussed the possibility of a sexual dimorphic pair of *Emericiceras* (?macroconch) and *Dissimilites* or *Acrioceras* (?microconchs). Bert (2009) noted that, proposing an evolutionary lineage of *Acrioceras–Dissimilites*, no macroconch (e.g., *Emericiceras*) would show the transition from *K. compressissima* Zone up to *M. moutonianum* Zone. Vermeulen (2008) disagrees with the idea of dimorphic pairs in *Acrioceras* and *Emericiceras*, mostly due to different stratigraphic occurrences and absence of *Acrioceras* macroconchs in the Hauterivian and *Dissimilites* macroconchs in the Lower Barremian, and likewise for uppermost Barremian to Aptian descendants, such as *Toxoceratoi*-

← Fig. 4. Drawings of acrioceratid ammonoids *Dissimilites intermedius* sp. nov. (A–C, F, G, I) and *Dissimilites trinodosum* (d'Orbigny, 1852) (D, E, H, J–N) figured in Fig. 3. A–C. Holotype NHMW 2012/0002/0001 (A), paratypes NHMW 2012/0002/0003 (B), and NHMW 2012/0002/0002 (C); note spines. D. NHMW 2011/0159/0026, lateral view (D<sub>1</sub>), external views of retroversum (D<sub>2</sub>), and flexus (D<sub>3</sub>). E. NHMW 2011/0159/0026, lateral view (E<sub>1</sub>), external views of flexus (E<sub>2</sub>). F. PZO-CP56, lateral views (F<sub>1</sub>, F<sub>2</sub>); white star indicates the start of body chamber, note suture line in F<sub>2</sub>. G. Paratype NHMW 2011/0159/0027, lateral view. H. PZO-CP68; lateral view. I. NHMW 1886/0008/0017, lateral view (I<sub>1</sub>, I<sub>2</sub>), external view of flexus (I<sub>3</sub>). J. NHMW 2005z0245/0036, lateral view (J<sub>1</sub>), external views of proversum (J<sub>2</sub>), retroversum (J<sub>3</sub>), and flexus (J<sub>4</sub>). K. NHMW 2012/0003/0002, lateral view. L. NHMW 2012/0003/0003, lateral view. M. NHMW 2012/0003/0001, lateral view. N. NHMW 2011/0159/0055, lateral view. All drawings in natural size.

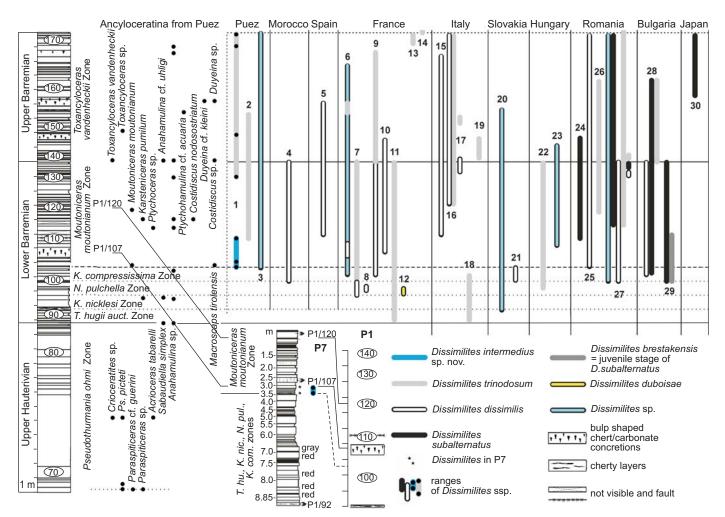


Fig. 5. Ranges of ancyloceratid ammonoids occurring in the Puez sections (P1 and P7) with indicated biostratigraphic ammonite zonation from Upper Hauterivian and Barremian. Distribution and range of *Dissimilites* species from Puez and all known occurrences worldwide. 1, this paper; 2, Lukeneder 2012; 3, Lukeneder and Aspmair 2006; 4, Company et al. 2008; 5, Company et al. 1995; 6, Vermeulen, 2005; 2006; 7, Vermeulen 1995, 1998; Ebbo et al. 2000; 8, Ebbo et al. 2000; 9, Busnardo 1963; 10, d'Orbigny 1842, 1852; Busnardo 2006; 11, Kilian 1910; 12, Vermeulen 2009; 13, Bert 2009; 14, Bert et al. 2010; 15, Uhlig 1887; 16, Haug 1887; 17, Haug 1889; 18, Stöhr 1993; 19, Cecca and Pallini 1994; 20, Uhlig 1883; 21, Vašíček 2006; 22, Nagy 1967; 23, Fülöp 1960; Fözy and Janssen 2009; 24, Simionescu 1998; 25, Avram 1976; 26, Avram 1994; 27, Avram 2001; 28, Breskovski 1966, 1975; 29, Dimitrova 1967; 30, Obata and Matsukawa 2007. Note indicated biostratigraphic ammonite zonation. Ammonite occurrences and ranges are marked by solid black circles. Bold horizontal line, stage boundary; dashed horizontal lines, zonal boundaries; dotted horizontal lines, subzonal boundaries. Taxonomic abbreviations: *Ps., Pseudothurmannia; T., Taveraidiscus; K., Kotetishvilia; N., Nicklesia; hu., hugii; nic., nicklesi; pul., pulchella; com., compressissima*.

*des* (see also Vermeulen and Bulot 2007). Notably, in most entirely preserved specimens of *Acrioceras* a helical spire stage is present. In *Dissimilites intermedius* sp. nov. this character is visible in very few specimens, which are entirely preserved. Nevertheless, this character demonstrates that *Dissimilites* is a member of the family Acrioceratidae Vermeulen, 2004 and the subfamily Helicancyclinae Hyatt, 1894.

This is of importance if we consider *Dissimilites* as a possible ancestor of *Toxoceratoides*, which, to the opinion of numerous authors such as Casey (1959, 1961), shows a shallow helical juvenile spire, similar to *Acrioceras* (Hyatt 1894; Wright et al. 1996), *Helicancylus* (Gabb 1869; Kakabadze 1981; Aguirre Urreta 1986), and *Argvethites* (Rouchadze 1933). These phylogenetic aspects were also discussed by Vermeulen (2005), who suggested an evolutionary lineage from Upper Hauterivian–Lower Barremian *Acrioceras* over

Lower–Upper Barremian *Dissimilites* to uppermost Barremian–Aptian *Toxoceratoides*, giving rise to Aptian *Argvethites* (Fig. 6). Vermeulen's (2008) *Dissimilites duboise* co-occurs with the last members of the genus *Acrioceras*, i.e., *Acrioceras terveri* Astier, 1851.

