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Authors: Saether, Kristian P., Little, Crispin T.S., and Campbell, Kathleen A.

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A new fossil provannid gastropod from Miocene hydrocarbon seep deposits, East Coast Basin, North Island, New Zealand

KRISTIAN P. SAETHER, CRISPIN T.S. LITTLE, and KATHLEEN A. CAMPBELL

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*Provanna marshalli*sp. nov. is described from Early to Middle Miocene−age fossil hydrocarbon seep localities in the East Coast Basin, North Island, New Zealand, adding to 18 modern and three fossil species of the genus described. Modern species are well represented at hydrothermal vent sites as well as at hydrocarbon seeps and on other organic substrates in the deep sea, including sunken wood and whale falls. Described fossil *Provanna* species have been almost exclusively re− ported from hydrocarbon seep deposits, with a few reports of suspected fossil specimens of the genus from other chemosynthetic environments such as sunken wood and large vertebrate (whale and plesiosaurid) carcasses, and the old− est occurrences are dated to the Middle Cenomanian (early Late Cretaceous). The New Zealand fossil species is the most variable species of the genus described to date, and its shell microstructure is reported and found to be comparable to the fossil species *Provanna antiqua* and some modern species of the genus.

Key words: Mollusca, Gastropoda, Provannidae, *Provanna*, hydrocarbon seeps, Miocene, East Coast Basin, New Zealand.

Kristian P. Saether [k.saether@auckland.ac.nz] and Kathleen A. Campbell [ka.campbell@auckland.ac.nz], School of Environment, Faculty of Science, University of Auckland, Private Bag 92019, Auckland Mail Centre, Auckland 1142, New Zealand;

Crispin T.S. Little [c.little@see.leeds.ac.uk], School of Earth and Environment, University of Leeds, Leeds, LS2 9JT, United Kingdom.

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Introduction

Modern submarine hydrothermal vents and hydrocarbon seeps supporting highly productive chemosynthesis−based faunal communities, quite distinct ecologically from the sur− rounding sea floor, have become a well known global phe− nomenon since their discovery in the late 20th century (Lons− dale 1977; Paull et al. 1984; Van Dover 2000; Levin 2005; Suess 2010). These modern "extreme environments" have allowed recognition of vent−seep deposits in the geological record (Campbell 2006) that had previously been enigmatic (e.g., Gilbert and Gulliver 1894; Stanton 1895; Van Winkle 1919; Tanaka 1959; Danner 1966; Moroni 1966; Ager 1986). Establishing the composition and palaeoecological structure of fossil vent−seep communities contributes to better under− standing of evolutionary trends in deep−sea faunas, with applications for palaeobiogeographical reconstruction and modelling.

Provanna Dall, 1918 is one of the more widely repre− sented genera associated with hydrothermal vent and hydrocarbon seep environments, with three fossil and 18 extant species described, as well as several undescribed

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modern species (Anders Warén, personal communication 2009). Modern *Provanna* species are common at vent sites, as well as at hydrocarbon seeps, and it is typical for one or two species to occur in any one chemosynthesis−based com− munity (Warén and Bouchet 1993). Their main mode of nu− trition involves grazing on filamentous bacteria, especially amongst mytilid beds, vestimentiferan tube worm bushes, and upon bacterial mats (e.g., Levin and Michener 2002; Sahling et al. 2002; Levin 2005; MacAvoy et al. 2005; Olu et al. 2009). Gut contents including crustacean fragments, polychaete bristles, sponge spicules, and the tests of plank− tonic organisms such as diatoms and radiolarians have been observed in some species, e.g., *Provanna admetoides* Warén and Ponder, 1991 and *Provanna laevis* Warén and Ponder, 1991, suggesting they also feed on detritus (Warén and Bouchet 1986; Warén and Ponder 1991). At least one species, *Provanna variabilis* Warén and Bouchet, 1986, may harbour symbiotic bacteria (Bergquist et al. 2007). Species of *Provanna* are not thought to have many natural predators within the communities in which they live (Berg− quist et al. 2007), but shell remains have been found in the gut contents of opportunistic octopuses that frequent hydro−