Vermeulen (2005) points out the known ranges as Upper Hauterivian (*Plesiodiscus ligatus* Zone) up to Lower Barremian (*Nicklesia pulchella* Zone) for *Acrioceras* and *N. pulchella* Zone to *T. vandenheckii* Zone (e.g., = lower *Heinzia sayni* Zone of Vermeulen, 2005; see Reboulet et al. 2009) for *Dissimilites*. Vermeulen (2005) gave a range for *Argvethites*, spanning from *Gerhardtia sartousiana* Zone into the Aptian and *Imerites giraudi* Zone (e.g., *I. giraudi* to *Martelites sarasini* Zone of Vermeulen 2005; see Reboulet et al. 2009). Reboulet et al. (2009) incorporated the *M. sarasini* Zone as the middle subzone into the *I. giraudi* Zone. Vermeulen (2005) noted and figured (e.g., by dashed range lines) that several first or last occurrences of these genera are still uncertain, making the phylogenetic junction points arguable.

It seems that there was a trend in general shell elongation within the *Dissimilites*; from *D. dissimilites* towards *D. trinodosum*. This was accompanied by increasing tuberculation of the hook and retroversum. A second lineage shows a reduction of size concurrently with an increasing curvation of the proversum; the latter developed towards *Toxoceratoides* (Fig. 6). There is agreement that *D. trinodosum* is the descendant of *D. dissimilis* (Vermeulen 2004; Bert 2009). Possible implication of ontogenetic heterochronies can not be considered due to the scarcity of material (14 specimens from Puez), which prevents from any reliable performance of statistics to clarify intraspecific variability.

## Geographic and stratigraphic occurrences of *Dissimilites*

In the following paragraphs, the occurrences of the species of the genus *Dissimilites* are discussed according to geographical regions, from East (Morocco) to West (Japan). A biostratigraphic and palaeogeographic synopsis is given in Figs. 5–7. Within each geographic area, occurrences and references are arranged in the order of their publication dates. For those localities described by several authors interrelationships are established.

Africa.—Company et al. (2008) reported *Dissimilites dissimilis* from the upper Lower Barremian of the western High Atlas of Morocco (Sidi M'Bark section, Imsoune section, Addar section, Tamri section).

In summary, *D. dissimilis* occurs from the lowermost *K. compressissima* Zone up to the top of the *M. moutonianum* Zone (= middle *Coronites darsi* Zone in Company et al. 2008; see Reboulet et al. 2009) in this region (Fig. 7).

Walchendorf et al. (1967) reported *Acrioceras dissimilis* from the Lower Aptian of South Mozambique (e.g., Chalala) and stated that *Acrioceras* Hyatt, 1900 is the senior synonym of *Dissimilites* Sarkar, 1955, an opinion that was followed by Arkell (1957), but is rejected herein. In contrast to Walchendorf et al. (1967) the African specimens were included in *Toxoceratoides krenkeli* by Klein et al. (2007), which is followed herein.

**Europe**.—Occurrences of *D. dissimilis* from southeastern Spain (e.g., Subbetic Domain) were reported from the *K. compressissima* Zone up to the *T. vandenheckii Zone* (up to lower *Barrancyloceras barremense* Subzone) by Company et al. (1995).

D'Orbigny (1852) described two species: *D. dissimilis* (= *Hamulina dissimilis* in d'Orbigny 1852) and *D. trinodosum* (= *Hamulina trinodosa* in d'Orbigny 1852), from Barremian beds near Escragnolles in southern France (Ravin de Saint-Martin). From Ravin de Saint-Martin (section 062), Vermeulen (1996) described Barremian faunas with D. trinodosum and D. dissimilis (e.g., both bed 18), which he included to the M. moutonianum Zone (lowermost Coronites darsi Zone in Vermeulen 1995, 1996, 1998; M. moutonianum Zone in Reboulet et al. 2009). Busnardo (2006) revised d'Orbigny's (1842) type material of Hamites dissimilis from Escragnolles (Lower Barremian; SE France). Busnardo (2006: pl. 20: 8, 9) figured two specimens as D. dissimilis, which was followed by Klein et al. (2007). We herein do not follow this assignment, due to the enormous difference in the ribbing pattern of the two specimens. The lectotype designated by Busnardo (2006: pl. 20: 8) shows 2–3 intercalatories between main trituberculate ribbing and strong bi-, tri- and sometimes even quadrifurcating ribbing on the retroversum, while the paralectotype (Busnardo 2006: pl. 20: 9) shows no or only a single intercalatory rib between strongly looping primaries on the proversum and only single to bifurcating ribbing on the retroversum. D'Orbigny (1852) figured a somewhat idealized specimen from his collection of Escragnolles, but, as noted by Busnardo (2006), shows looping primary ribs on the retroversum.

Lory and Sayn (1895) mentioned a specimen as *Crioceras trinodosum* from SE France (Châtillon-en-Diois), which is based on the accompanying fauna, most probably Late Barremian in age. Kilian (1910) and Paquier (1900) mentioned *Crioceratites trinodosum* from the Lower Barremian of SE France (Drôme and Basses Alpes, Diois, and Baronnies area). The specimen from the Lower Aptian of southeast France (Homme-d'Armes) that was figured by Kilian and Reboul (1915: 75, pl. 1: 4) as *Crioceras dissimile* (sub. *Hamulina*), was interpreted as a misclassification by Klein et al. (2007), a standpoint followed herein. It may rather be incorporated into the genera *Anahamulina* Hyatt, 1900, *Amorina* Vermeulen, 2005, or *Duyeina* Vermeulen, 2005.

Sarkar (1955) reported several specimens from south-east France: D. cf. dissimilis ("forme n° 1" and "forme n° 2"), D. trinodosum and a new species Dissimilites piveteaui. He suggests D. cf. dissimilis ("forme n° 1") to be of Barremian age, D. cf. dissimilis ("forme n° 2") of Late Barremian, and D. piveteaui of Barremian and Early Barremian age.