Fig. 1. Overview of relevant geology of the East Coast Basin, North Island, New Zealand, showing locations of the fossil seep deposits from which type specimens and a possible specimen of *Provanna marshalli* sp. nov. were collected. 1. Puketawa. 2. Rocky Knob. 3. Ugly Hill. CP, Cape Palliser; HSZ, Hikurangi subduction zone; M, Marlborough; TVZ, Taupo Volcanic Zone. Also shown are some of the modern offshore seeps on the Hikurangi margin, in− cluding Ritchie Ridge, from which several modern *Provanna* specimens have been reported. Inset shows the Hikurangi subduction zone as part of the transpressive boundary between the Indo−Australian and Pacific Plates. Figure modified from Campbell et al. (2008).

thermal vents (Voight 2000, 2008). Kiel (2006) also re− ported a healed shell injury in a single specimen of *Pro− vanna antiqua* Squires, 1995 from the Oligocene Lincoln Creek Formation, north−western USA, that may have re− sulted from a predatory crab attack. Species of *Provanna* are often reported as endemic to vent−seep environments

(e.g., Hessler and Lonsdale 1991; Warén and Bouchet 1993; Carney 1994; Yamamoto et al. 1999; Fujikura et al. 2002; Cordes et al. 2009), but some species have been re− ported from other organic substrates such as sunken drift− wood (Warén and Bouchet 2001, 2009) and whale falls (Smith et al. 2002; Smith and Baco 2003). These more ephemeral deep−sea habitats may act as stepping stones for dispersal and colonisation between the geographically re− stricted but more stable and long−lived chemosynthetic en− vironments at vents and seeps (e.g., Distel et al. 2000).

Institutional abbreviations.—GNS, GNS Science, Lower Hutt, New Zealand; GSNZ, Geological Society of New Zea− land, Lower Hutt, New Zealand; LEMAS, Leeds Electron Microscopy and Spectroscopy Centre, Institute for Materials Research, University of Leeds, Leeds, United Kingdom; NMNZ, Te Papa Tongarewa, Wellington, New Zealand; UOA, Geology, School of Environment, University of Auckland, Auckland, New Zealand.

Other abbreviations.—AU (preceding five−digit number), collection number for material deposited in the palaeontol− ogy collection, UOA; D, shell diameter; d, apertural diame− ter; G (preceding four−digit number), gastropod specimen number, palaeontology collection, UOA; GS (preceding five−digit number), fossil locality number, GNS; H, shell height; h, apertural height; TM (preceding four−digit num− ber), type specimen number, palaeontology collection, GNS; Y16/f (preceding four−digit number), fossil locality number, registered in archival New Zealand Fossil Record File main− tained jointly by GNS and GSNZ.

Materials and methods

29 fossil gastropod specimens were collected during a num− ber of field trips to both the northern and southern Hawke's Bay Miocene seep localities (Fig. 1) over several years (1997–2009), adding to material obtained from the bulk palaeontology collections at GNS. Shell images were obtained via scanning electron microscopy at UOA and LEMAS. Specimens described and figured herein are de− posited in the palaeontology collections of UOA. Modern comparative material is housed at NMNZ. Measurements were taken of all of the specimens with vernier callipers to within an error of 0.1 mm.

Geological setting

At least 16 geographically isolated ancient hydrocarbon seep deposits are present along 300 km of the uplifted East Coast Basin forearc, eastern North Island (Campbell et al. 2008). They occur in two distinct regions in the Raukumara Penin− sula and near Dannevirke, to the north and south of the Hawke's Bay region, respectively, where they occur as dis− crete pods or lenses of authigenic carbonate enclosed within thick siliciclastic mudstone deposits (Fig. 2).

The East Coast Basin formed as a result of regional tectonism that has been ongoing since the Early Miocene (Ballance 1976). Extending roughly 650 km north to south, the East Coast Basin varies in width from 60 to 110 km,

Fig. 2. Generalised stratigraphy of the East Coast Basin, showing the Mio− cene−age seep carbonates that occur as isolated lenses within the mudstones of the Tolaga Group. The organic−rich, underlying Waipawa Black Shale is thought to be one of the likely source rocks. Abbreviations: Bx, Bexhaven Limestone; Ih, Ihungia Limestone; To, Tolaga Group; Wb, Weber Forma− tion; Wn, Wanstead Formation; Wp, Waipawa Black Shale. Figure modi− fied from Campbell et al. (2008).

bounded to the west by NNE–SSW trending axial mountain ranges and extending offshore to its eastern boundary along the similarly trending Hikurangi Trough (Fig. 1). It repre− sents an exhumed forearc, its sediments having been gener− ated during a period of oblique convergence along the Hiku− rangi subduction zone throughout the Cainozoic (Ballance 1976; Barnes et al. 2002), during the onset of intense defor− mation, the appearance of andesitic volcanism, and abrupt changes in sedimentation rate and character (Rait et al. 1991). Today, the Hikurangi margin constitutes the southern ex− tremity of the Tonga−Kermadec−Hikurangi system, an east− facing subduction zone oriented roughly parallel to the East Coast of North Island, along which the Pacific Plate is mov− ing north−westward underneath the overriding Australian Plate (e.g., Ansell and Bannister 1996).

We report morphologically highly variable specimens of the genus *Provanna* from two of the northern region New Zealand hydrocarbon seep deposits (Fig. 1): Puketawa and Rocky Knob (late Early to Middle Miocene age). All but a single specimen comes from Rocky Knob, mostly from within a group of closely scattered float boulders originat− ing from the main scarp of the site (Fig. 3A, B), found in as− sociation with worm tubes and bathymodioline and large lucinid bivalves (Fig. 3C, D), as well as small, unidentified ?columbellid gastropods and acmaeid limpets. A poorly preserved, possibly juvenile gastropod specimen from the southern region Ugly Hill seep (Early Miocene age) may also belong to the same species, but we have not used it here in the species description. The ages of these deposits are ap− proximate due to stratigraphic correlation limitations of the limestones within the monotonous, voluminous mudstone sequence, which is structurally complex with major, longactive, low−angle normal faults predominating in some ar− eas, especially in the north (Campbell et al. 2008).

Fig. 3. Field photographs showing the Middle Miocene Rocky Knob type locality and collection location of the assemblages that yielded most of the speci− mens of *Provanna marshalli* sp. nov. **A**. Main scarp of the deposit, white asterisk marks the collection location. **B**. Slab in which an assemblage (arrowed) was found, located a few metres from a smaller block (not pictured) that yielded the majority of the specimens; locations of photos C and D marked by squares. **C**, **D**. Other taxa (arrowed) found in association with the assemblage within the boulder. **C**. Worm tubes (wide arrow) and lucinid bivalves (thin arrows). **D**. Bathymodioline bivalves. Photographs taken by KAC (1997: fig. 3A) and KPS (2008: fig. 3B–D).

Systematic palaeontology

Phylum Mollusca Linnaeus, 1758

Class Gastropoda Cuvier, 1797

Order Caenogastropoda Cox, 1960

Superfamily Abyssochrysoidea Tomlin, 1927

Discussion.—Placement of the provannids within this super− family follows Kaim et al. (2008a, 2009). Bouchet et al. (2005) did not recognise this superfamily and placed the provannids alongside the Abyssochrysidae Tomlin, 1927

within the paraphyletic "zygopleuroid group" amongst other caenogastropods of uncertain taxonomic position.

Family Provannidae Warén and Ponder, 1991 Genus *Provanna* Dall, 1918

Type species: ?*Trichotropis (Provanna) lomana* Dall, 1918. Recent; US Pacific coast; by monotypy.

Provanna marshalli sp. nov.