The first individual of *D*. cf. *trinodosum* from the Barremian stratotype in Angles (SE France) was reported by Busnardo (1963) from the Upper Barremian interval (beds 152–161; in Busnardo 1963: 106–107). Additionally, a find of *D*. *dissimilis* was reported as Lower Barremian from a nearby, more western section (Busnardo 1963: 108, cote 981, niveau 5). Considering recent literature (Vermeulen 2005), the bed numbers given by Busnardo (1963: beds 152–161), would correspond to the *Hainzia sayni* Zone (Vermeulen 2005) and thus to the adequate lower *T. vandenheckii* Zone (Reboulet et al. 2009). Bed 152 corresponds to the last finding of *Dissimilites* at Angles by Vermeulen (2005: 154, bed 152) identified as *Dissimilites* sp.

Vermeulen (2005) reported the occurrence of *Dissimilites* from the Barremian stratotype over an interval of approx. 20 m from the upper *K. compressissima* Zone (e.g., bed 120 with *Dissimilites* sp.) up to the middle *T. vandenheckii* Zone (e.g.,

= lower Heinzia sayni Zone of Vermeulen, 2005; see Reboulet et al. 2009). He noted D. cf. dissimilites from the lower part of the *M. moutonianum* Zone (e.g., bed 126, lower Coronites darsi Zone of Vermeulen 2005; see Reboulet et al. 2009) and D. cf. trinodosum from up to the lower T. vandenheckii Zone (e.g., bed 144, Holcodiscus uhligi Zone of Vermeulen 2005; see Reboulet et al. 2009). Vermeulen (2008) erected the most recent species of the genus Dissimilites from southeast France (Clos de Barral), namely Dissimilites duboise Vermeulen, 2009. It occurs in the lower parts of the N. pulchella Zone, representing the oldest known species of Dissimilites so far. From the Nauvin section (SE France, Moustiers Sainte-Marie area), Bert et al. (2010) described D. trinodosum from his Gassendiceras alpinum Subzone (= equivalent to the Barrancyloceras barremense Subzone; after Reboulet et al. 2009), corresponding to the T. vandenheckii Zone.

Uhlig (1887: 100) described a specimen of Crioceras dissimile in his paper on Neocomian fossils of the Gardenazza Plateau, northern Italy, which includes also the Puez area. Haug (1887: 260) listed only in his synopsis the occurrence of Crioceras trinodosum and Crioceras dissimile from the "Upper Neocomian" of the Puez area in northern Italy (= "Oberneokom von der Puezalpe"). More detailed descriptions of Crioceras trinodosum and Crioceras dissimile were given by Haug (1889) from the same locality (Puez area) and attributed to the same age, Upper Neocomian (e.g., Hauterivian–Barremian). The figured specimen of C. trinodosum (Haug 1889: pl. 13: 3) shows clear trituberculation on the proversum, hook, and retroversum. It also shows only a single intercalatory rib on the proversum, and only single, straight, sharp ribs between the main ribs shortly after the hook. The latter specimens are all "rock-fall samples" which cannot be assigned to any particular horizon.

Rodighiero (1919: 306) published a list in which he included *Crioceras mojsisovicsi* n. sp. from the Barremian of Colle Muntiella (= Col de la Muntejela; Dolomites, N. Italy), which is located on the Gardenaccia Plateau near Puez. This specimen is as attributed to *D. dissimilis* herein. Reithofer (1928) summarized the faunistic knowledge at that time from Puez, including *Crioceras dissimile* and *Crioceras trinodosum*, which correspond to findings of Uhlig (1887) and Haug (1887).

Stöhr (1993) reported three specimens from the Lower Barremian of northern Italy (La Stua) and figured a single *D. trinodosum* (Stöhr 1993: pl. 3: 7; his nr. 91.1). The latter author gives a range from the Hauterivian–Barremian boundary up to the Upper Lower Barremian for *D. trinodosum* at La Stua (Stöhr 1993: fig. 1). The first occurrence of *D. trinodosum* at the Hauterivian–Barremian boundary proposed by Stöhr (1993) is most probably based on a species misidentification. The short range at the top is probably due to poor outcrop situation or unsuccessful sampling. Stöhr (1993) described and figured a specimen showing trituberculate ribbing on both pro- and retroversum, suggesting assignment to *D. trinodosum*. Additional specimens are preserved only as internal molds with trituberculated proversa (Stöhr 1993); this is insufficient for species determination and therefore for stratigraphic implications.

Cecca and Pallini (1994) reported *D*. cf. *trinodosus* from the Gorgo a Cerbare section within the Umbria-Marche Apennines of Italy. The latter authors do not figure the specimen but refer it to the lowermost upper Barremian, the *T. vandenheckii* Zone.

Lukeneder and Aspmair (2006) identified several specimens from the Puez area as *Acrioceras (Dissimilites) dissimilis* (pl. 6: 7), *Acrioceras (Dissimilites) trinodosum* (pl. 6: 8), and *D. (Dissimilites) trinodosum* (pls. 6: 9, 7: 1). All were suggested to be typical Barremian forms (Lukeneder and Aspmair 2006).

Lukeneder (2012b) described specimens of *D. trinodo*sum from northern Italy with precise stratigrafical positions (?*N. pulchella*, *M. moutonianum*, and *T. vandenheckii* zones; Lukeneder 2012b: figs. 6A, B, 8H, respectively). All these specimens have been reevaluated and incorporated in the synonymy lists herein (see above). As discussed above, the biostratigraphic zonation has changed due to new findings. No members of *Dissimilites* have been found in strata older than the *M. moutonianum* Zone at Puez.

Uhlig (1883: pls. 25: 2, 3, 25: 4a-d, respectively) figured D. dissimile (= Crioceras dissimile in Uhlig 1883) from Niedek, Slovakia and from Escragnolles, France, and mentioned D. trinodosum (i.e., Crioceras trinodosum in Uhlig 1883: 270) from Kozy in Poland. Klein et al. (2007) agreed with the classification of the Slovakian specimens but incorporated Uhlig's (1883) specimen from France into the synonymy of D. subalternatus without providing any reasons. This was probably because it showed only a single intercalatory rib on the retroversum and no lateral tubercle on the hook and retroversum, and is therefore plausible. No precise stratigraphic range can be given for Uhlig's (1883) specimens due to the sampling method (not bed-by-bed) and lack of knowledge of the exact localities. Based on the accompanying fauna, however, the stratigraphic range of Dissimilites in Uhligs's (1883) material from Slovakia can be estimated with upper Lower Barremian to lower Upper Barremian.

Vašíček (2006) reported a single *D. dissimilis* specimen from the Lower Barremian *K. compressissima* Zone of the Central Western Carpathians (Slovakia, Butkov Quarry). The specimen more closely resembles *D. subalternatus* according to looping, tuberculation style, and intercalatories.