Figs. 4, 5.

Etymology: Named for Bruce A. Marshall, malacologist and collections manager of Mollusca at NMNZ, who has published extensively on New SAETHER ET AL.—MIOCENE PROVANNID GASTROPOD FROM NEW ZEALAND 511

Fig. 4. Variation in provannid gastropod *Provanna marshalli* sp. nov. as seen from the apertural (upper two rows) and apical (lower two rows) views of specimens from the Middle Miocene Rocky Knob type locality. **A**. Paratype G7105. **B**. Paratype G7110. **C**. Paratype G7108. **D**. Paratype G7106. **E**. Para− type G7109. **F**. Paratype G7107. **G**. Paratype G7111. **H**. Holotype G7103. **I**. Paratype G7126.

Zealand molluscan taxonomy, and has been involved in the identifica− tion of the modern New Zealand seep faunas since the first official seep collections were made in 1996.

Type material: Holotype: G7103 (Fig. 4H), well preserved shell lacking protoconch and earliest whorls. Paratypes: one moderately preserved specimen, TM8704, from Puketawa (Y16/f0580); nine moderately to

well preserved specimens from Rocky Knob: Y16/f1027, AU19605, seven specimens (G7105–7109, G7113, G7126); Y16/f1029, AU15834, one specimen (G7110); Y16/f1031, AU19610, one specimen (G7111). All specimens at UOA, TM8704 borrowed from palaeontology collec− tion, GNS.

Type locality and horizon: Rocky Knob (Y16/f1027), northern Hawke's Bay area (38°30.58'S, 177°93.10'E). Fossil hydrocarbon seep deposit of the Bexhaven Limestone Formation, Tolaga Group, Middle Miocene.

Material.—19 poorly to well preserved specimens from Rocky Knob: Y16/f1027, AU19605, 14 specimens (G112, G114–125); Y16/f1028, AU19606, two specimens (G128, one unnumbered specimen); Y16/f1030, AU19609, three specimens (G7104, G7129, one unnumbered specimen). All specimens at UOA.

Dimensions.—See Table 1.

Table 1. Selected measurements of specimens of *Provanna marshalli* sp. nov., this study. Italicised values are approximate. Non−type material is labelled "specimen".

Diagnosis.—Shell medium−sized, ovate, fusiform; sculpture highly variable from nearly smooth to strongly cancellate; body whorl with up to 11 spiral ribs and 30 axial riblets, often forming strong nodes at their intersections, especially along tabulate shoulder; suture inclined at roughly 5° from horizontal; at least three teleoconch whorls present.

Description.—Shell medium−sized (D up to 6.7 mm; H up to ca. 10.8 mm), ovate (D/H \approx 0.60–0.82), fusiform; shell microstructure consists of at least one (outer) simple pris− matic layer (ca. 5–10 μm) above thicker crossed lamellar layer (ca. 30–40 μm); periostracum not preserved. Sculp− ture highly variable, ranging from nearly smooth, with only very weak spiral ribs or very weak axial riblets formed by rugose growth lines, to strongly cancellate with several spi− ral ribs, widely but evenly spaced. At least three teleoconch whorls present, protoconch and very earliest whorls not preserved (see remarks below); whorls rather convex with tabulate shoulder delineated by spiral rib, often strongly noded; suture distinct, inclined at angle roughly 5° from horizontal (shell axis vertical); last whorl moderately glo− bose with distinct neck having up to three spiral ribs, mod− erate to weak; 0–5 spiral ribs on penultimate whorl, moder− ate to weak; on last whorl number of ribs varies from none (or possibly two very weak, one on shoulder, one on neck) to ca. 11, with increased strength towards neck and shoul− der; 0–30 axial riblets on last whorl, less on earlier whorls, which often form strong nodes where crossing spiral ele− ments, especially prominent on tabulate shoulders; strong sculpture can be consistent from aperture through to early whorls, or fade out in middle whorls, being expressed only in antepenultimate (possibly earlier) whorl and in latest shell, but usually with some weak sculpture in between. Ap− erture higher than wide $(d/h = 0.61-0.94$, probably less variable in life before taphonomic alteration), usually just over half height of shell; basal and outer lips often eroded, where intact forming almost circular continuation; basal lip comprises lowest point of shell from apertural perspective, passing to distinct siphonal canal, visible in a few speci− mens where not eroded. Operculum unknown.