In his synopsis on the Lower Cretaceous of the Bakony Mountains Fülöp (1960) noted Barremian specimens of *"Hamulina dissimilis* ORB." from several drill cores (e.g., Sümeg, Városlőd, and Szenta). Nagy (1967: 66, pl. 4: 2, 3) reported specimens of *Hamulina dissimile* from the Lower Barremian of the Gerecse Mountains (Hungary). Fözy and Janssen (2009) reported *Dissimilites* sp. from the Lower Barremian of the Gerecse Mountains, ranging from the upper *N. pulchella* Zone to the *T. vandenheckii* Zone in the Bersek Quarry (sections A–E).

Simionescu (1898: pl. 4: 8a, b) figured a Barremian ("unknown horizon" in Avram 1976) specimen as *Crioceras dis*- *simile* from the Murierii Valley, Romania and reported two specimens of *Crioceras trinodosum* (not figured; same locality). The first specimen was refigured by Avram (2001: pl. 2: 21a, b), who noted that this specimen is most probably a *D*. aff. *subalternatus*, an opinion followed by Klein et al. (2007) by incorporating it in the synonymy of *D. subalternatus*. That concept is followed herein, due to strong zigzagging between umbilical and ventro-lateral tubercles, looping of ribbing on the proversum and the fewer intercalatories compared with *D. dissimilis*. In his synopsis of Lower Cretaceous ammonite faunas, Simionescu (1900: 133, 136 respectively) also mentioned *Crioceras dissimile* from France, Spain, Italy, the Czech Republic, and Romania, and *Crioceras trinodosum* from Italy, France, the Czech Republic, and Romania.

Avram (1976) described an "almost entire" specimen (e.g., without spire and early stages of proversum) of *D. dissimilis* from the Barremian of Romania (Comarniac, Outer Eastern Carpathians). Avram figured three specimens of *D. trinodosus* (1994: pl. 15: 3–5) from the Lower Barremian (Temeneacia valley) and lower Upper Barremian (Orsova-Moldova Noua and Tiganski valley) outcrops near Svinita in the South Carpathians of Romania.

Avram (2001) reported several specimens of the Barremian genus *Dissimilites* and Aptian genus *Toxoceratoides* from Romania. The latter author proposed a stratigraphic range from the *K. compressissima* Zone to *Gerhardtia sartousiana* Zone for *Dissimilites*. Avram (2001) reported *D. dissimilis* (d'Orbigny, 1842) as characteristic for the top of the Lower Barremian, whereas *D. trinodosum* and *D. subalternatus* were found at the top of the Lower Barremian and in the lower Upper Barremian. Avram (2001) stated that the spire of *Dissimilites* is not trituberculate as it is in *Acrioceras*. This is incorrect as seen in our holotype, which shows spire trituberculation.

Breskovski (1966, 1975) in his paper on zonal and subzonal subdivisions of the Barremian in Bulgaria provided also ranges of species belonging to Dissimilites: upper Lower Barremian (e.g., Holcodiscus perezianus Subzone after Breskovski 1975) for D. dissimilis (= Hamulina dissimilis in Breskovski 1975), upper Lower Barremian to lower Upper Barremian (e.g., Holcodiscus perezianus-Melchiorites cassidoides Subzones after Breskovski 1975) for D. subalternatus (= Hamulina subalternata in Breskovski 1975), lower Upper Barremian (e.g., M. cassidoides Subzone after Breskovski 1975) for D. brestakensis (= Hamulina brestakensis in Breskovski 1975), and lower Upper Barremian (e.g., M. cassidoides Subzone according to Breskovski 1975) for D. trinodosus (= Hamulina trinodosa, Breskovski 1966). To the opinion of Avram (2001), D. brestakensis represents a juvenile stage of D. subalternatus, an opinion which has also been adopted by Klein et al. (2007). Accordingly, the material of D. subalternatus from Bulgaria (Breskovski 1975) ranges from upper Lower Barremian to lower Upper Barremian. Dimitrova (1967) noted the occurrence of Dissimilites brestakensis (=D. subalternatus) in the Upper Barremian of Bulgaria (e.g., Brestag near Varna). Dimitrova (1967) described the co-occurrence with "Pulchellia" compressissima (= Kotetishvilia armenica

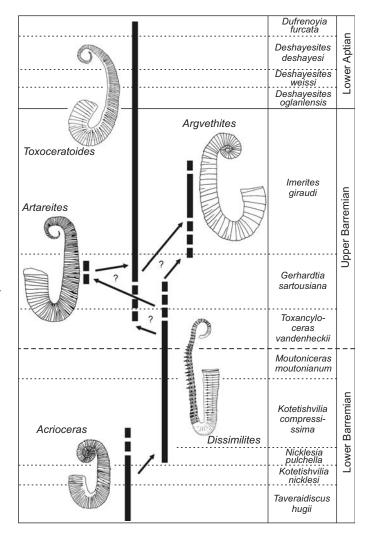


Fig. 6. Evolutionary line from *Acrioceras* over *Dissimilites* to *Toxoceratoides* with speculative position of the connective forms *Artareites* and *Argvethites*. Indicated stratigraphy to the right is in accordance with Reboulet et al. (2009). Bold horizontal line, stage boundary; dashed horizontal lines, substage boundaries; dotted horizontal lines, zonal boundaries.

after Vermeulen and Klein 2006), which most probably shows a stratigraphic range of upper Lower Barremian (e.g., base of *M. moutonianum* Zone; see Reboulet et al. 2009).

The specimen described and figured as *Crioceratites* aff. *dissimile* by Kazansky (1914: 42, pl. 1: 12, 13) from Lower Aptian beds (N 4b; beds with *Hoplites deshayesii* in Kazansky 1914; = *D. deshayesi* Zone after Reboulet et al. 2009) of Daghestan in the Northern Caucasus (e.g., Zoudakar; southern Russia) does not belong to *D. dissimilis*. As noted by Klein et al. (2007: 104) the Caucasian specimen is most probably a member of *Toxoceratoides* Spath, 1924. Reasonable doubts are with the identification of *Dissimilites* from the Aptian of Transcaucasia and the Caucasus to the north, reported by Kotetichvili (1983). The material probably also corresponds to the genus *Toxoceratoides*.

**Asia**.— The specimen published by Obata and Matsukawa (2007: figs. 5, 7) from the Barremian of Japan as *Hamulina* 

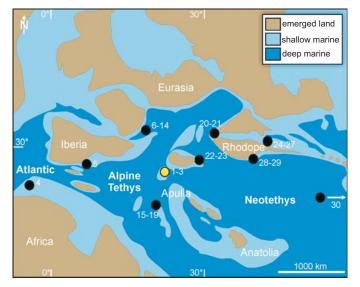


Fig. 7. Paleogeographical map of the Mediterranean area during the Early Barremian based on Dercourt et al. (1993) and Barrier and Vrielynck (2008). Numbers correspond to Fig. 5 and indicate localities of *Dissimilites* occurrences. Yellow spot shows locality of Puez, other localities are in black. 1–3, Puez; 4, Morocco; 5, Spain; 6–14, France; 15–19, Italy; 20, 21, Slovakia; 22, 23, Hungary; 24–27 Romania; 28, 29, Bulgaria; 30, Japan.

aff. *brestakensis* is more likely *D*. aff. *subalternatus*. The distribution of ammonites provided by Obata and Matsukawa (2007: fig. 2), points to an Upper Barremian occurrence of their *D*. aff. *subalternatus*.

## Conclusions

A new species within the heteromorph genus Dissimilites Sarkar, 1954 is proposed, based on comprehensive material, including the most complete specimen known so far (spire, proversum, flexus, retroversum, peristome). Dissimilites intermedius sp. nov. (upper Lower Barremian, Moutoniceras moutonianum Zone) shows intermediate morphological features of the ancestral Dissimilites dissimilis and its descendant Dissimilites trinodosum. The new species combines the ribbing style of *D. dissimilis* (bifurcating ribs intercalated with simple ribs) with the tuberculation style of D. trinodosum (trituberculation on entire shell). We could also confirm the previously observed trend in shell length decrease in Dissimilites from the upper Lower Barremian (M. moutonianum Zone) to lower Upper Barremian (Toxancyloceras vandenheckii Zone). This size reduction is accompanied by a simultaneous increase of shell curvation during the same time interval. The shallow-helical spire with its consistent ornamentation with trituberculated main ribs shows a close relationship to the assumed ancestor Acrioceras, whereas the increasing curvation of the younger forms transfers to the descendant Toxoceratoides. Thus, the presented data support the concept of a direct evolutionary lineage from Acrioceras over Dissimilites to Toxoceratoides. The evolutionary trend in Acrioceras, Dissimilites, and Toxoceratoides in the Upper Barremian (e.g., Gerhardtia sartousi*ana* and *Imerites giraudi* zones) is strongly supported by the connective genera *Artareites* and *Argvethites*.

Additional ammonite findings show that the M. moutonianum Zone has a longer range than expected and that the Taveraidiscus hugii, Kotetishvilia nicklesi, Nicklesia pulchella, and Kotetishvilia compressissima zones are more condensed than previously thought. The fauna from the Puez locality comprises only two genera of the family Acrioceratidae: Acrioceras and Dissimilites. The genus Dissimilites starts at Puez with D. intermedius within the M. moutonianum Zone. The survey of the literature data revealed that Dissimilites appeared during the N. pulchella Zone, represented by D. duboise, and later replaced by D. dissimilis. Two morphological forms develop in the Kotetishvilia compressissima Zone-the first one with a shortened spire, very similar to Acrioceras, and the second one with a very long shaft and juvenile spire, as seen in the new species D. intermedius. Later on, in the M. moutonianum Zone, D. intermedius sp. nov. is replaced by D. subalternatus. Besides, D. trinodosum occurs, which seems to be ancestral to the genus Toxoceratoides.

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## References

- Aguirre Urreta, M.B. 1986. Aptian ammonites from the Argentinian Austral Basin. The subfamily Helicancylinae Hyatt, 1894. *Annals of the South African Museum* 96: 271–314.
- Arkell, W.J. 1957. Mollusca 4 (Cephalopoda, Ammonoidea). In: R.C. Moore (ed.), Treatise on Invertebrate Paleontology L, 1–490. University of Kansas Press, Lawrence.
- Astier, M.J.-E. 1851. Catalogue descriptif des Ancyloceras appartenant à l'étage Neocomien d'Escragnolles et de Basses-Alpes. Annales de la Sociéte Nationale d'Agriculture, d'Histoire Naturelle et des Arts Utiles de Lyon 3 (2): 435–456.
- Avram, E. 1976. Les fossils du flysch éocrétecé et des calcaires tithoniques des hautes vallées de la Doftana et du Tirlung (Carpates Orientales). Mémoires de l'Institut de Géologie et de Géophysique Bucarest 24: 5–74.
- Avram, E. 1994. Lower Cretaceous (Valanginian–Early Aptian) ammonite succession in the Svinita region (SW Rumania). *Géologie Alpine, Hors Série* 20: 113–167.