Remarks.—A single, variable species is herein recognised based upon the following points: (i) the similarity in general shell shape between the specimens, with variations being ob− served chiefly in ornamentation; (ii) a continuum of ornamental strength, with smooth and strongly cancellate end− member specimens separated by intermediate specimens of incremental ornamental strength that grades from one end to the other; (iii) the frequency of observed morphological vari− ation of a similar nature in other described species of the ge− nus; (iv) the collection of the majority of specimens from a single assemblage within one site. Modern species of this ge− nus are commonly rather simple and variable in shell shape and form, which can make it difficult to compare them across localities, although there is usually some obvious shell mor− phological difference where two or more species are found at a given site (Warén and Bouchet 1993).

Many of the *Provanna marshalli* specimens are well preserved, shown by the retention of original shell micro− structure (Fig. 5), although the protoconchs and very earli− est whorls are never present. It is rare to collect the fully in− tact larval shell of modern species of *Provanna*, because it is thought that the apical whorls of the protoconch are dis− carded after the veliger phase (Kaim et al. 2008a; Warén and Bouchet 2009). Furthermore, the corrosive condition of their environment often leads to loss of further early whorls in life, and secondary growth of a calcareous plug to replace

the protoconch is common (Kaim et al. 2008a; Warén and Bouchet 2009). These environmental factors, along with preservational constraints during fossilisation, make obser− vations of this key taxonomic shell feature in fossil material highly unlikely. Indeed, only once has the protoconch I been reported in fossil material, in two specimens of *Pro− vanna antiqua* (Kiel 2006: fig. 5.1–4). A decollate proto− conch II also was reported in a Late Cretaceous−age juve− nile *Provanna* specimen by Kaim et al. (2008a), where ex− quisite preservation occurred by early diagenetic silicifica− tion. Moreover, of 18 described modern species, only four species have been figured with an intact protoconch I, and these always from juvenile specimens (Warén and Ponder 1991; Warén and Bouchet 1993, 2009; Gustafson and Lutz 1994; Kaim et al. 2008a).

simple prismatic layer.

The smooth specimens of *Provanna marshalli* resemble the species *Provanna chevalieri* Warén and Bouchet, 2009 from off West Africa, *Provanna glabra* Okutani, Tsuchida, and Fujikura, 1992 from Sagami Bay, Japan, and *Provanna laevis* Warén and Ponder, 1991 from the Gulf of California, but none of these species ever develops the strong ornamen− tation seen in other specimens of *P. marshalli*. Two further species may have smooth−shelled individuals: *Provanna antiqua* and *Provanna variabilis*, and strong morphological variation is seen in both species. However, morphological variation in *P. marshalli* is even greater than that seen in *P. antiqua* and *P. variabilis*, with a similar smooth end−member but ranging up to a far greater number of axial riblets. The su− ture in *P. marshalli* is different from *P. antiqua* in its greater inclination from the horizontal (where shell axis is vertical).

Moreover, the shells of *Provanna marshalli* are considerably larger, reaching nearly twice the height of *Provanna antiqua*, are wider but less squat overall, and also with a larger maxi− mum height than *Provanna variabilis*. The new species has similarities to the "Shosanbetsu *Provanna* sp." from a Mio− cene whale fall community in Japan, which was figured and briefly, but not formally, described by Amano and Little (2005). However, the Japanese species has a maximum mea− sured height of 6.1 mm, which is significantly less than *P. marshalli* (ca. 10.8 mm). It also has weaker sculpture that ex− hibits less variability (22–26 axial ribs on the last whorl), and a higher maximum number of spiral cords (16) than is seen in any specimen of *P. marshalli*. *Provanna admetoides* is the only species of the genus to possess more axial riblets (35–45 on last whorl) than the maximum observed in specimens of *P. marshalli*. *Provanna admetoides* differs in this character, as well as in its lesser variability, and far stronger and less nu− merous spiral ribs, which gives it a much more distinctly spi− rally keeled appearance than *P. marshalli*. Furthermore, the reticulations in the ornament of *P. admetoides*, formed by the intermeshing of axial and spiral elements, are elongated axi− ally, as opposed to more or less equilateral in *P. marshalli*. Specimens of *P. marshalli* with stronger, more cancellate sculpture can resemble *Provanna nassariaeformis* Okutani, 1990, but have a larger maximum number of axial and spiral ribs, a less globose shell, and more variability in ornamenta− tion overall than this species. Specimens of *P. marshalli* with intermediate ornamentation often resemble *Provanna lo− mana* (Dall, 1918), but the latter does not display the variation in its sculpture, and spiral ornamentation disappears in whorls earlier than the last whorl, unlike in *P. marshalli*, in which it persists, albeit more weakly, in some specimens.

The shell microstructure of *P. marshalli*, where preser− ved, is shown to be similar to that observed in some modern and fossil specimens of the genus. Kiel (2004) described the microstructure of the shell material of two *P. variabilis* spec− imens from the Juan de Fuca Ridge, which consists of an outer, organic periostracum slightly separated (probably during preparation) from an outer simple prismatic layer, beneath which is a thicker layer with a complex crossed lamellar structure. Another simple prismatic layer forms the innermost portion of the shell, but this layer is variably pres− ent or absent depending upon location on the shell. Kiel (2006) reported a similar shell microstructure in *P. antiqua* and Kaim et al. (2009) also showed this pattern in the shell of a related gastropod, *Hokkaidoconcha hikidai* Kaim, Jenkins, and Warén, 2008a. The periostracum is never preserved in the New Zealand fossil specimens, but there are several places in the shell of one specimen (Fig. 5A) where the outer simple prismatic layer is visible above a thicker crossed lamellar layer. Since the shell microstructure has only been observed in commarginal exposures, it cannot be confirmed whether the nature of the crossed lamellar layer is simple (i.e., two non−vertical dip directions) or complex (i.e., three or more non−vertical dip directions) (cf. Carter 1990). In an− other specimen from this study (Fig. 5B), a distinct simple

prismatic layer is preserved above layers that recrystallised during diagenesis. The boundary is not particularly sharp be− tween the outer simple prismatic and complex crossed lamellar layers in modern specimens (see Kiel 2004: figs. 30–32). In the fossil material, the transition between these layers also is indistinct (Fig. 5A), although there is a rather distinct boundary between the preserved simple prismatic and lower recrystallised layer in one specimen (Fig. 5B). The inner simple prismatic layer, which can form a sharp bound− ary with the complex crossed lamellar layer in modern speci− mens (see Kiel 2004: fig. 30) is never preserved in the fossil specimens. This is what might be expected, where the or− ganic periostracum would have protected the outer simple prismatic layer until it naturally decayed after burial, the in− ner simple prismatic layer afforded no such protection. Weathering of internal shell layers likely occurred on the sea floor or during early burial, leaving the two inner shell layers relatively unscathed where conditions conducive to such preservation prevailed thereafter.

Stratigraphic and geographical range.—Early to Middle Miocene seep carbonates of Hawke's Bay, North Island, New Zealand. Known mostly from the type locality at Rocky Knob (28 specimens), with one confirmed specimen from Puketawa and one possible juvenile specimen from Ugly Hill.