- Avram, E. 2001. Representatives in Romania of the genera *Dissimilites* Sarkar, 1954 and *Toxoceratoides* Spath, 1924 (Ancyloceratina, Ammonoidea). *Acta Palaeontologica Romaniae* 3: 23–30.
- Baccelle, L. and Lucchi-Garavello, A. 1967. Ammonite dei livelli Cretacici di La Stua (Cortina d'Ampezzo). Annali dell' Università di Ferrara, Nuova Serie 4 (9): 117–166.
- Barrier, E. and Vrielynck, B. 2008. Palaeotectonic Maps of the Middle East, CCGM, 14 maps. Commission for the Geologogical Map of the World/UNESCO, Paris.
- Bert, D. 2009. Description de Artareites landii nov. (Ammonoidea) du Barrémien supérieur de Majastre (Sud-Est de la France) et discussion sur les Helicancylidae Hyatt, 1894. Annales de Paléontologie 95: 139–163.
- Bert, D., Busnardo, R., Delanoy, G., and Bersac, S. 2010. Problems in the identity of "Crioceras" barremense Kilian, 1895 (Ancyloceratina, Late Barremian). Carnets de Géologie, Notebooks on Geology Article 2010/01 (CG2010\_A01): 1–17.
- Breskovski, S. 1966. Biostratigraphie du Barrémien au sud du village de Brestak, dans la région de Varna. *Trudove V'rkhu Geologiyata na Balgariya, Seriya Palaeontologiya* 8: 71–121.
- Breskovski, S. 1975. Les zones et sous-zones ammonitiques dans l'étage Barrémien en Bulgarie du Nord-Est. *Geologica Balcanica* 5 (2): 47–66.
- Busnardo, R. 1963. Le stratotype du Barrémien. Lithologie et macrofaune. In: Colloque sur le Crétacé inférieur (Lyon, Septembre 1963). Mémoire du Bureau de Recherches Géologiques et Minières, Orléans 34: 101–116.
- Busnardo, R. 2006. Dissimilites dissimilis. In: H. Gauthier, R. Busnardo, R. Combémorel, G. Delanoy, J.-C. Fischer, F. Guérin-Franiatte, B. Joly, W.J. Kennedy, J. Sornay, and H. Tintant (eds.), Révision critique de la Paléontologie française d'Alcide d'Orbigny, Vol. IV, Céphalopodes crétacés, 1–292. Backhuys Publisher, Leiden.
- Casey, R. 1959. A monograph of the ammonoidea of the Lower Greensand—Part I. Palaeontographical Society (1958): 1–44.
- Casey, R. 1961. A monograph of the ammonoidea of the Lower Greensand—Part II. Palaeontographical Society (1960): 45–118.
- Cecca, F. 1998. Early Cretaceous (pre-Aptian) ammonites of the Mediterranean Tethys: palaeoecology and palaeobiography. *Palaeogeography*, *Palaeoclimatology*, *Palaeoecology* 138: 305–323.
- Cecca, F. and Pallini, G. 1994. Latest Hauterivian–Berremian Ammonite biostratigraphy in the Umbria-Marche Apennines (Central Italy). *In*: L.G. Bulot, M. Argot, and H. Arnaud (eds.), Lower Cretaceous Cephalopod biostratigraphy of the Western Tethys. *Géologie Alpine, Mémoire Hors Série* 20: 205–217.
- Cita, M.B. 1965. Jurassic, Cretaceous and Tertiary Microfacies from the Southern Alps (Northern Italy). 99 pp. E.J. Brill, Leiden.
- Cita, M.B. and Pasquaré, G. 1959. Osservationi micropaleontologiche sul Cretaceo delle Dolomiti. *Rivista Italiana di Paleontologia e Stratigrafia* 65: 385–443.
- Company, M., Sandoval, J., and Tavera, J.M. 1995. Lower Barremian ammonite biostratigraphy in the Subbetic Domain (Betic Cordillera, southern Spain). *Cretaceous Research* 16: 243–256.
- Company, M., Sandoval, J., Tavera, J.M., Aoutem, M., and Ettachfini, M. 2008. Barremian ammonite faunas from the western High Atlas, Morocco—biostratigraphy and palaeobiogeography. *Cretaceous Research* 29: 9–26.
- Costamoling, H. and Costamoling, W. 1994. *Fossilien des Gardertales*. 111 pp. Verlagsanstalt Athesia Ges. m.b.H., Bozen.
- Cuvier, G. 1798. *Tableau élémentaire de l'histoire naturelle de Animaux*. 710 pp. De l'Institut National de France, Paris.
- Dercourt, J., Ricou, L.E., and Vrielynck, B. 1993. Atlas Tethys Palaeoenvironmental Maps. 307 pp. (with 14 maps). Gauthier-Villars, Paris.
- Delanoy, G. 2003. *Toxancyloceras gen. nov*. (Ammonoidea, Ancyloceratina) un nouveau genre du Barrémien supérieur. *Annales du Muséum d'Histoire naturelle de Nice* 18: 1–19.
- Delanoy, G. and Bulot, L. 1990. Révision des types et figures des collections Matheron et Reynes. 3. Genres: Acrioceras, Heteroceras, Kutatissites et Pseudocrioceras (Ancyloceratina, Cephalopoda). Mésogée 50: 15–21.
- Delanoy, G., Ropolo, P., Magnin, A., Autran, G., Poupon, A., and Gonnet, R. 1995. Sur les dimorphisme chez les Ancyloceratina (Ammonoidea)

du Crétacé inférieur. *Comptes Rendus de l'Academie des Sciences de Paris* 321: 537–543.

- Dimitrova, N. 1967. Les fossiles de Bulgarie IV. Crétacé inférieur. Cephalopoda (Nautiloidea et Ammonitina). 424 pp. Académie Bulgare des Sciences, Sofia.
- D'Orbigny, A. (1840–1842). Paléontologie francaise. Description zoologique et géologique de tous les animaux mollusqes et royonnés fossiles de France. Terrains Crétacés. I. Céphalopodes. 662 pp, 148 pls. Alcide d'Orbigny & Arthus-Bertrand, Paris.
- D'Orbigny, A. 1852. Notice sur la genre Hamulina. Journal de Conchyologie 3: 207–228.
- Ebbo, L., Poupon, A., and Delanoy, G. 2000. Nouvelles données paléontologiques sur le genre Dissimilites Sarkar, 1954. Annales de Paléontologie 86: 91–99.
- Fözy, I. and Janssen, N.M.M. 2009. Integrated Lower Cretaceous biostratigraphy of the Bersek Quarry, Gerecse Mountains, Transdanubian Range, Hungary. *Cretaceous Research* 30: 78–92.
- Fülöp, J. 1960. Unterkreide-Bildungen (Berrias-Apt) des Bakony-Gebirges. *Geologica Hungarica* 13: 1–194.
- Gabb, W.N. 1869. *Palaeontology of California. Volume 2, Cretaceous and Tertiary Fossils.* 396 pp. Geological Survey of California, Sherman and Company, Philadelphia.
- Gill, T. 1871. Arrangement of the families of mollusks. Smithsonian Miscellaneous Collections 227: 1–49.
- Haug, E. 1887. Die geologischen Verhältnisse der Neocomablagerungen der Puezalpe bei Corvara in Südtirol. Jahrbuch der Kaiserlich-Königlichen Geologischen Reichsanstalt 37: 245–280.
- Haug, E. 1889. Beitrag zur Kenntniss der oberneocomen Ammonitenfauna der Puezalpe bei Corvara (Südtirol). Beiträge zur Paläontologie und Geologie Österreich-Ungarns 7 (3): 193–229.
- Hoernes, R. 1876. Neocomfundorte in der Gegend von Ampezzo und Enneberg in Südtirol. Verhandlungen der kaiserlich-königlichen Geologischen Reichs-Anstalt 7: 140–141.
- Hölder, H. 1956. Über Anomalien an jurassischen Ammoniten. Paläontologische Zeitschrift 30: 95–107.
- Hyatt, A. 1894. Phylogeny of an acquired characteristic. American Philosophical Society Proceedings 32: 349–647.
- Hyatt, A. 1900. Cephalopoda. In: K.A. von Zittel (ed.), Textbook of Palaeontology, English ed., translated by Eastman C.R., 502–592. Macmilla, London.
- Kakabadze, M.V. [Kakabadze, M.V.] 1981. Ancyloceratids of the south of the USSR and their stratigraphical significance [in Russian]. *Trudy Geologičeskogo Instituta Imeni A.I. Dzanelidze, Akademiâ Nauk Gruzinskoj SSR* 71: 1–195.
- Kazansky, P.A. [Kazancky, P.A.] 1914. Description d'une collection des Céphalopodes des terrains Crétacées du Daghestan [in Russian]. Izvestiâ Tomskogo Technologičeskoga Instituta 32 (4): 127 pp.
- Kilian, W. 1910. 1. Das bathyale Palaeocretacicum im südöstlichen Frankreich In: F. Frech (ed.), Das Mesozoikum. Lethaea geognostica, pt. 2, Das Mesozoicum. Vol. 3, Kreide. Sect. 1, 169–288. Schweizerbart'schen Verlagsbuchhandlung, Stutgart.
- Kilian, W. and Reboul, P. 1915. Contribution à l'étude des faunes paléocrétacées du Sud-Est de la France. I. La faune de l'Aptien inférieur des environs de Montelimar (Drôme). II. Sur quelques ammonites de L'Hauterivien de la Bégude (La Bégüe) [Basses-Alpes]. Mémoires pour servir à l'Explication de la Carte Géologique détaillée de la France 14: 1–288.
- Klein, J., Busnardo, R., Company, M., Delanoy, G., Kakabadze, M., Reboulet, St., Ropolo, P., Vašíček, Z., and Vermeulen J. 2007. Bochiantitoidea, Protancyloceratoidea, Ancyloceratoidea, Ptychoceratoidea. *In:* W. Riegraf (ed.), *Fossilium Catalogus I: Animalia 144*, 1–381. Blackhuys Publishers, Leiden.
- Kotetishvili, E.V. 1983. Sur la paléozoogéographie des bassins eocrétacés du Caucase. Zitteliana 10: 375–386.
- Lory, P. and Sayn, G. 1895. Sur la constitution du système Crétacé aux environs de Chatillon-en-Diois. Annales de l'Université de Grenoble 8: 9–26.
- Lukeneder, A. 2008. The ecological significance of solitary coral and bivalve