Discussion

Provanna **in the fossil record**.—*Provanna antiqua* was the first formally described fossil species of the genus, from hydrocarbon seep deposits in Washington, USA (Squires 1995), and it has since been extensively reported from seep deposits of this region (e.g., Squires and Goedert 1995; Rigby and Goedert 1996; Peckmann et al. 2002; Goedert and Benham 2003; Kiel 2006; Table 2). Two more fossil *Pro− vanna* species were recently described from Japan, namely *Provanna tappuensis* Kaim, Jenkins, and Warén, 2008a and *Provanna nakagawaensis* Kaim, Jenkins, and Hikida, 2009. There are other reports of the genus from fossil seep deposits where details have been unavailable for species−level identi− fication (Table 2). The oldest formally described species of the genus is from the Middle Cenomanian (upper Late Creta− ceous) "Kanajirisawa" deposit, Yezo Group, northern Japan (Kaim et al. 2008a), although Kiel and Campbell (2005) re− ported gastropod fossils that may belong to *Provanna* from the Valanginian (Early Cretaceous) Crack Canyon Forma− tion, Great Valley Group, California. Of the reports of pro− vannids in the seep fossil record where it has been possible to identify the genus, all but the recently described *Desbruye− resia kanajirisawensis* Kaim, Jenkins, and Hikida, 2009 be− long to *Provanna*. "Provannids" have been reported from Palaeocene seep deposits of California (Schwartz et al. 2003), but these are instead aporrhaids (Kaim et al. 2008a; CTSL, personal observation). There also are several seep

fossil species described from localities in the Antarctic, France, Japan, and the USA that belong to closely related families of the "zygopleuroid group", the majority of which are now included within Hokkaidoconchidae (see Goedert and Kaler 1996; Kiel and Peckmann 2007; Kiel 2008; Kiel et al. 2008; Kaim et al. 2008a; Kaim and Kelly 2009; Kiel et al. 2010). Of the relatively few reports of *Provanna* from fossil deposits other than hydrocarbon seeps, the only verified spe− cies assignment is *Provanna antiqua*, reported from wood fall deposits of the Late Eocene Lincoln Creek Formation by Kiel and Goedert (2006a, b). Undescribed fossils inferred as belonging to *Provanna*, or at least members of the zygo− pleuroid group, also have been reported from Cretaceous to Miocene−age fossil communities associated with hydrother− mal vents (Little et al. 1998, 1999; Little 2002; Little and Vrijenhoek 2003), whale and plesiosaurid carcasses (Amano and Little 2005; Amano et al. 2007; Kaim et al. 2008b), and wood fall deposits (Kiel et al. 2009).

Fossil versus modern New Zealand *Provanna*.—Undes− cribed *Provanna* species have been reported from modern New Zealand seeps (e.g., Lewis and Marshall 1996; Warén and Bouchet 2001), but not by Baco et al. (2010), who made the largest and most recent attempt to characterise the mod− ern NZ seep faunas. Lewis and Marshall (1996) suggested the presence of two distinct modern species, their spp. A and B, based on four specimens trawled from 1100–1200 m at *Calyptogena* Bank, Ritchie Ridge (between 39-26.4'S, 178-23.6'E and 39-23.6'S, 178-24.7'E; Fig. 1). They also reported a more tentatively assigned "?*Provanna* sp. C" from two specimens collected from 960 m depth off south−eastern

South Island at Goomes Hill, Puygesur Bank (46°57.64'S, 165°25.21'E). Each of these specimens is smaller than the specimens of this study. Ongoing studies of the fossil content of New Zealand's East Coast Basin seep deposits suggest that the composition of seep communities from the Hiku− rangi margin has remained relatively stable over the past ca. 20 Ma, with several fossil groups showing a close relation to those found in analogous modern environments (KPS, un− published data). Therefore, it is possible that the modern *Provanna* specimens reported by Lewis and Marshall (1996) represent a single variable modern species, perhaps con− specific with, or a descendent of, *Provanna marshalli*. This hypothesis requires confirmation through analysis of a larger amount of modern New Zealand seep material and complete formal characterisation of the New Zealand Miocene seep fauna, which is outside the scope of this study.

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