epibionts on Lower Cretaceous (Valanginian–Aptian) ammonoids from the Italian Dolomites. *Acta Geologica Polonica* 58: 425–436.

- Lukeneder, A. 2010. Lithostratigraphic definition and stratotype for the Puez Formation: formalisation of the Lower Cretaceous in the Dolomites (S. Tyrol, Italy). *Austrian Journal of Earth Sciences* 103: 138–158.
- Lukeneder, A. 2011. The Biancone and Rosso Ammonitico facies of the northern Trento Plateau (Dolomites, Southern Alps, Italy). *Annalen des Naturhistorischen Museum Wien* 113A: 9–33.
- Lukeneder, A. 2012a. Computed 3D visualisation of an extinct cephalopod using computer tomographs. *Computers & Geosciences* 45: 68–74.
- Lukeneder A. 2012b. New biostratigraphic data of an Upper Hauterivian– Upper Barremian ammonite assemblage from the Dolomites (Southern Alps, Italy). *Cretaceous Research* 35: 1–21.
- Lukeneder, A. and Aspmair, C. 2006. Stratigraphic implication of a new Lower Cretaceous ammonoid fauna from the Puez area (Valanginaian– Aptian, Dolomites, Southern Alps, Italy). *Geo.Alp* 3: 55–91.
- Nagy, I.Z. 1967. Unterkretazische Cephalopoden aus dem Gerecse-Gebirge I. Annales Historico-Naturales Musei Nationalis Hungarici. Pars Mineralogica et Palaeontologica 59: 53–79.
- Obata, I. and Matsukawa, M. 2007. Barremian–Aptian (Early Cretaceous) ammonoids from Choshi Group, Honshu (Japan). Cretaceous Research 28: 363–391.
- Paquier, V.L. 1900. Recherches géologiques dans le Diois et les Baronnies orientales. Travaux du Laboratoire de Géologie de la Faculté des Sciences de l'Université de Grenoble 5: 373–806.
- Pozzi, E. 1993. Die Fossilien der Dolomiten. 176 pp. Tappeiner Verlag, Lana.
- Reboulet, S., Klein, J., Barragan, R., Company, M., Gonzalez-Arreola, C., Lukeneder, A., Raissossadat, S.N., Sandoval, J., Szives, O., Tavera, J.M., Vašiček, Z., and Vermeulen, J. 2009. Report on the 3rd international Meeting on the IUGS Lower Cretaceous Ammonite Working Group, the "Kilian Group" (Vienna, Austria, 15th April 2008). *Cretaceous Research* 30 (2): 496–592.
- Reithofer, O. 1928. Geologie der Puezgruppe (Südtiroler Dolomiten). Jahrbuch der Geologischen Bundeanstalt 78: 257–326.
- Rodighiero, A. 1919. Il sistema Cretaceo del Veneto Occidentale compreso fra l'Adige e il Piave con speziale reguardo al Neocomiano dei Sette Comuni. *Palaeontographica Italica* 25: 39–125.
- Rouchadze, J. 1933. Les ammonites aptiennes de la Géorgia occidentale. Bulletin de l'Institut de géologique de Géorgie 1: 165–273.
- Sarasin, C. and Schöndelmayer, C. 1902. Étude monographique des ammonites du Crétacique inférieur de Chantel-Saint-Denis. Mémoires de la Société Géologique de France 29: 95–195.
- Sarkar, S.S. 1954. Some new genera of uncoiled ammonites from Lower Cretaceous. *Science and Culture* 19 (12): 618–620.
- Sarkar, S.S. 1955. Révision des ammonites déroulées du Crétecé inférieur du Sud-Est de la France. Mémoires de la Société Géologique de France 72: 1–176.
- Scotese, C.R. 2001. Atlas of Earth History. Paleomap Project. 52 pp. Arlington Publishers, Texas.
- Simionescu, I. 1898. *Studii Geologice și Paleontologice din Carpații Sudici*. 133 pp. Academia Româna, Publicatiunile Fundului Vasilie Adamachi II, Bucharest.
- Simionescu, I. 1900. Synopsis des ammonites néocomiennes [Infravalanginien (Berriasien)–Aptien (incl.)]. *Travaux du Laboratoire de Géologie de Grenoble* (1899–1900) 5 (1): 109–143.
- Spath, L.F. 1924. On ammonites of the Speeton Clay and the subdivision of the Neocomian. *Geological Magazine* 61: 73–89.
- Stampfli, G. and Mosar, J. 1999. The making and becoming of Apulia. In: 3<sup>rd</sup> Workshop on Alpine Geology. Mémoires Science Géologie (University of Padova), Special volume 51 (1): 151–154.
- Stampfli, G.M., Borel, G.D., Marchant, R., and Mosar, J. 2002. Western Alps geological constraints on western Tethyan reconstructions. *In*: G. Rosenbaum and G.S. Lister (eds.), Reconstruction of the evolution of the Alpine-Himalayan Orogen. *Journal of Virtual Explorer* 8: 77–106.
- Stöhr, D. 1993. Die Ammoniten der Kreide von La Stua (Dolomiten, Norditalien). 173 pp. Unpublished Ph.D. thesis, University of Giessen, Giessen.
- Stöhr, D. 1994. Ammonoidea aus Schwarzschiefern von La Stua (Norditalien, Provinz Belluno). Giessener Geolgische Schriften 51: 291–311.

- Thomel, G. 1965. Contribution à la connaissance des Céphalopodes crétacés du Sud-Est de la France. Note sur les ammonites déroulées du Crétacé inférieur vocontien. *Mémoires Société Géologique de France. Nouvelle Série* 43 (2), Mémoire 101 (1964): 1–80.
- Thomel, G., Delanoy, G., and Autran, G. 1987 Valeur taxinomique, position stratigraphique et relations phylétiques des genres d'Ammonoidea: *Acrioceras* Hyatt, 1900 et *Aspinoceras* Anderson, 1938 au cours des temps hauteriviens, ainsi que de leurs derives. *Comptes Rendus de la Académie des Sciences de Paris (série 2)* 305: 215–219.
- Thomel, G., Coulet, P., Davoux, J., Delanoy, G., Führ, F., Führ, M., Gazel, L., Giverso, J., Jullian, H., Jullian, V., Levan, G., Mascarelli, E., Maurel, L., Poupon, A., and Salomon, M. 1990. Le genre Acrioceras Hyatt, 1900 dans le Barrémien du Sud-Est de la France. Annales du Muséum d'Histoire Naturelle de Nice 7: 20–151.
- Uhlig, V. 1883. Die Cephalopodenfauna der Wernsdorferschichten. Denkschriften der Österreichischen Akademie der Wissenschaften, mathemathisch-naturwissenschaftliche Klasse 46: 127–290.
- Uhlig, V. 1887. Ueber neocome Fossilien von Gardenazza in Südtirol nebst einem Anhang über das Neocom von Ischl. *Jahrbuch der kaiserlich-königlichen geologischen Reichsanstalt* 37 (1): 69–108.
- Vašíček, Z. 1972. Ammonoidea of the Těšín-Hradiště Formation (Lower Cretaceous) in the Moravskoslezské Beskydy Mts. Rozpravy Ústredního Ústavu Geologického 38: 1–104.
- Vašíček Z. 2006. A remarkable assemblage of Early Barremian ammonites in the Central Western Carpathians (Butkov, Quarry, Slovakia). Acta Geologica Polonica 56: 421–440.
- Vermeulen, J. 1995. Nouvelle biozonation du Barrémien basée sur la famille des Pulchellidae (Ammonoidea). Géologie Alpine 71: 199–211.
- Vermeulen, J. 1996. Le Barremien de Saint-Martin, Escragnolles (Alpes-Maritimes, France). *Riviéra Scientifique* 22: 53–76.
- Vermeulen, J. 1998. Nouvelle biostratigraphie ammonitique du Barrémien (pro parte) du Sud-Est de la France. Comptes Rendus de l'Academie des Sciences, Paris 327: 619–624.
- Vermeulen, J. 2004. Vers une nouvelle classification à fondement phylogénétique des ammonites hétéromorphes du Crétacé inférieur méditerranéen. Le cas des Crioceratitidae Gill, 1871, nom. correct Wright, 1952, des Emericiceratidae fam. nov et des Acrioceratidae fam. Nov. (Ancylocerataceae Gill, 1871). *Riviera Scientifique* 88: 69–92.
- Vermeulen, J. 2005. Boundaries, ammonite fauna and main subdivisions of the stratotype of the Barremian. *In*: T. Adatte, A., Thierry, A. Godet, A. Arnaud-Vanneau, H. Arnaud, and K. Foellmi (eds.), The Hauterivian– Lower Aptian Sequence Stratigraphy from Jura Platform to Vocontian Basin: A Multidisciplinary Approach. *Géologie Alpine, série spéciale «colloques et excursions»* 7: 147–173.
- Vermeulen, J. 2006. Nouvelle classification à fondement phylogénétique des ammonites hétéromorphes du Crétacé inférieur méditerranéen. Annales du Muséum d'Histoire Naturelle de Nice 21: 137–178.
- Vermeulen, J. 2009. Ammonites nouvelles (Ancyloceratoidea) du Barrémien inférieur du sud-est de la France. *Riviéra Scientifique* 92: 49–64.
- Vermeulen, J. and Bulot, L. 2007. Données nouvelles sur la classification phylogénétique des ammonites du Crétacé inférieur. Annales du Muséum d'Histoire naturelle de Nice 22: 101–112.
- Vermeulen, J. and Klein, J. 2006. Endemocerataceae: Pulchelliidae. In: W. Riegraf (ed.), Fossilium Catalogus I: Animalia 141: 187–278. Blackhuys Publishers, Leiden.
- Wachendorf, H., Bettenstaedt, F., and Ernst, G. 1967. Zur Unterkreide-Stratigraphie von Süd-Mozambique. *Neues Jahrbuch für Geologie* und Paläontologie, Abhandlungen 129: 273–303.
- Wiedmann, J. 1966. Stammesgeschichte und System der posttriadischen Ammonoideen, ein Überblick (2. Teil). Neues Jahrbuch für Geologie und Paläontologie, Abhandlungen 127: 13–81.
- Wright, C.W., Calloman, J.H., and Howarth, M.K. 1996. Treatise on Invertebrate Paleontology, Part L, Mollusca 4, Revised (Cretaceous Ammonoidea). 362 pp. Geological Society of America, Boulder and University of Kansas Press, Lawrence.
- Zittel, K.A. von 1884. Cephalopoda. In: K.A. Zittel (ed.), Handbuch der Paläontologie, Band 1, Abteilung 2, Lieferung 3., 329–522. R. Oldenburg, Munich.